

Arctic Basin Hydrography

...& some circulation

Michael Steele

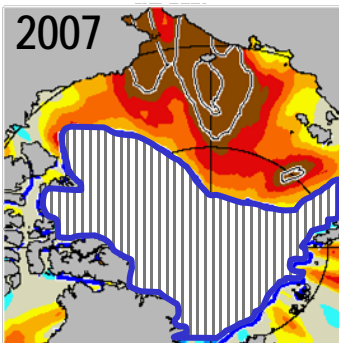
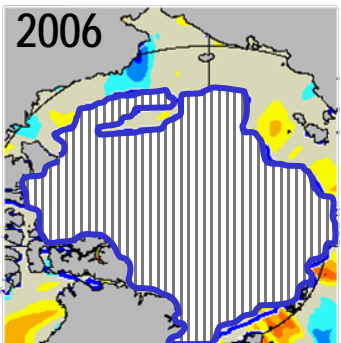
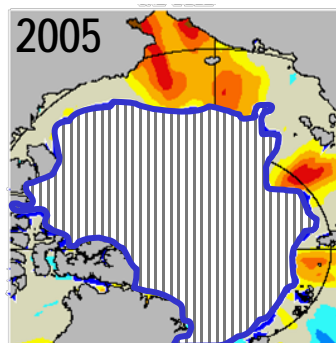
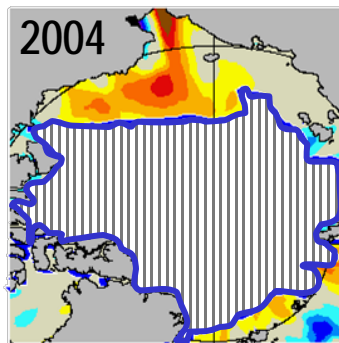
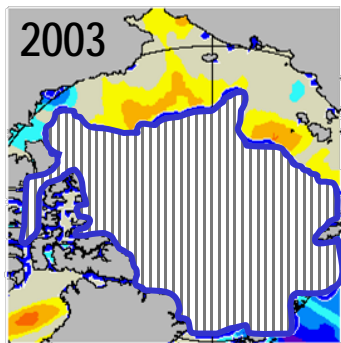
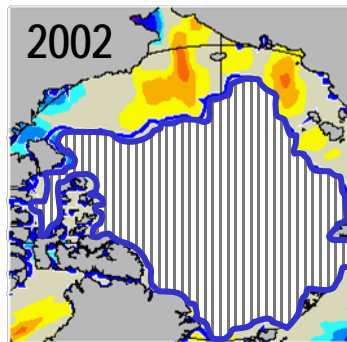
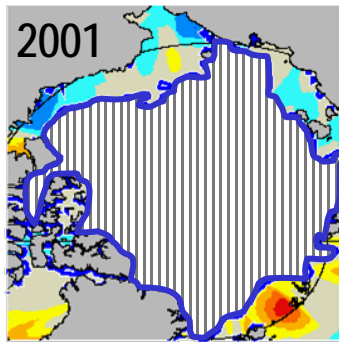
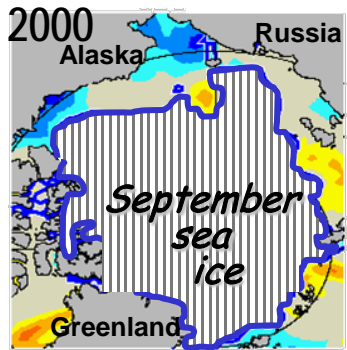
*Polar Science Center
Applied Physics Laboratory
University of Washington*

☞ **Large-scale,**

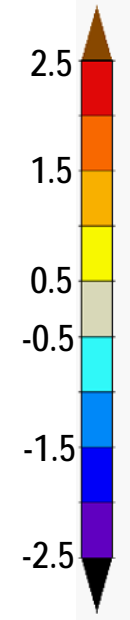
☞ **Pacific-Arctic focus**

- **Warming**
- **Freshening?**

Ice Retreat & Upper Ocean Warming



from Steele et al., GRL 2008



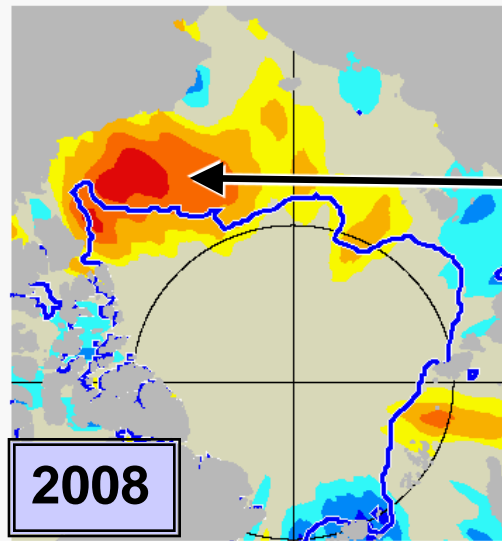
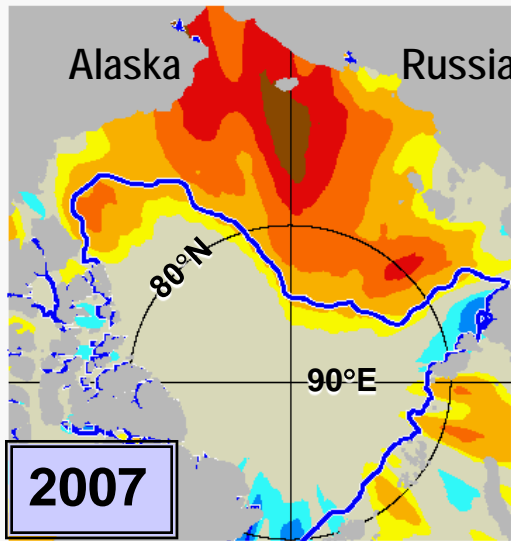
**Anomaly of Summer \equiv JAS
Sea Surface Temperature ($^{\circ}\text{C}$)**
(relative to 1982-2007 mean)

> 4 $^{\circ}\text{C}$ above mean

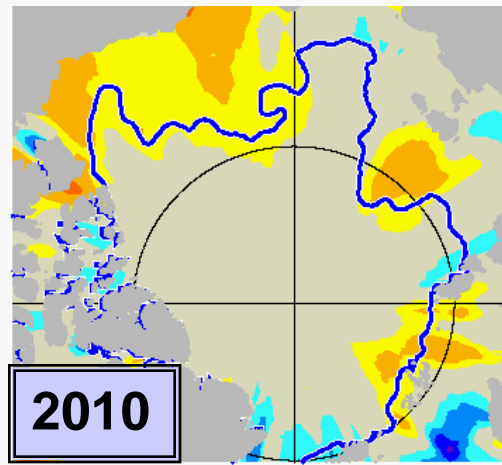
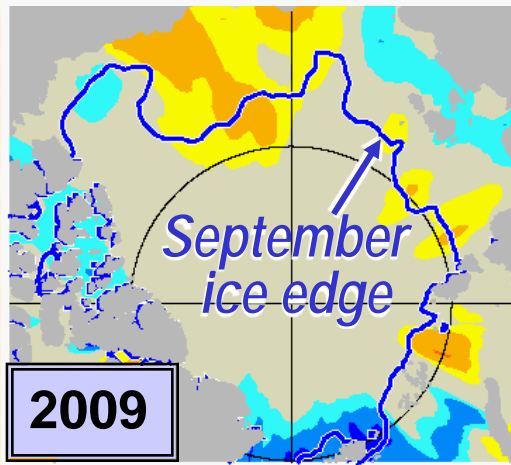
SSTs from monthly
mean AVHRR (Reynolds et al)

Summer SST Anomaly

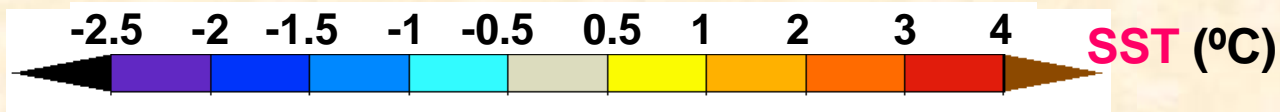
(relative to 1982-2006 mean)



Early ice retreat



Not as warm...



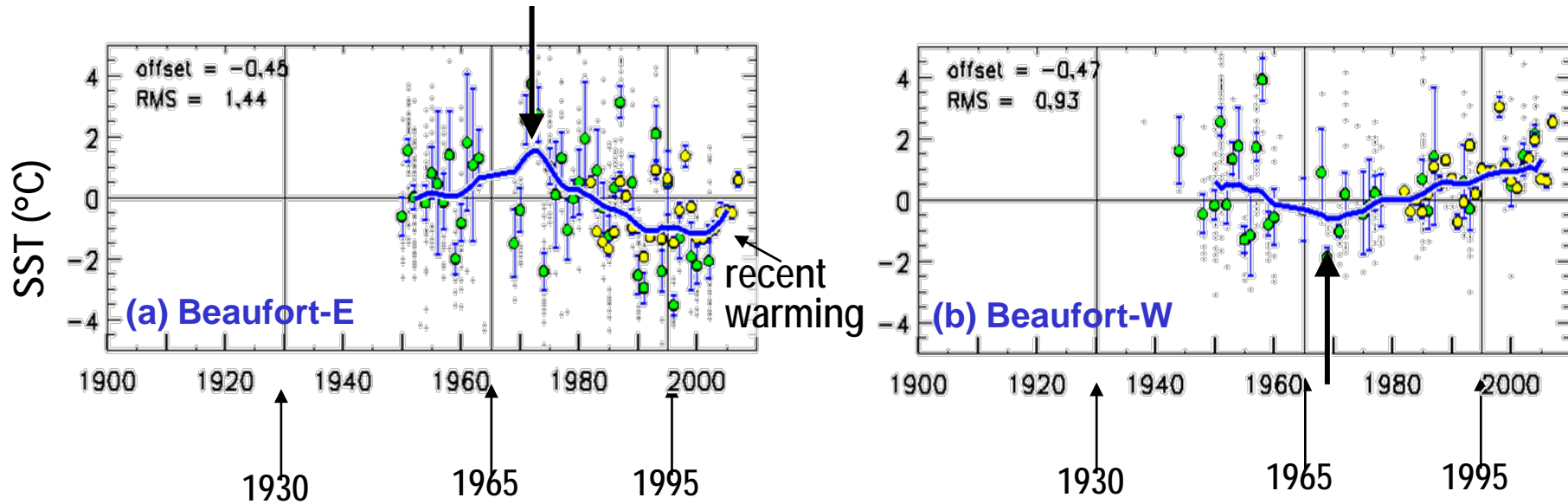
summer SST trends:

regional anomaly time series

- *in situ* WOD'05
- *satellite* Reynolds et al



smoothed



- Natural see-saw?
- Recent pan-arctic warming

Three Questions



1.) What causes upper **ocean warming**?



2.) How much sea **ice melt** is forced by this warm ocean?



3.) What is the **fate** of summer heat?

SUMMER

*(Steele et al. JGR
November 2010)*

FALL/WINTER

*(Steele et al. JGR
submitted)*

Pan-Arctic Ice–Ocean Modeling and Assimilation System (PIOMAS)

Jinlun Zhang, Univ. of WA

- TED (Thickness & Enthalpy Distribution) **sea ice** model
- POP (Parallel Ocean Program) **ocean** model, v. 1.4
- **No assimilation** in this study
- **Atmos forcing**: NCEP
- 30 vertical levels (5 m resolution in upper 30 m)
- 22 km resolution (average); 3 grid pts across Bering St.

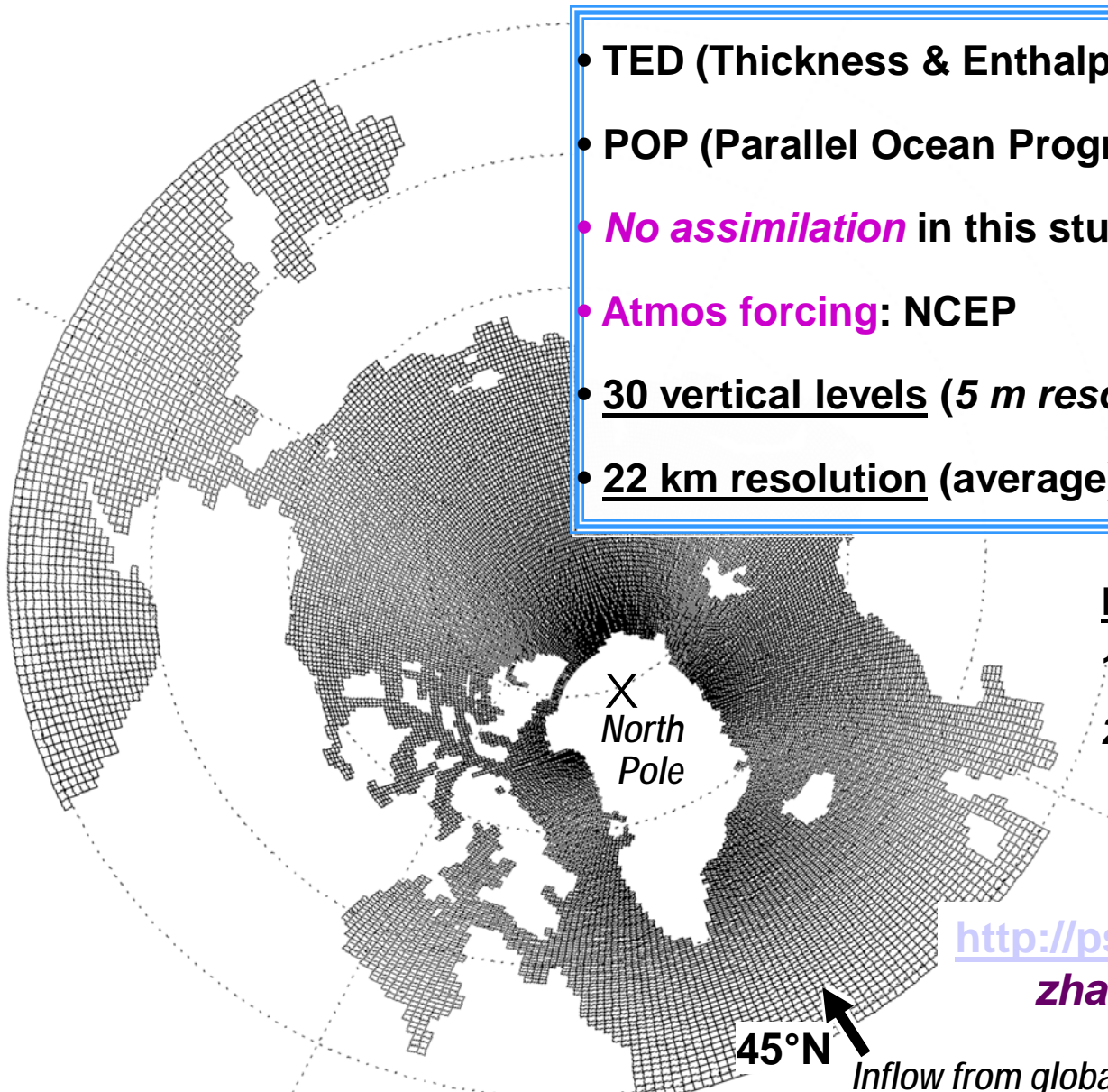
BESTMAS:

10 km resol. in Beaufort/Chukchi;
20 pts across Bering St.

<http://psc.apl.washington.edu/IDAO>
zhang@apl.washington.edu

45°N

Inflow from global model



Summer Upper Ocean Heat Balance

June 1 – Sept 15



F_{surf}

$$\frac{\partial H}{\partial t} =$$

Net air-ocean heat flux

Ice-penetrating solar heat flux

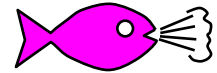


ice

Ice-ocean heat flux

ocean

$$\frac{\partial H}{\partial t}$$



F_{advec}
= lateral heat flux convergence

60 m depth

Lateral diffusive heat flux

Vertical advec/convec heat flux convergence

Vertical diffusive heat flux

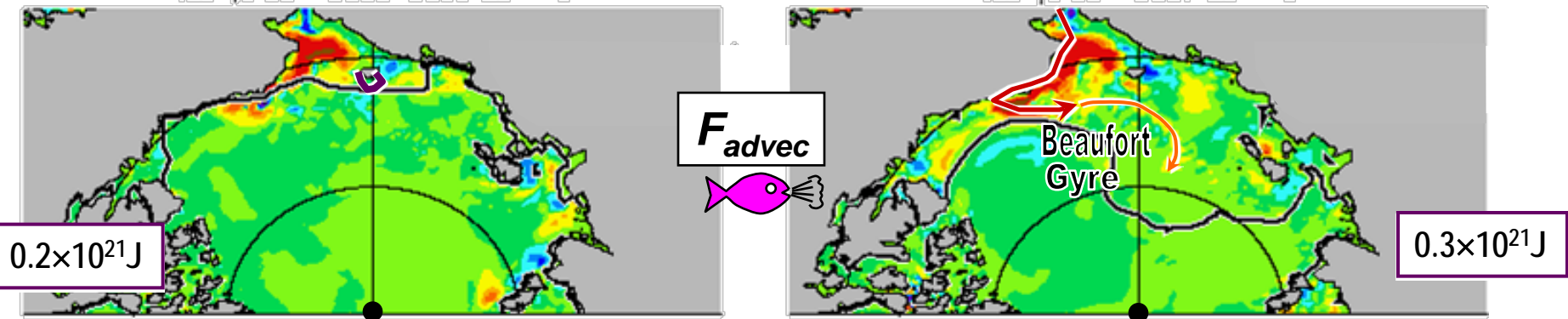
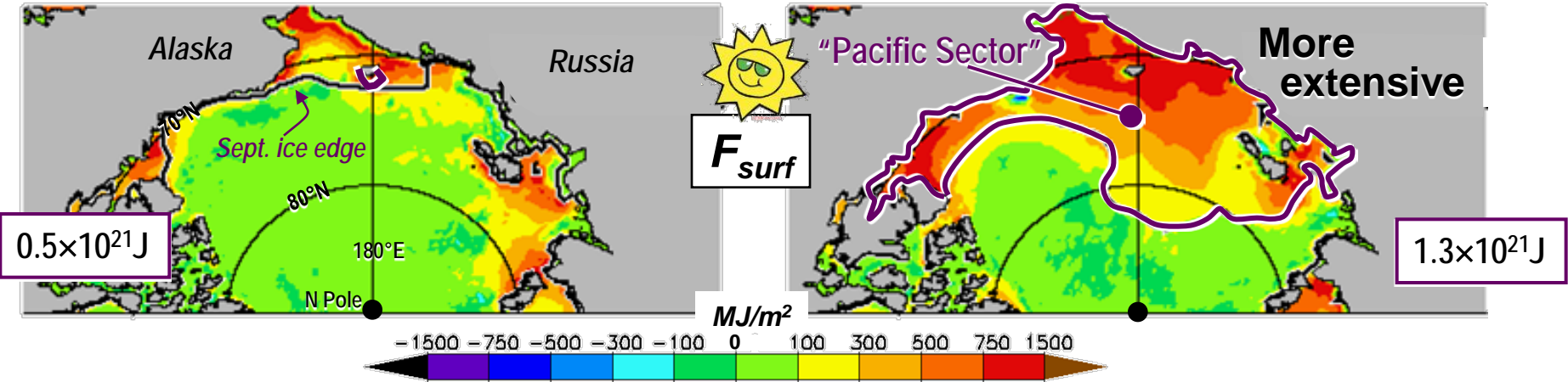
F_o

2000-2006 vs. 2007

Summer Ocean Heating

2000-2006

2007



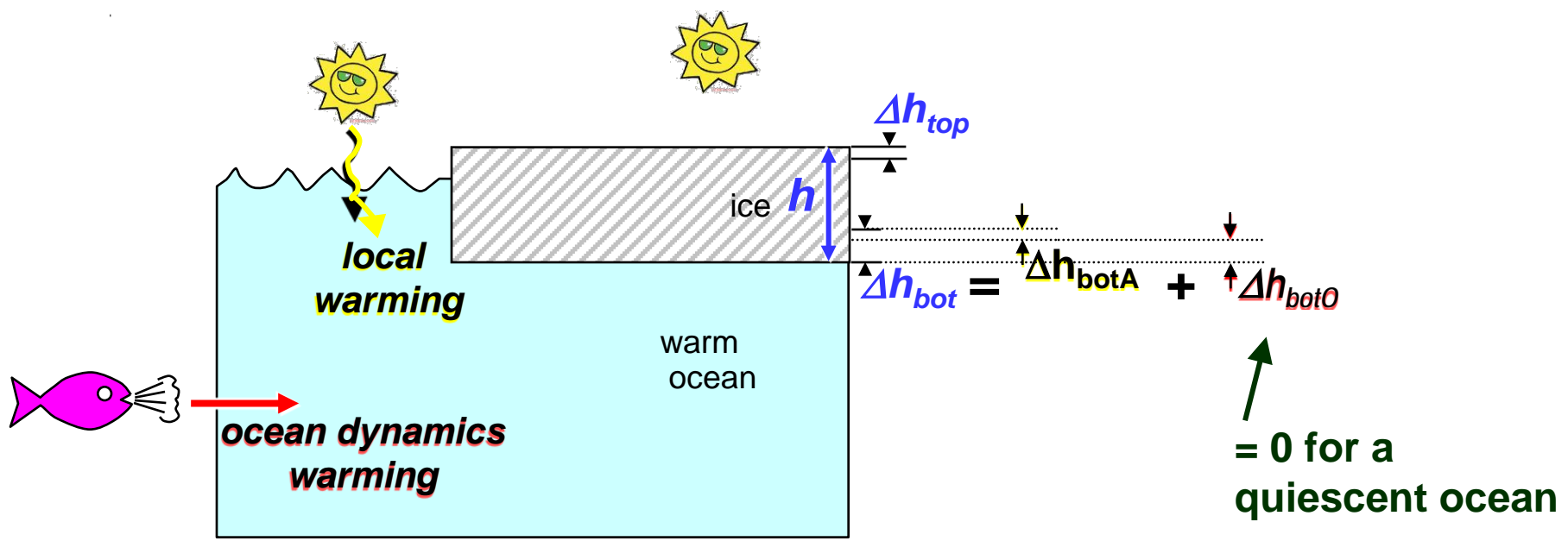
$F_0 < 0.1 \times 10^{21} \text{ J}$

•2007: Alaskan "retroreflection" pathway unchanged, but stronger

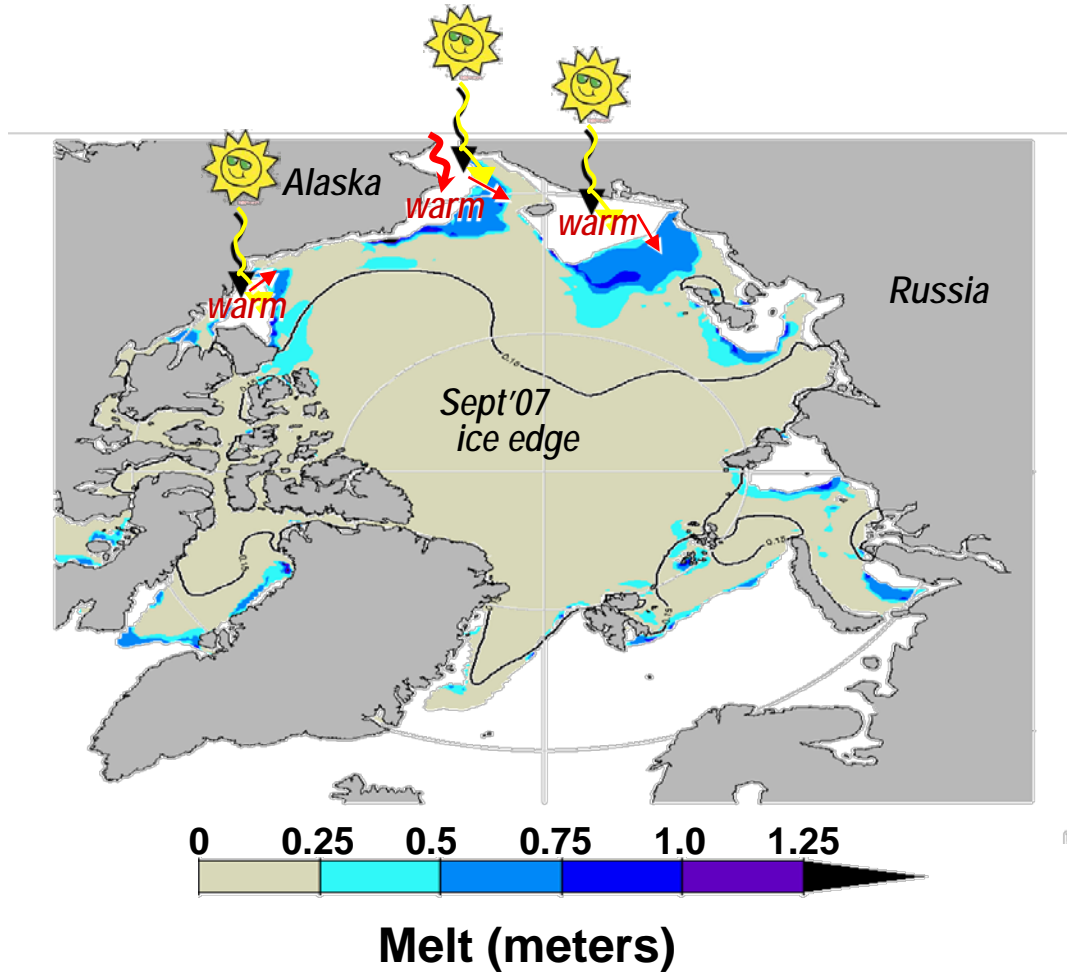
see Shimada et al., GRL 2006

F_{surf} = 70-80% of ocean surface warming

Summer Ice Mass Balance

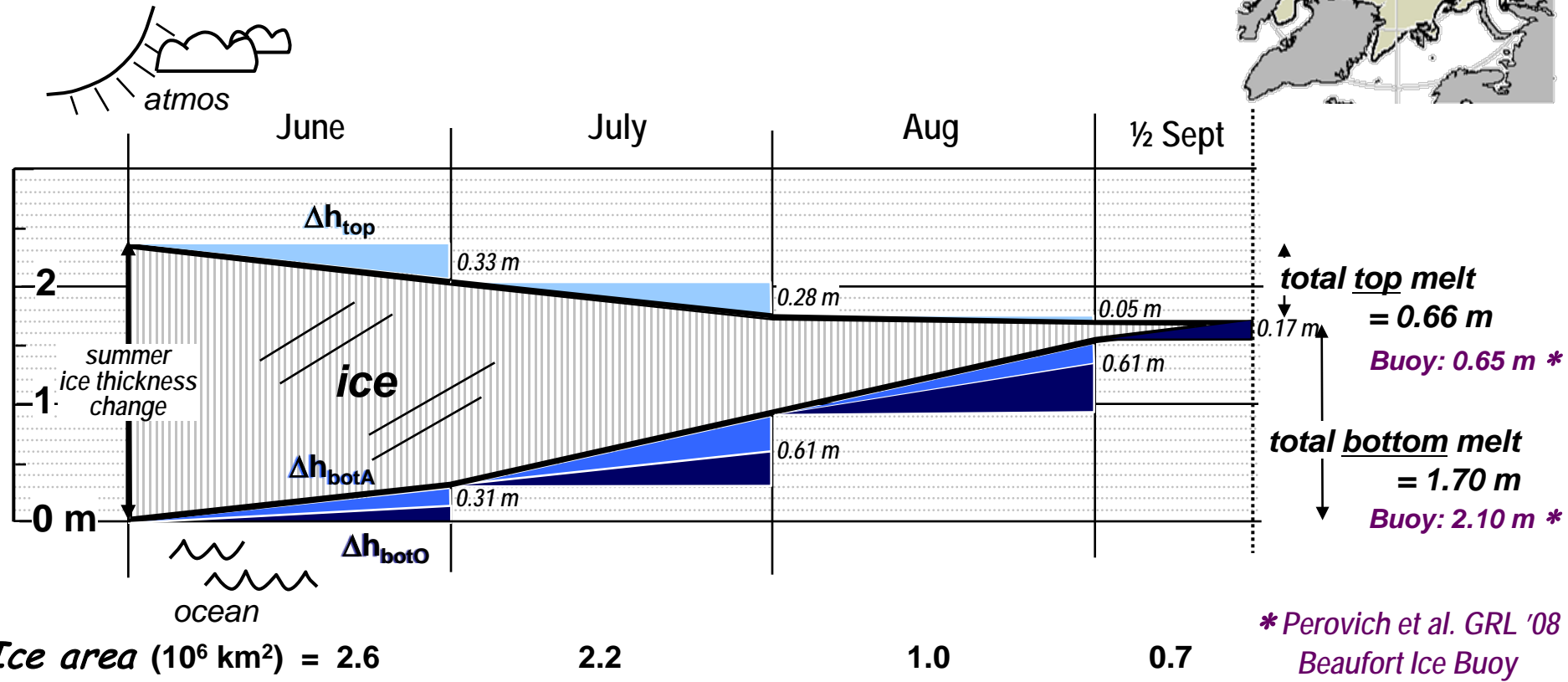


Δh_{bot0} = Ice melt from ocean dynamics
July '07



- **warm** SST
- **strong advec**

Ice thickness Budget: 2007

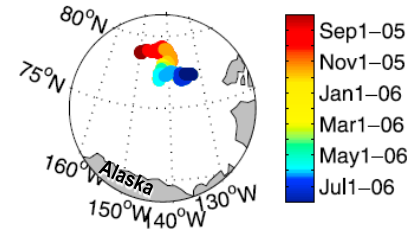
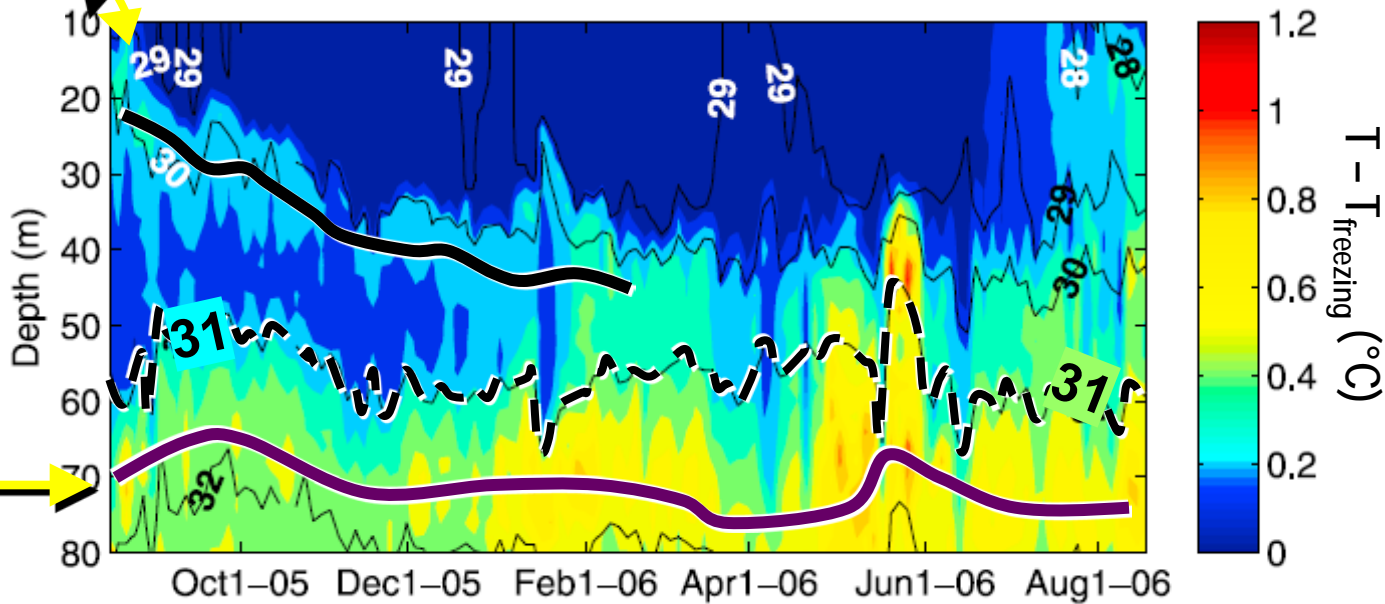


Next Question:

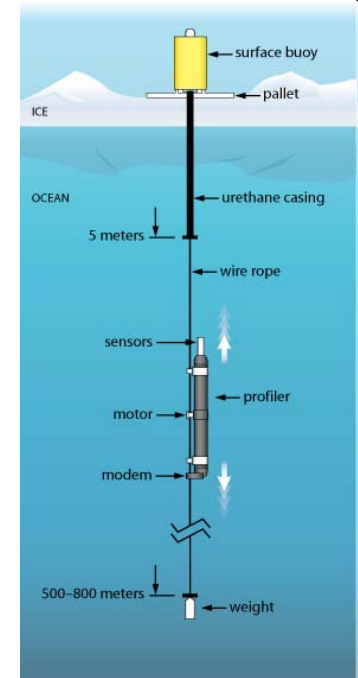
What is the FATE of summer ocean heat?



Jackson et al. JGR 2010



Data:
Ice-Tethered Profiler



Krishfield et al. J. Tech. 2008

contours: Salinity

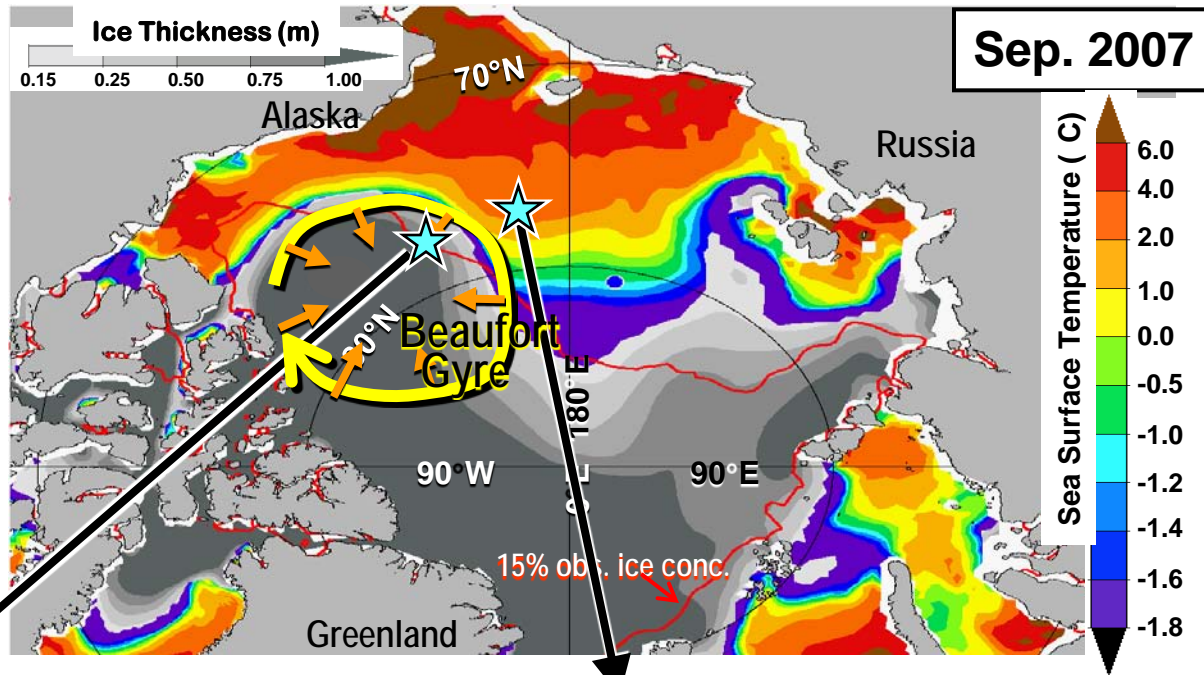
The “Near-Surface Temperature Maximum” (NSTM)
...a LOCAL Tmax layer, as opposed to...

“summer Pacific Water” (sPW)

ITP data from 2007-2008: NSTM survives through the winter!

Data are sparse:
What can we learn about the NSTM from a model?

Steele et al. submitted
 (2010)

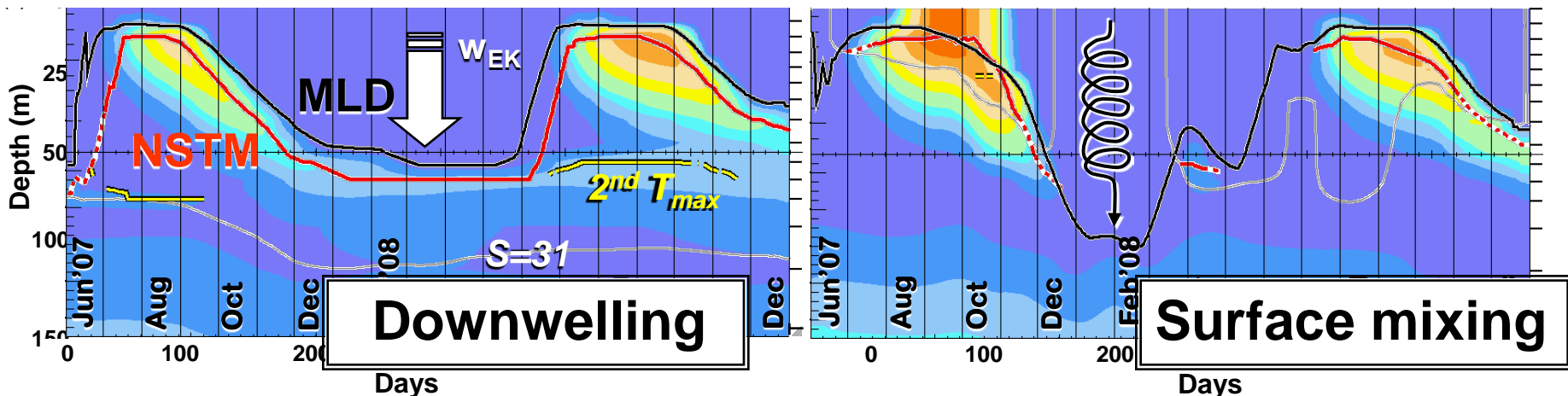


Case study:

**June 2007 –
 December 2008**

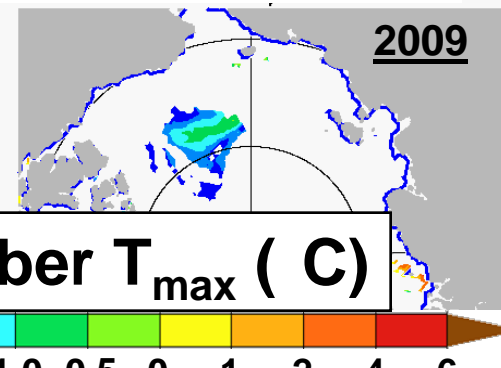
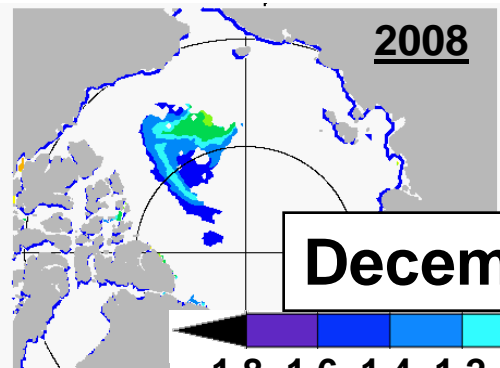
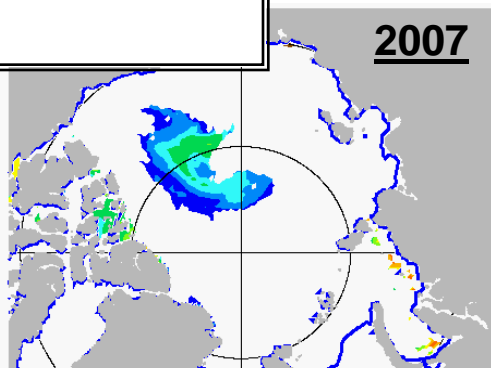
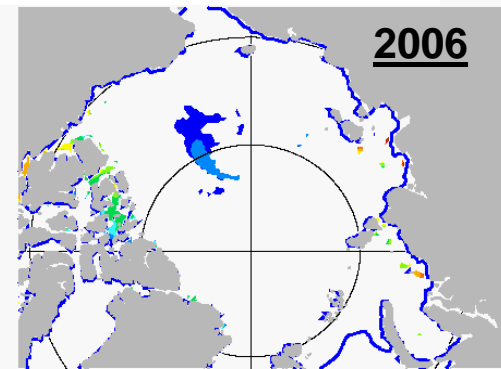
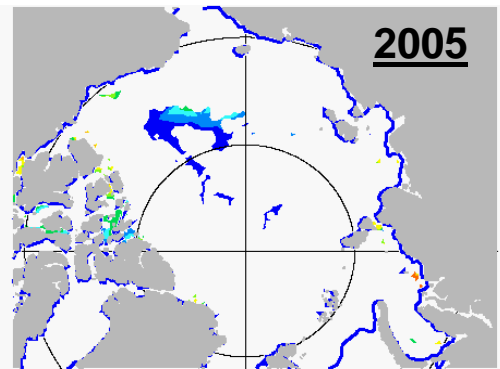
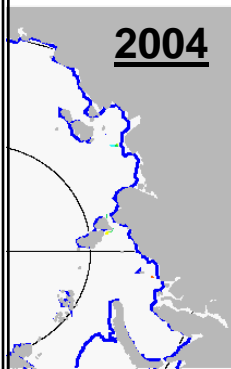
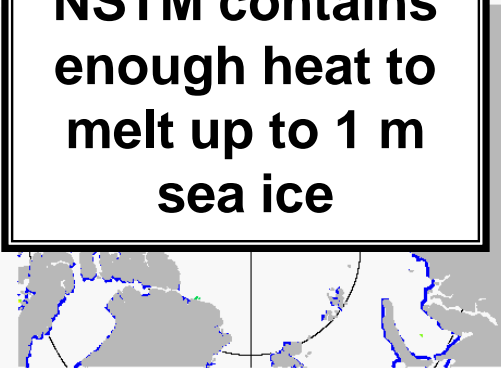
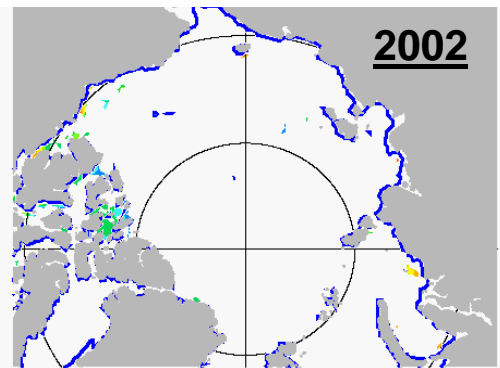
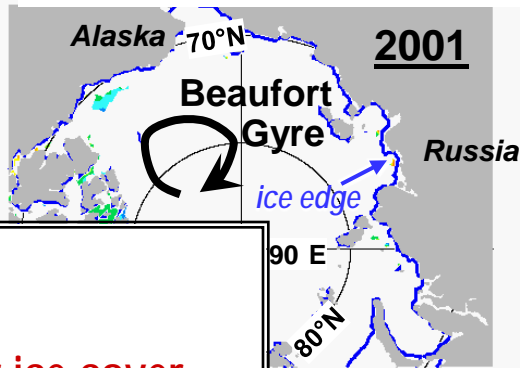
NSTM *survives through winter*

summer NSTM *destroyed by fall convection*



Recent winter survival of the NSTM: WHY?

NSTM contains enough heat to melt up to 1 m sea ice



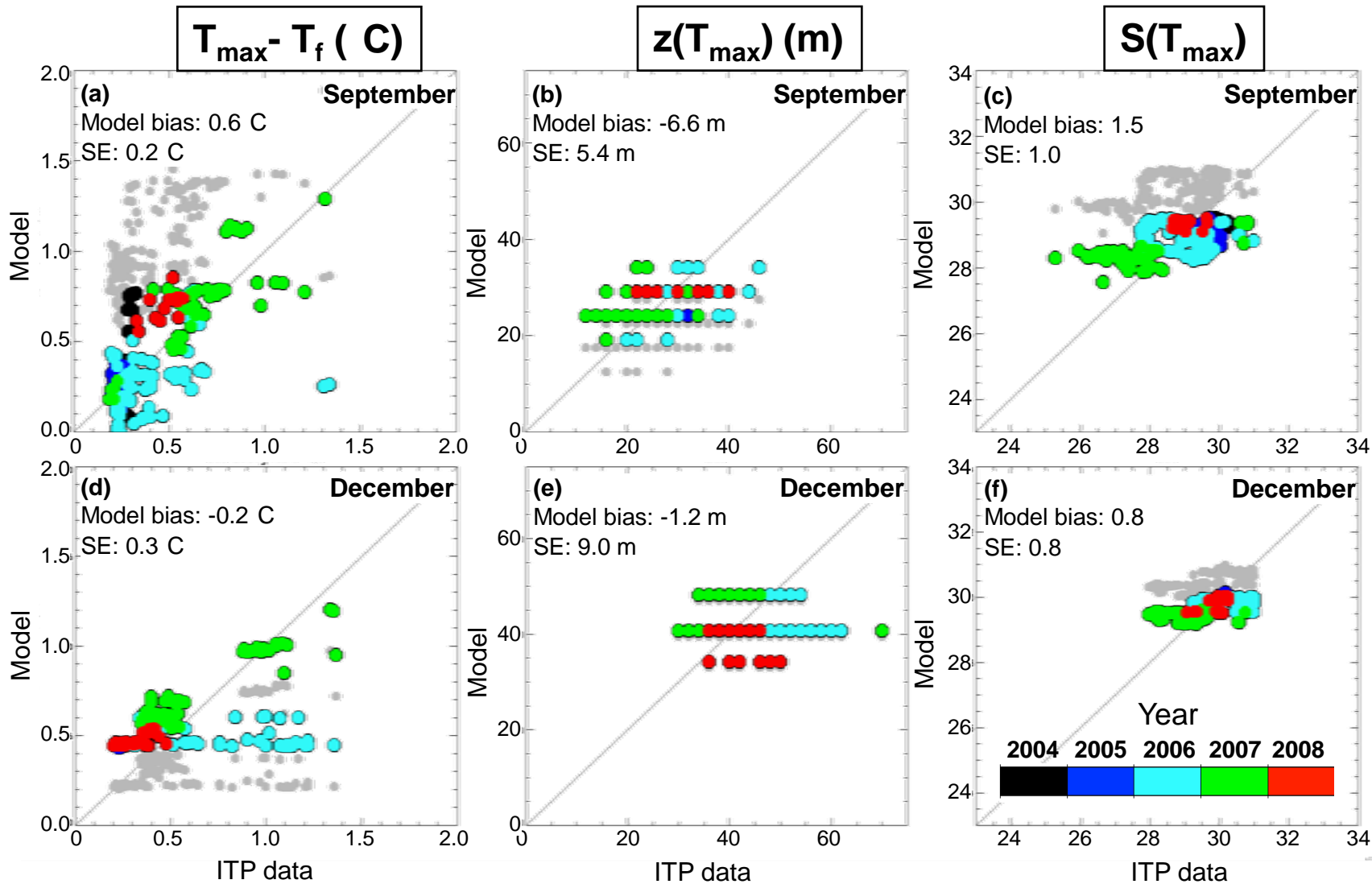
Three factors:

- 1.) Thinner, **looser ice cover** allows more summer heating
- 2.) Increasing Beaufort Gyre **stratification** suppresses surface mixing
- 3.) Increasing Beaufort Gyre **downwelling**

December T_{max} (C)



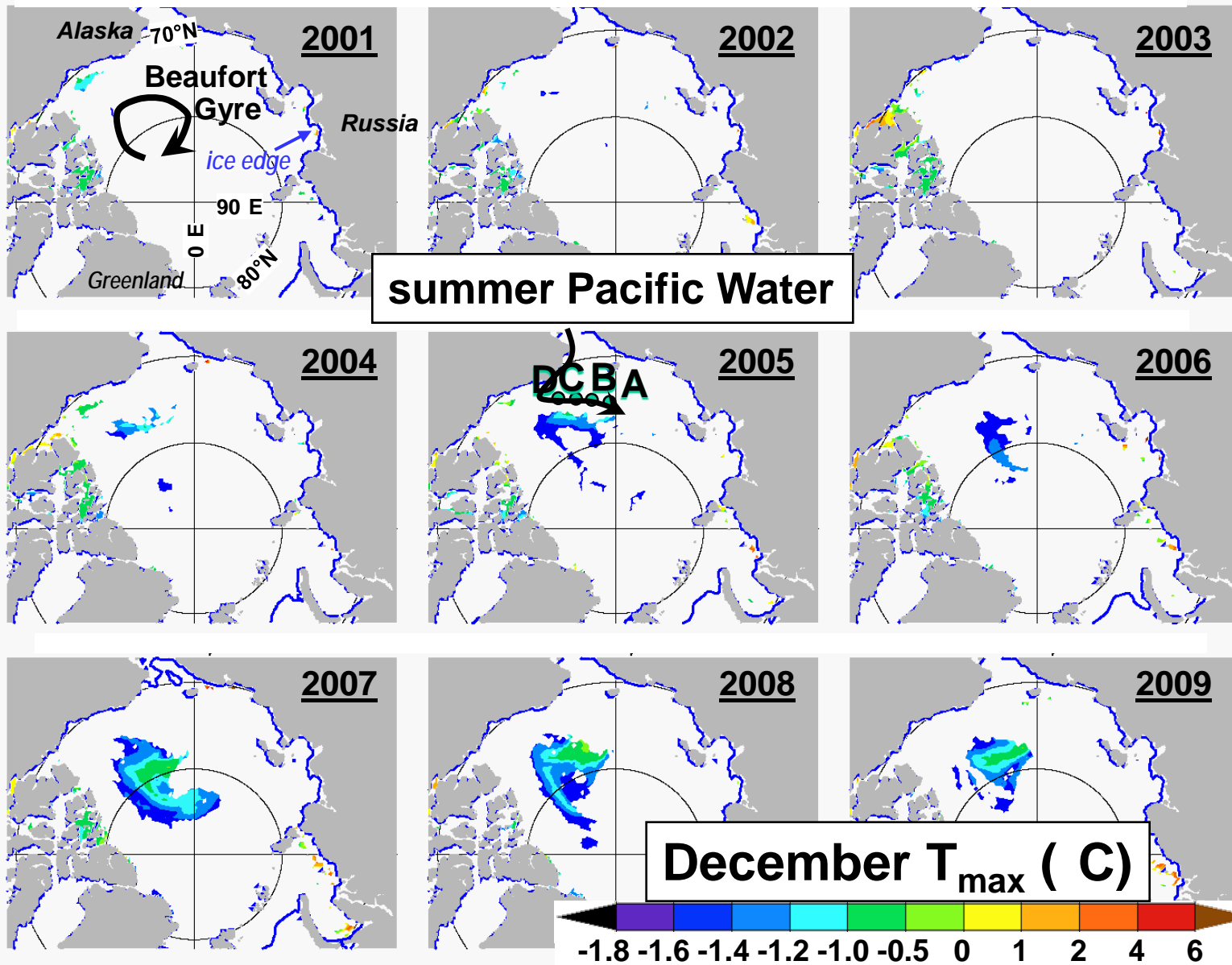
Validation: model vs. ITP NSTM

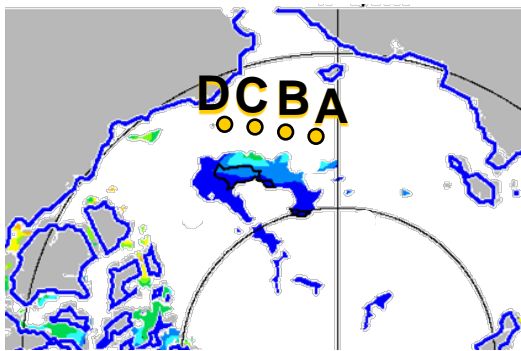


tenths of a °C error Model resolution error

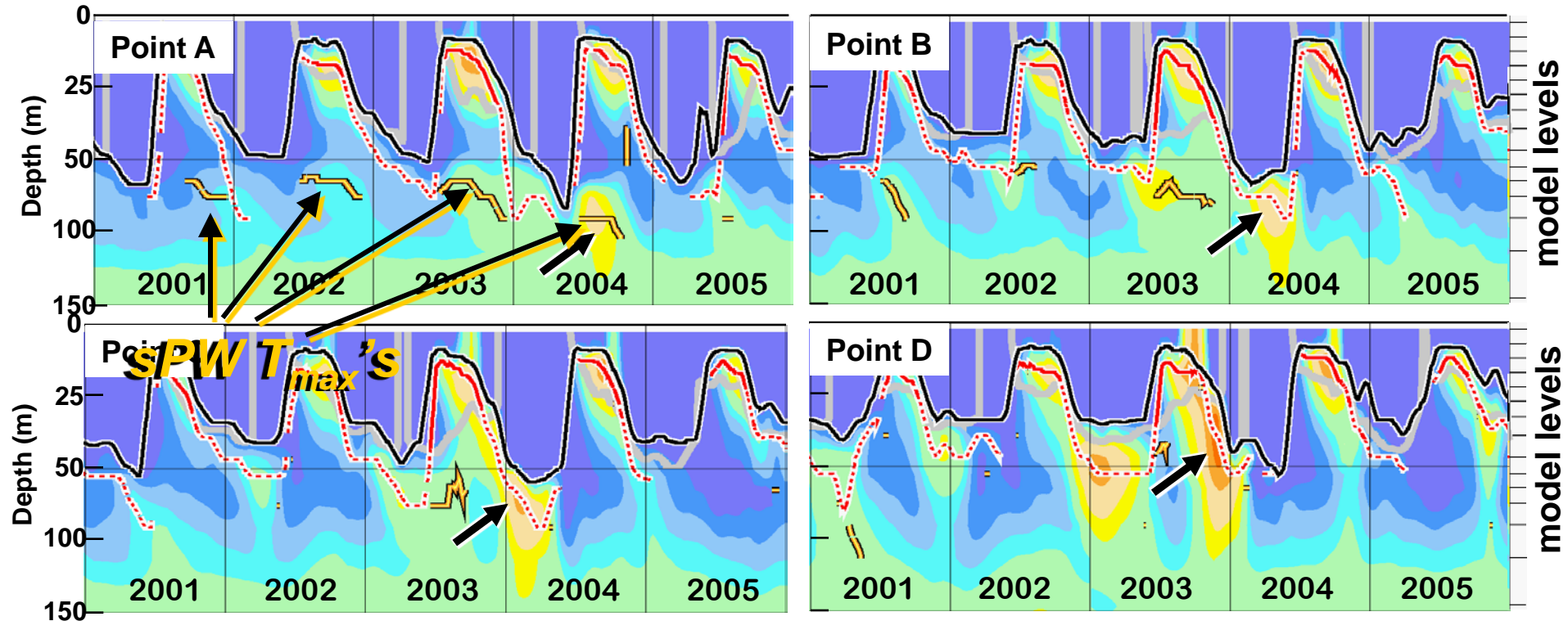
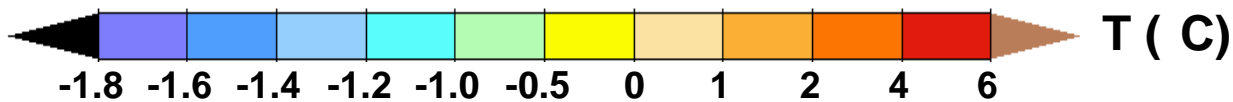
~1 ppt error

Pacific Water



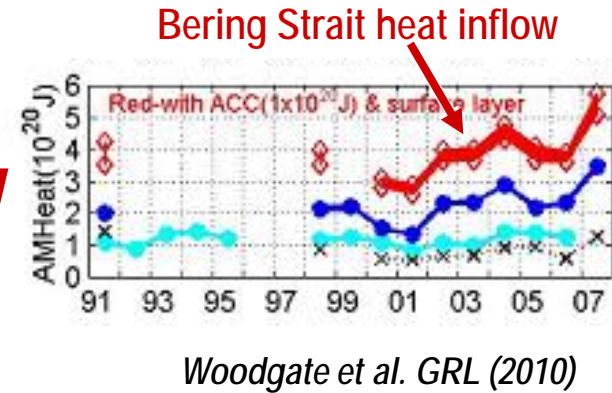
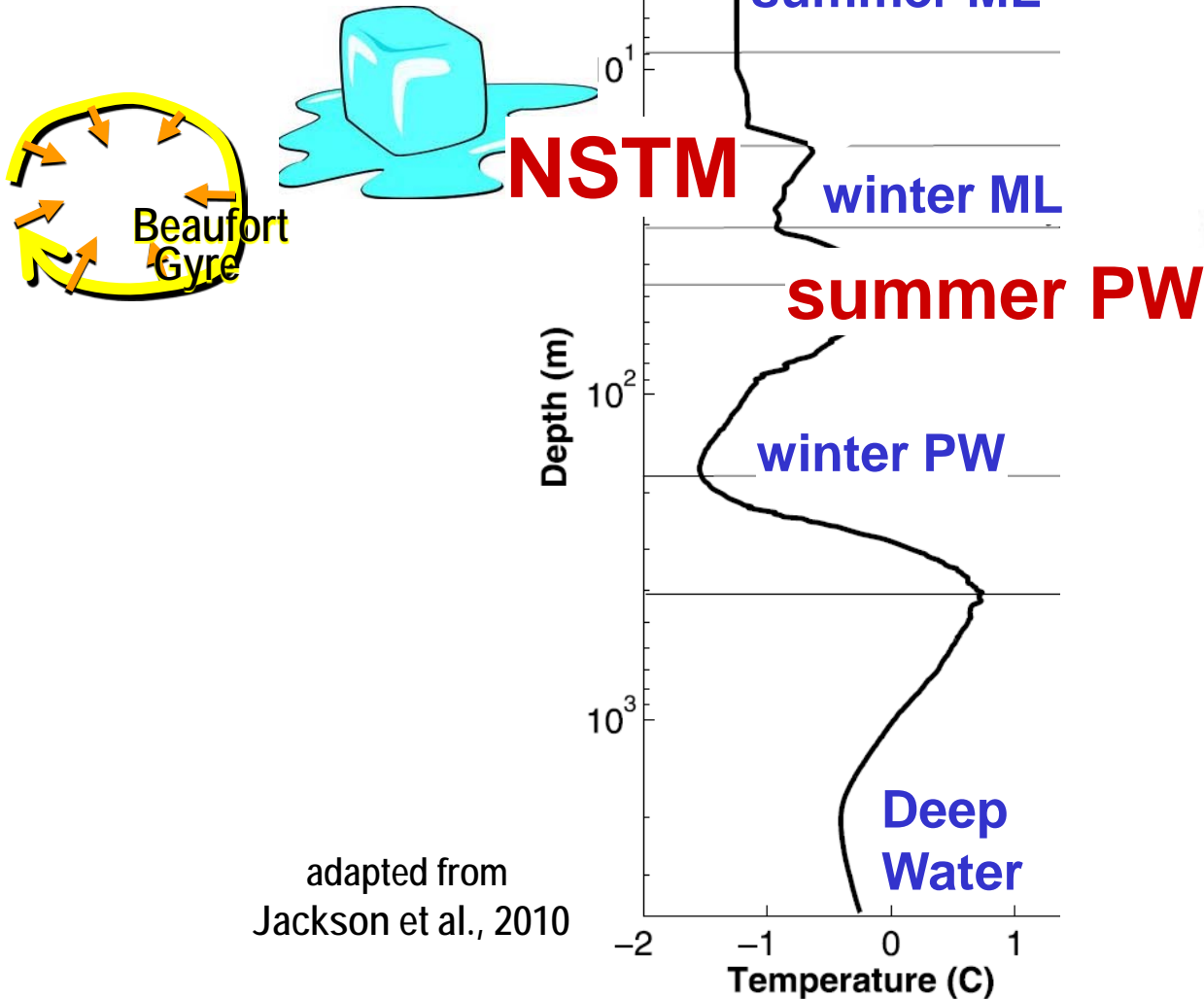


summer Pacific Water in the model: **A sad story**



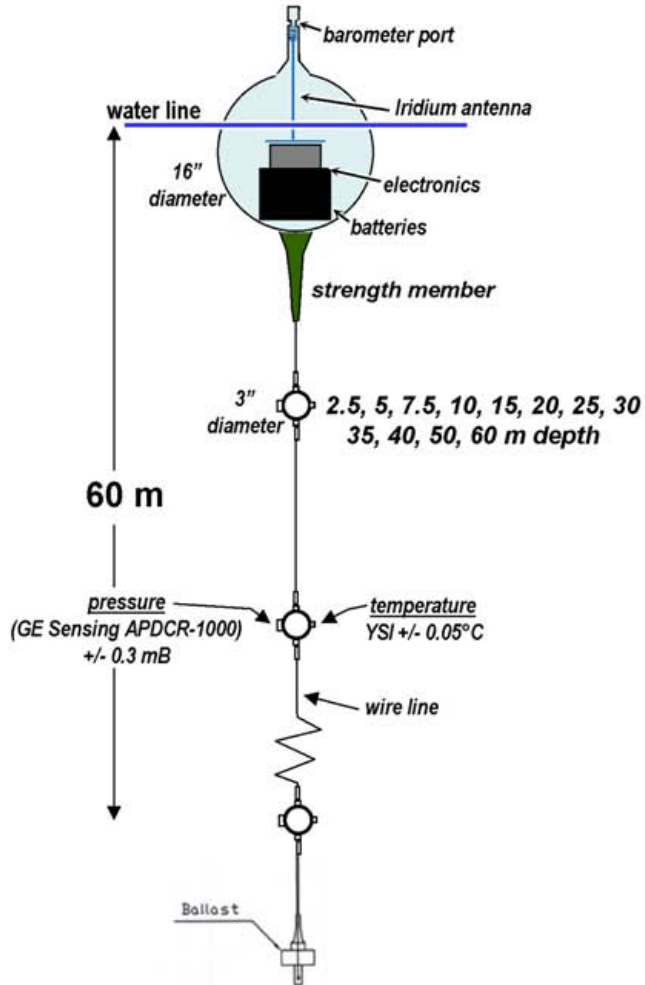
- 1.) sPW enters the Beaufort Gyre near NW Alaska, but...
- 2.) Model's resolution/mixing is not adequate to maintain a sPW layer at 70-100 m depth.

The future?



Measuring Upper Ocean Warming

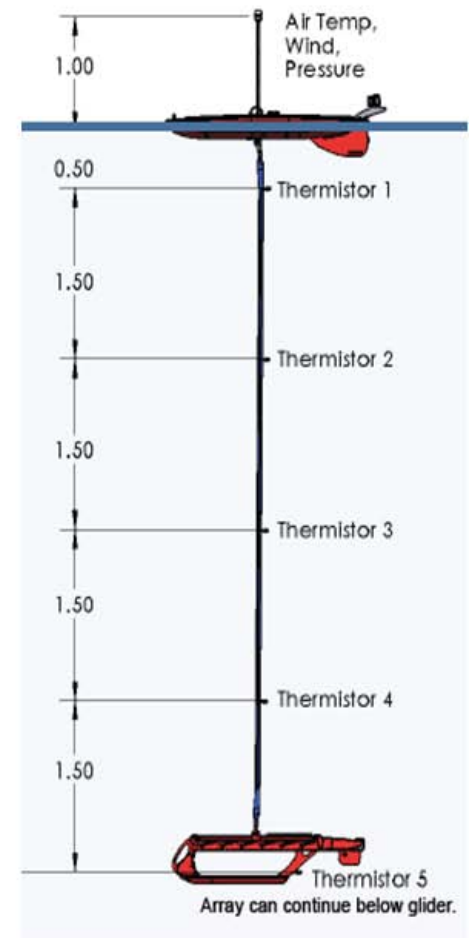
buoys



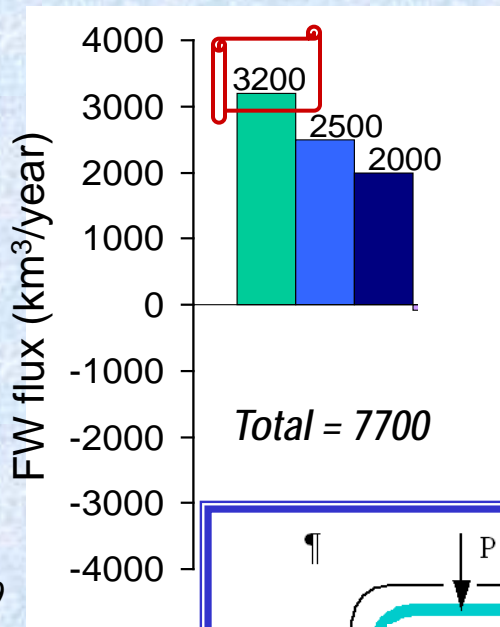
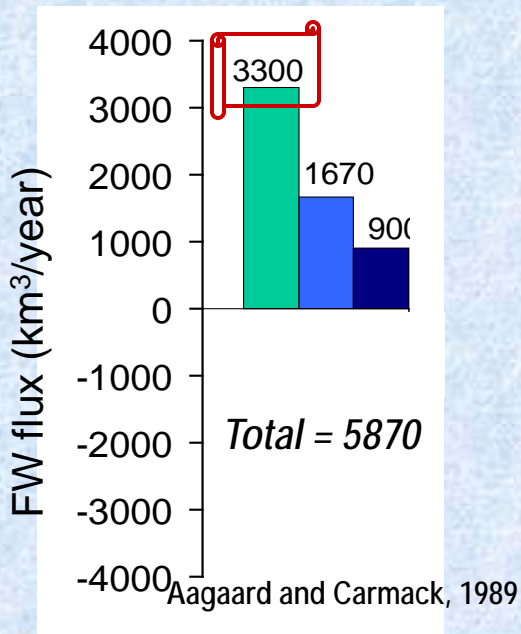
profiling floats



wave glider



Consider all the FW inputs



River Runoff

Bering Strait liquid

P-E

Woodgate et al., 2005

Russian ice camp data

Q: Where's the ice term?

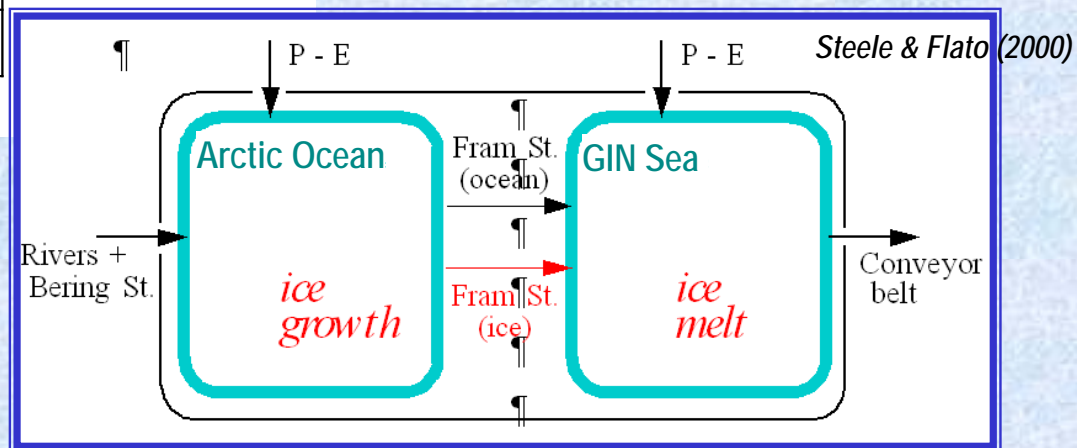
summer is short...

For the Arctic Ocean,

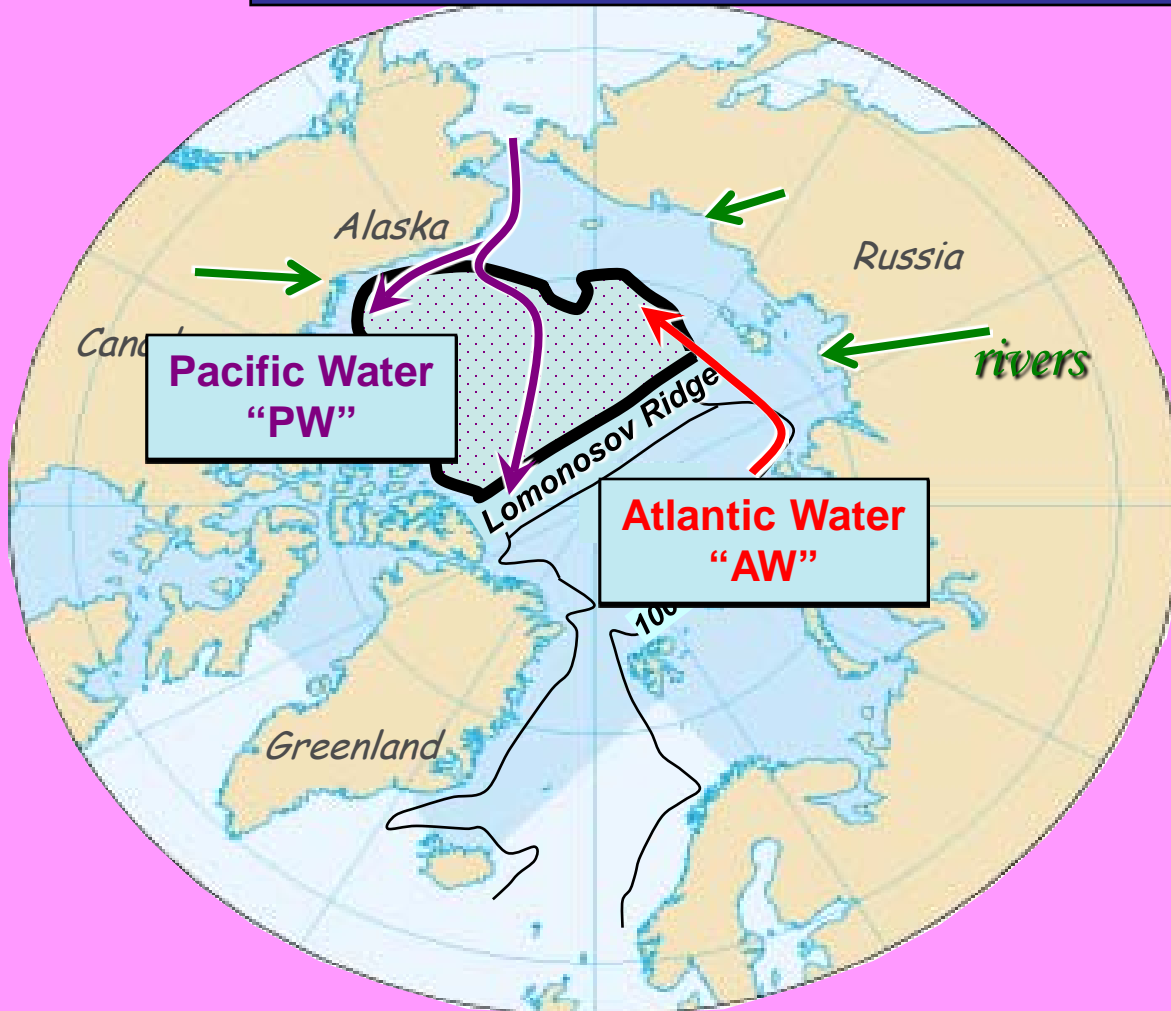
net sea ice GROWTH is a FW SINK $\cong 0.4 \text{ m/yr} \times 7 \text{ million km}^2 = 2800 \text{ km}^3/\text{yr}$



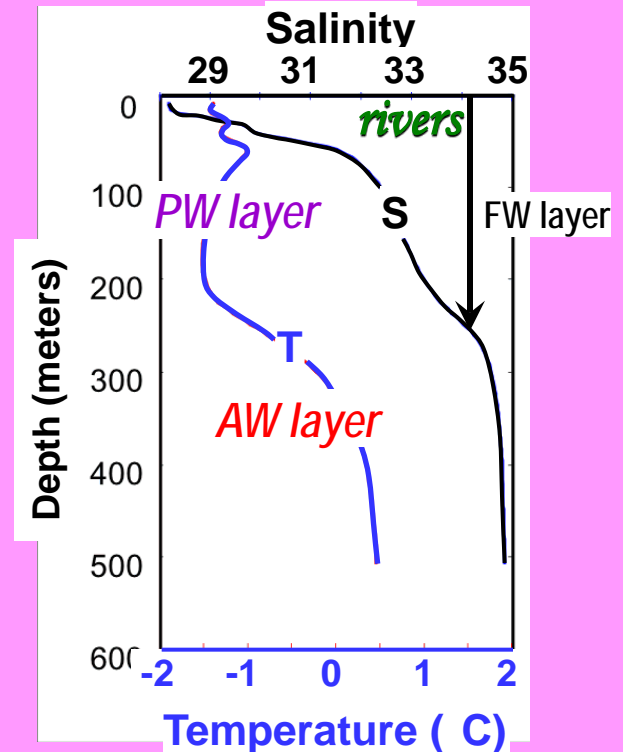
“river runoff → sea ice growth”



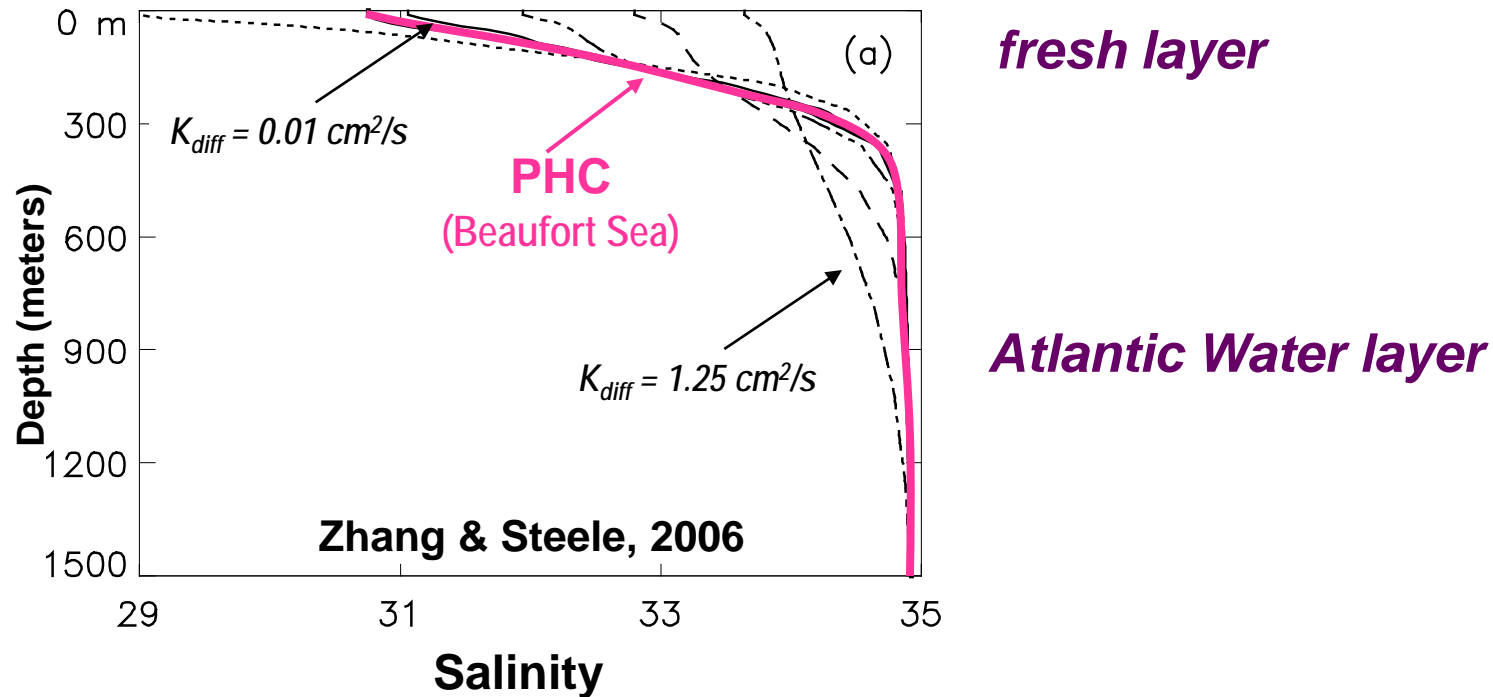
How it stacks up: Ocean water mass structure in the *Beaufort Sea*



Winter profiles



Getting the right stratification: The effect of “background mixing”

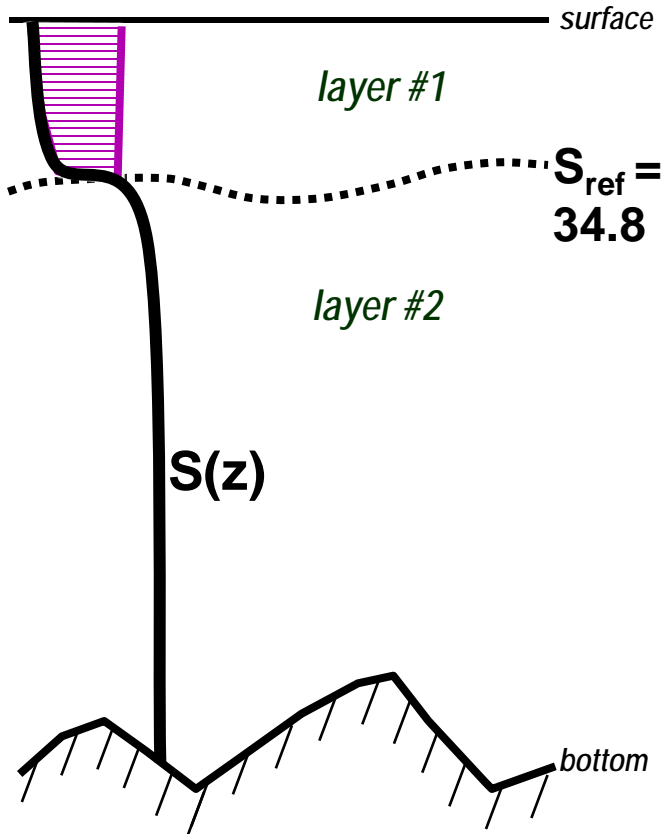


“Typical” $K_{diff} \cong 0.1 \text{ cm}^2/\text{s}$; Need to reduce *mixing* to levels 10 times **less!**

...it's ok: *sea ice suppresses mixing!* (D'Asaro & Morison, 1992)

...for now, anyway: (Rainville & Woodgate, 2009)

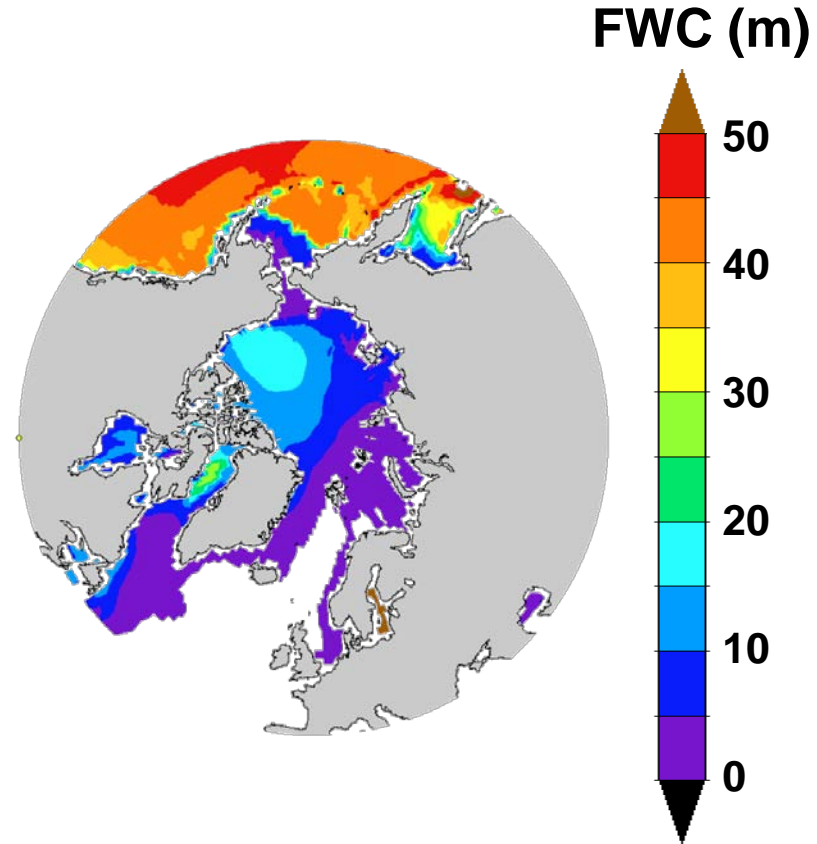
FWC = Vertical integral of FW

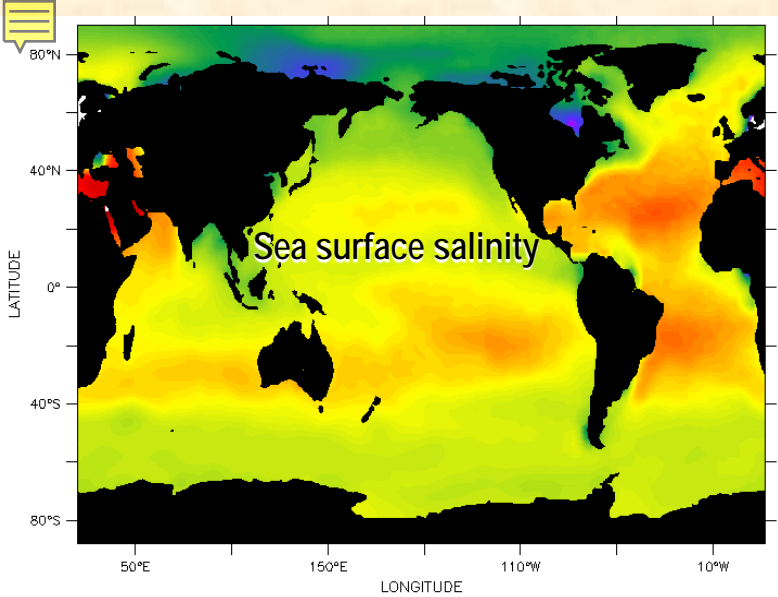


$$FWC \equiv \int_{S=34.8}^0 (S_{ref} - S)/S_{ref} dz$$

Units: meters

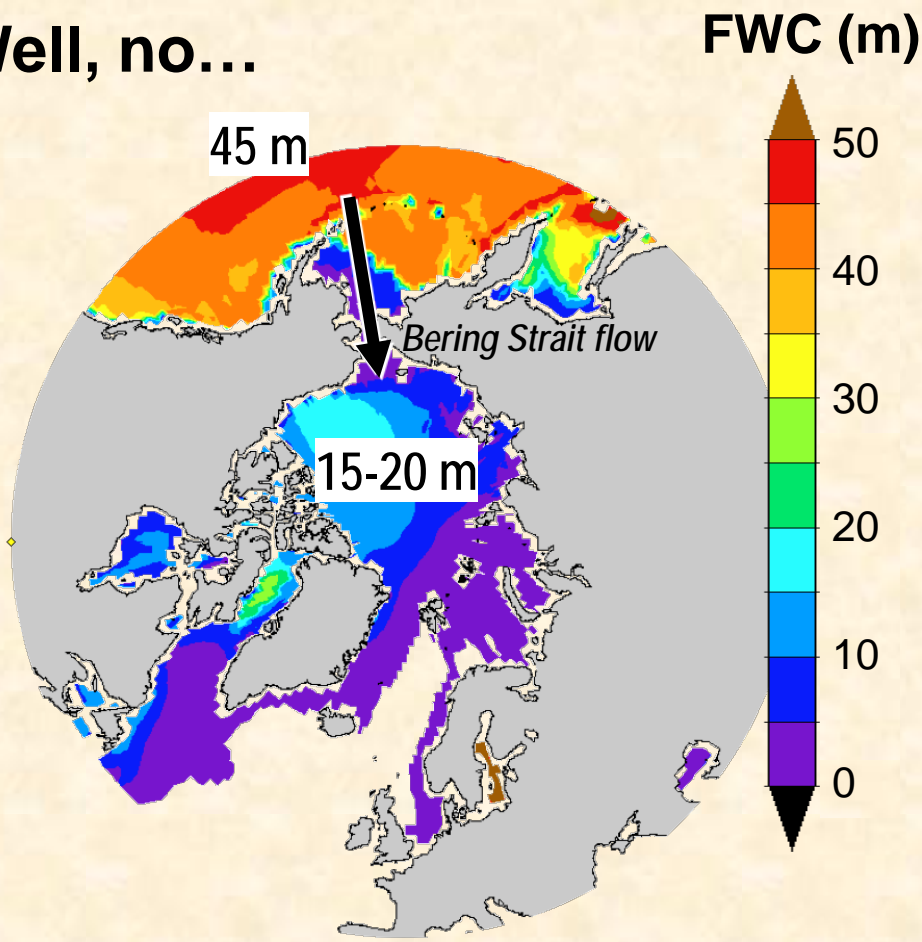
Sea ice melt: ~1 or 2 m/yr





“The Arctic is the freshest ocean!”

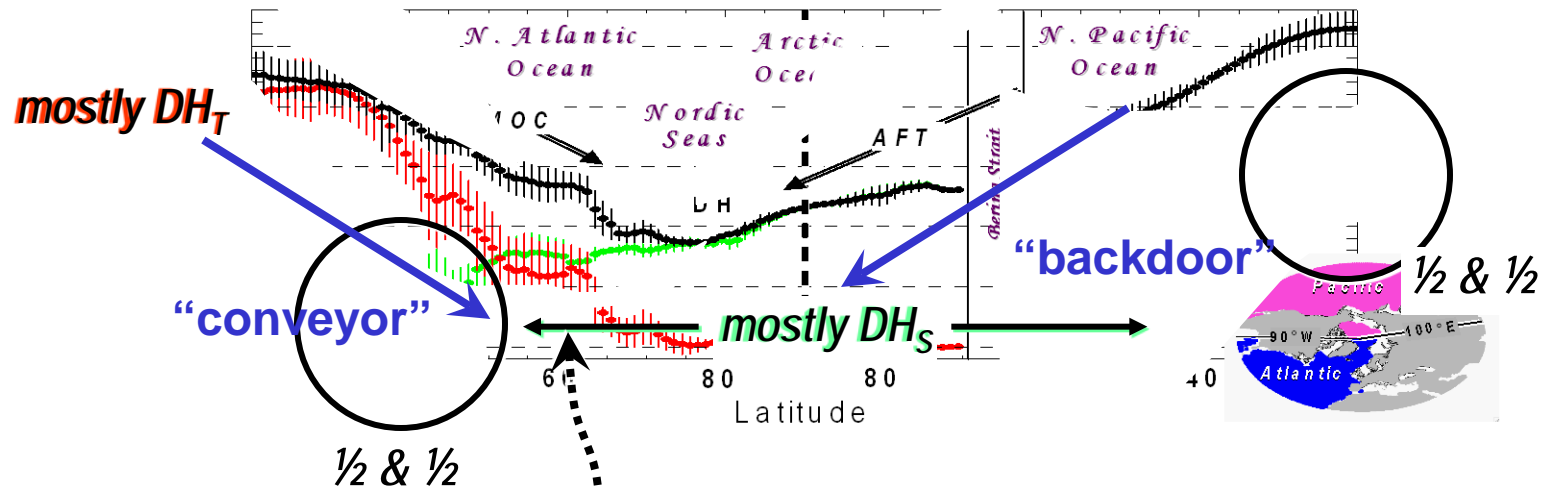
Well, no...



$FWC \cong \text{factor} * \text{steric height}$

DH: The zonal mean view

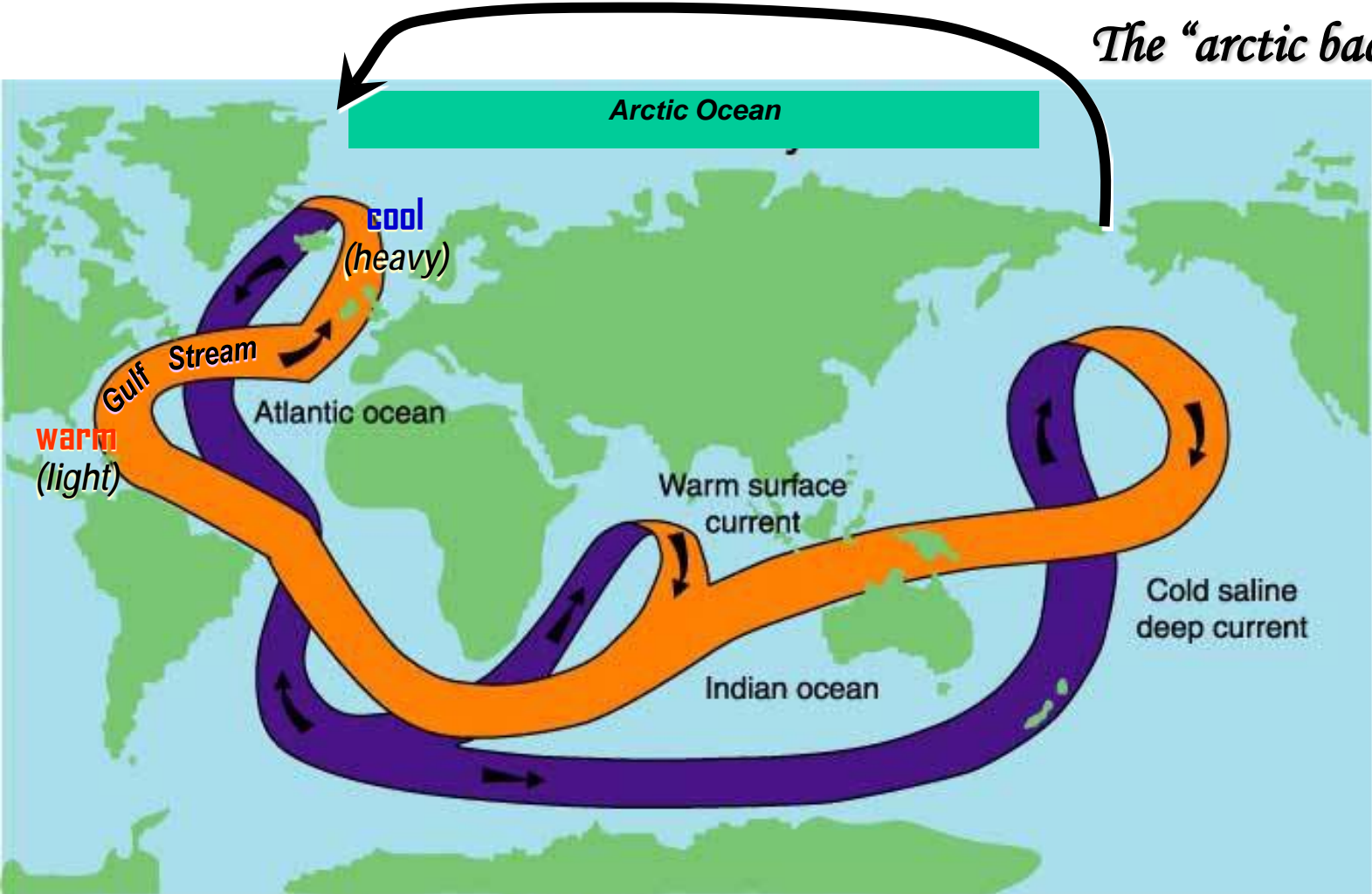
Figure 2



Nordic Seas "potential well"

Ocean circulation pathways

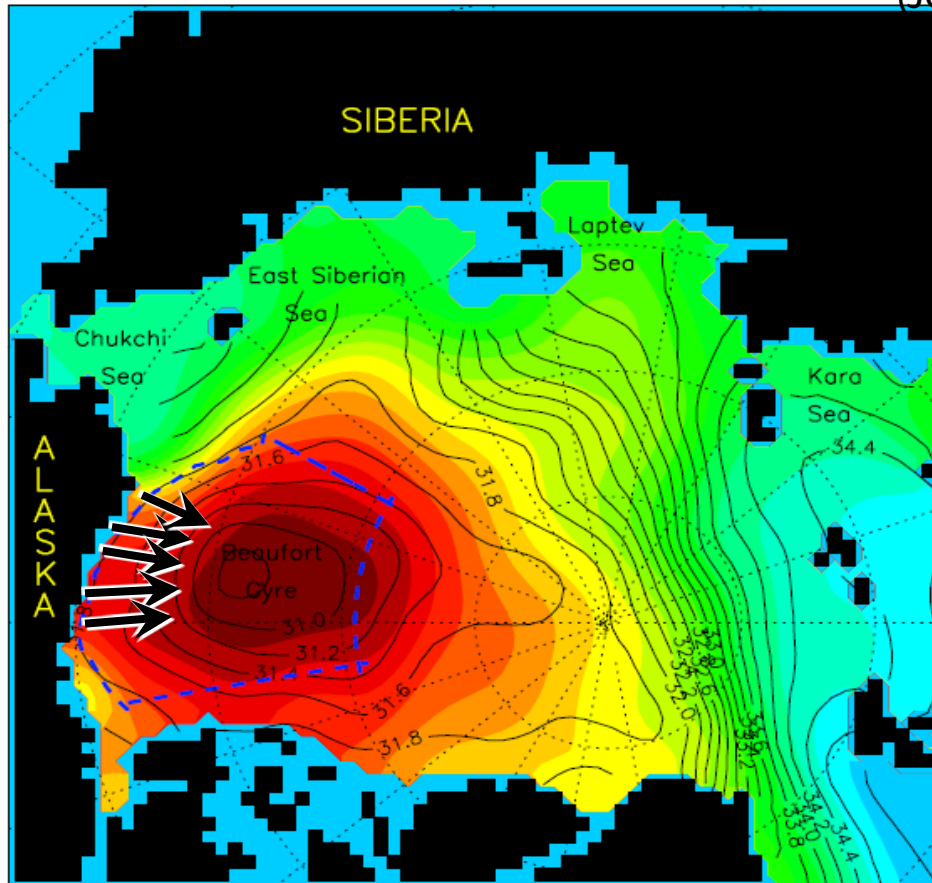
The "arctic backdoor"



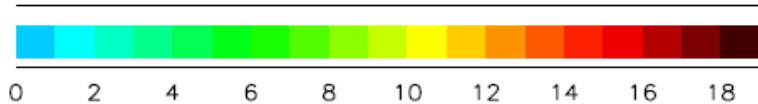
FWC: Seasonal changes

PROSHUTINSKY ET AL.: BEAUFORT GYRE FRESHWATER RESERVOIR

(JGR, 2009)



Freshwater content, meters



There's 2 seasonal maxima:

June/July:

FWC **peak #1** from **ice melt**

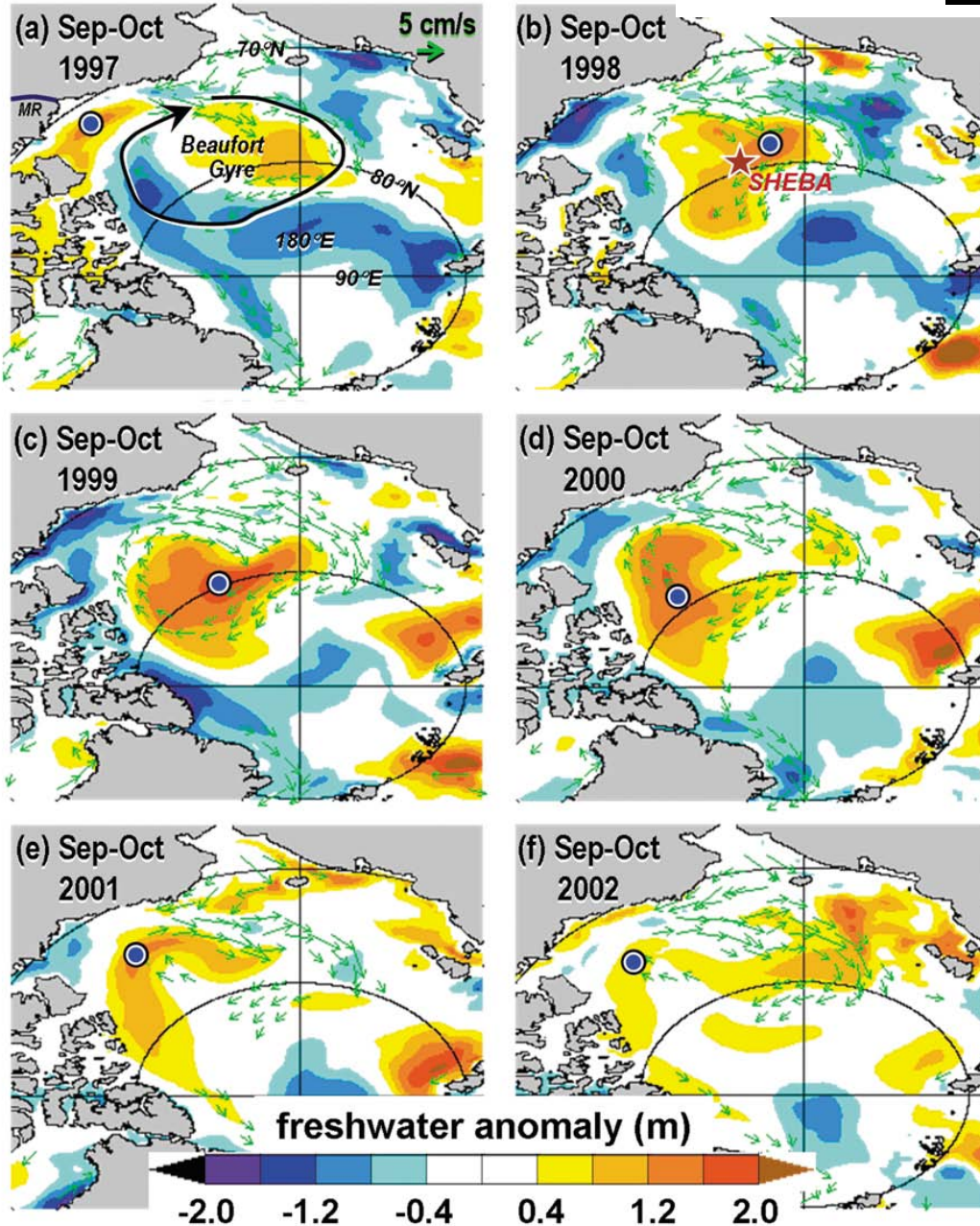
→ *upper layer freshening*

Nov/Dec/Jan:

FWC **peak #2** from **Ekman convergence**

→ *upper layer deepening*

FWC: Interannual variability



The SHEBA FW anomaly: A model study

- Mackenzie River *fed FW into the Beaufort Gyre, as observed*, forced by:
- anomalously *strong fall SE'lies* (i.e., strong BG)
- 2007: another strong summer/fall BG anti-cyclone!

Future of the BG anticyclone?

Longer term pan-Arctic Ocean FWC

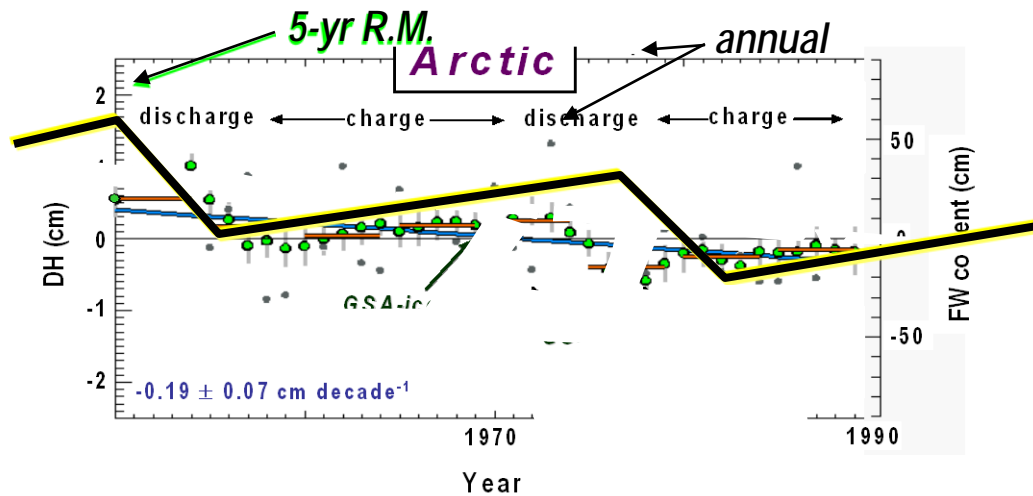
...anomalies

Steele & Ermold, 2007

Figure 4

"fw capacitor"

Proshutinsky et al. (2002)

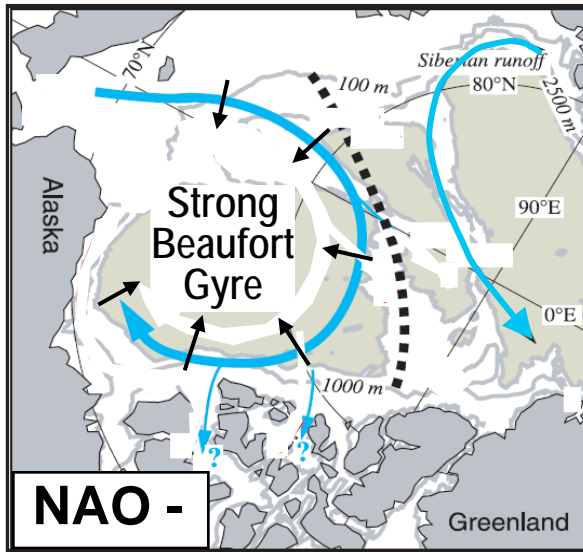


Polyakov et al. (2008):

Downward trend → 1990s

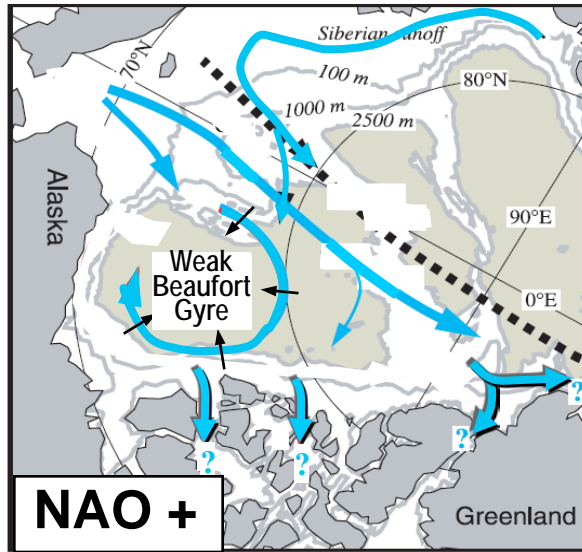
Arctic Ocean getting SALTIER! Why?

Arctic FW & the North Atlantic Oscillation



NAO -

FW storage



NAO +

FW release

Steele et al., 2004

Why did the Arctic Ocean get saltier?

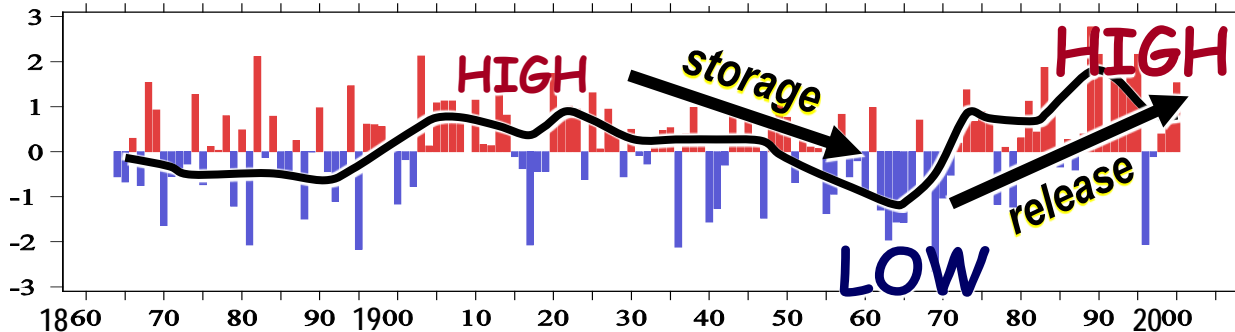
It was the atmospheric forcing!

(i.e., winds → ice, ocean FW release)

(non-anthropogenic?)

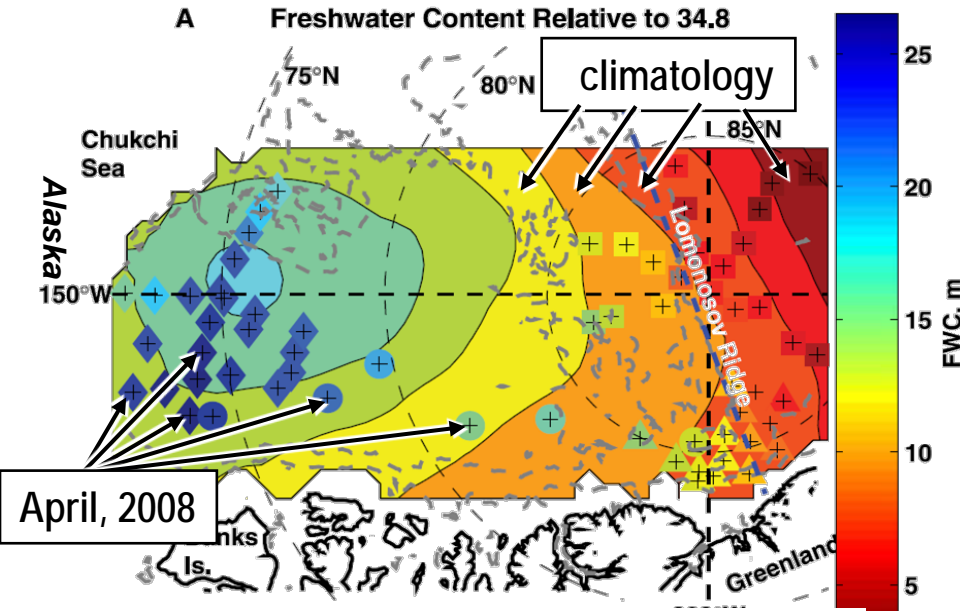
Incr. river discharge was a much smaller effect

NAO index, 1864-2001 (Hurrell)

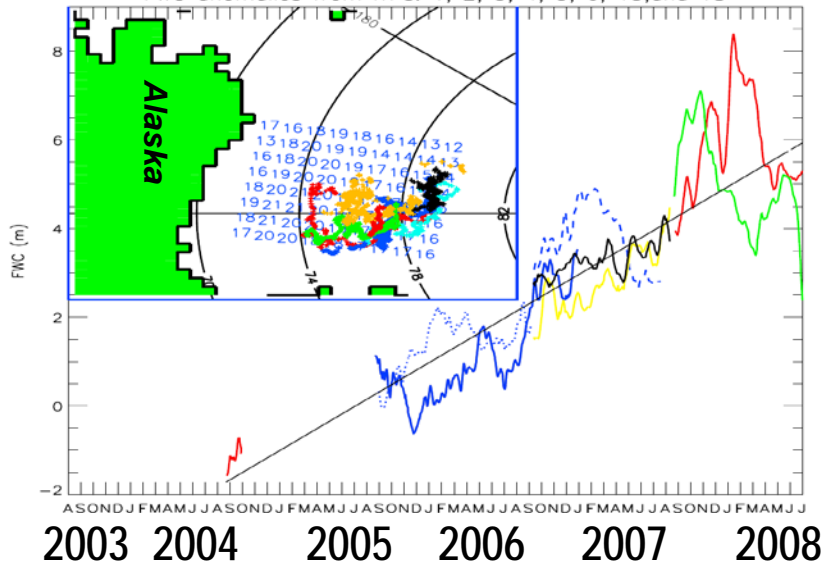


The Beaufort Gyre in the 2000's

McPhee et al. (2009)

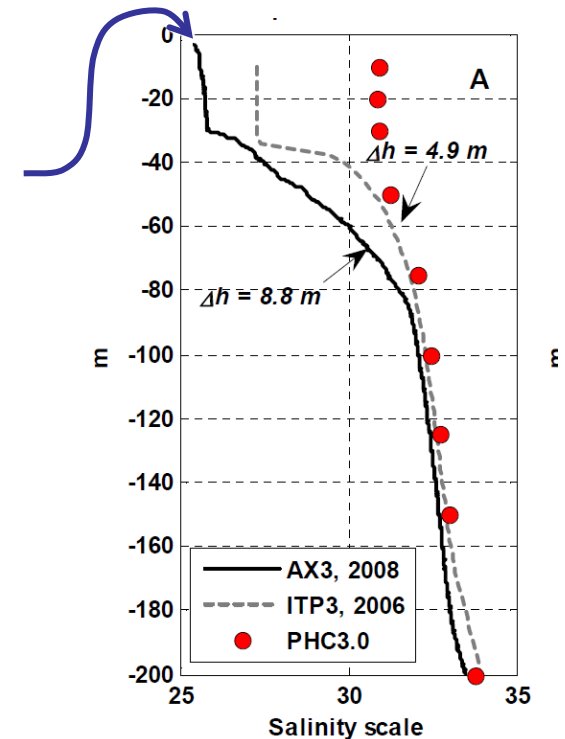


Proshutinsky et al. (2009)



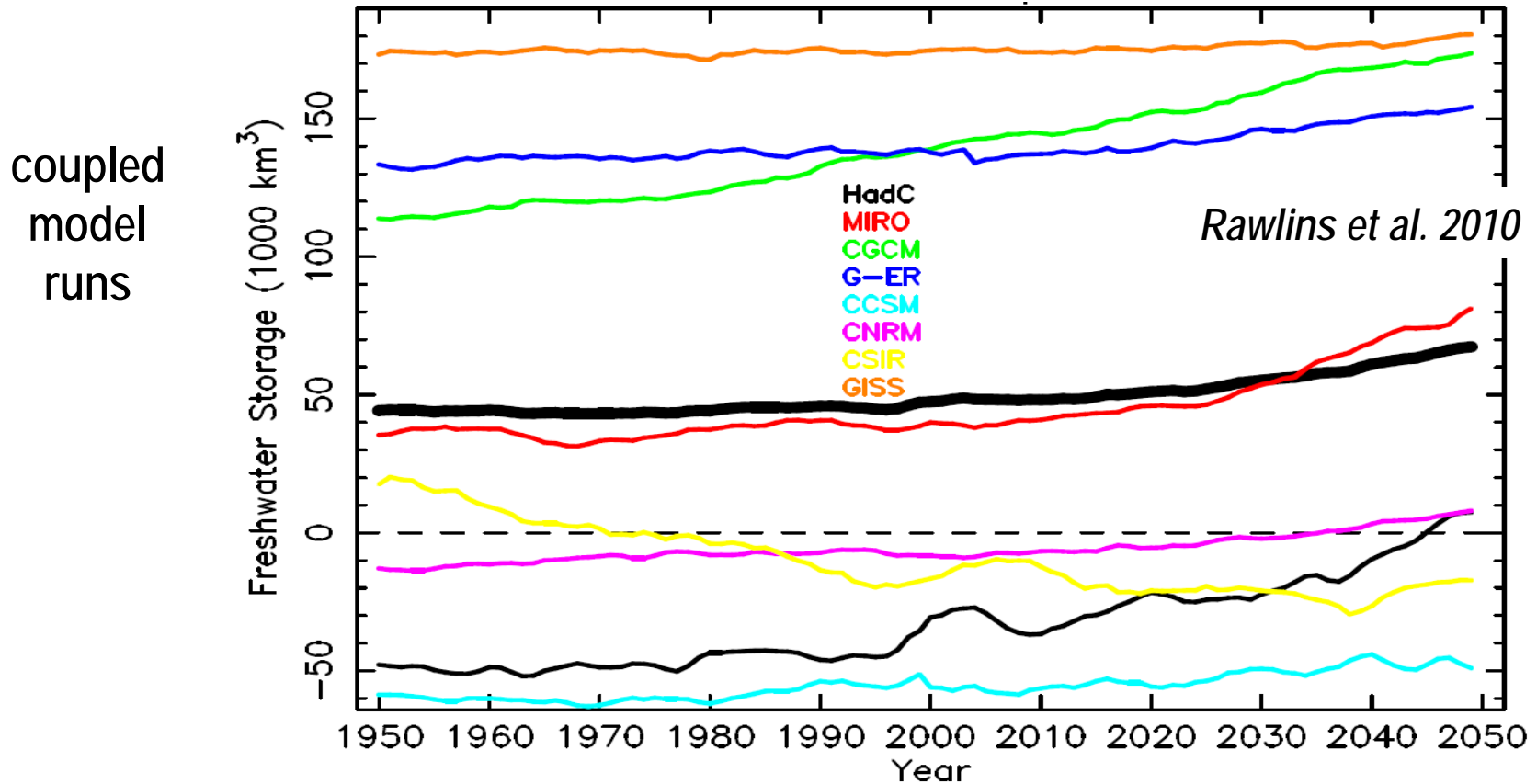
Increasing Beaufort Gyre FWC

winter $S \approx 26!$



The Longer-Term Future?

Arctic Ocean (liquid) FWC



The arctic will freshen ...
... but when?!

Factors influencing Arctic Ocean circulation

- sea ice thinning & retreat
 - vertical **momentum** flux changes (*currents, waves, mixing*)
 - surface **warming** (*ice melt, tracer, density*)
 - vertical **FW** flux changes? (*density*)
- changes in the global hydrologic cycle
 - incr. **FW input** (*density*)
 - altered **Pacific-Atlantic Δ SSH?** (*OMG!*)

Arctic Ocean circulation modeling: What's new?

- numerical improvements (e.g, **resolution**)
- better forcing (e.g., *atmos. reanalyses*)
- new tracers and **diagnostics** (e.g., *biology!*)



Thank You

**The arctic “sea ice refuge”
Lincoln Sea “Switchyard” project, 2009**

The Changing Arctic: Observation and Model Study

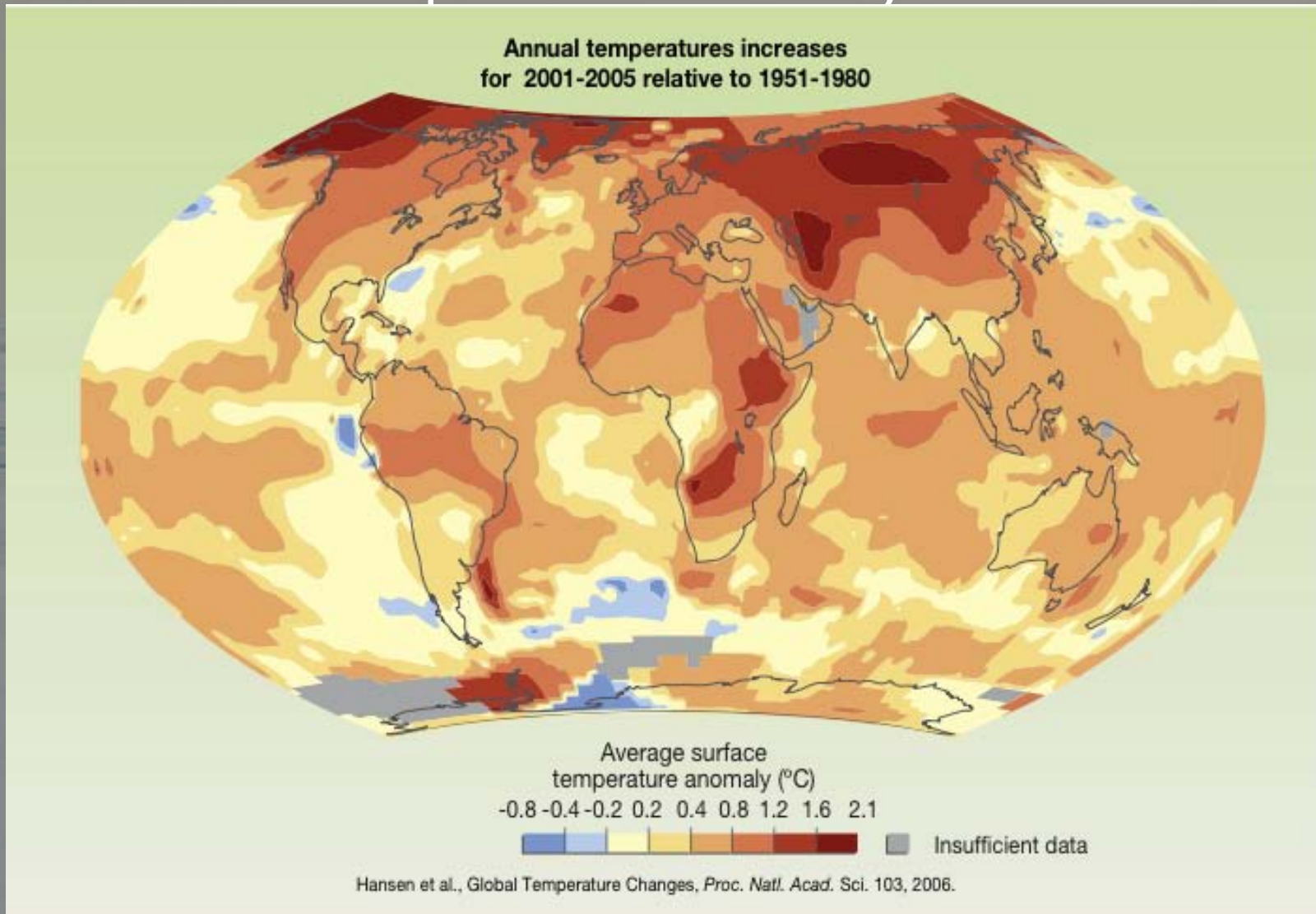
MUYIN WANG¹
& JAMES E. OVERLAND²

¹JISAO/UW, ²PMEL/NOAA

OUTLINE

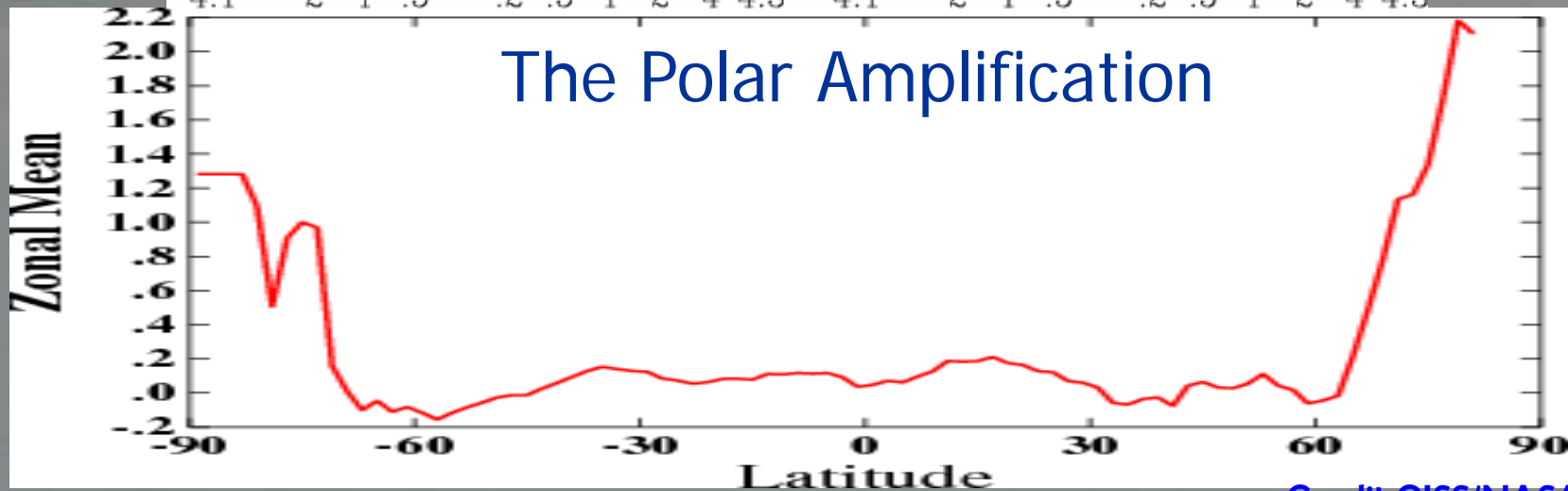
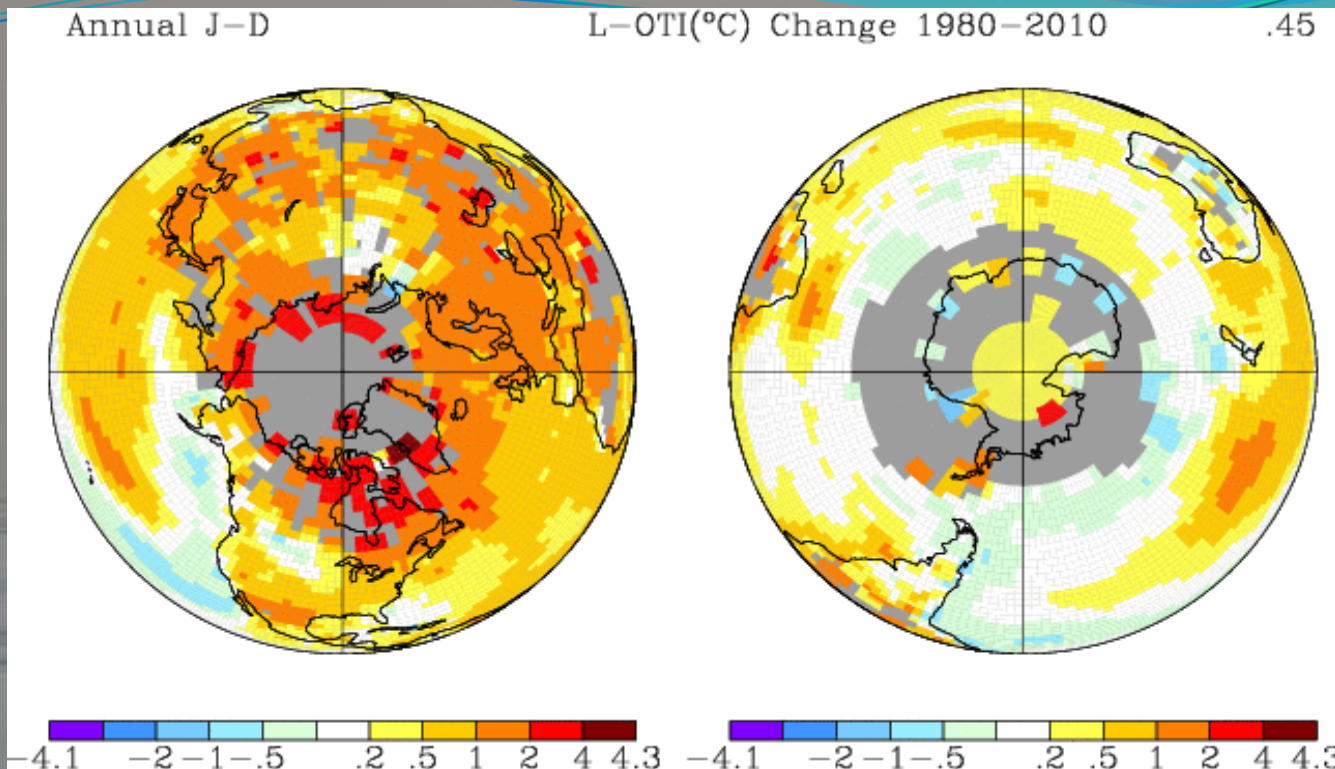
- The observed changes in the Arctic
- Feedback mechanisms
- Model culling and projections of future Arctic Sea ice condition

Surface Air Temperature Anomaly for 2001-2005



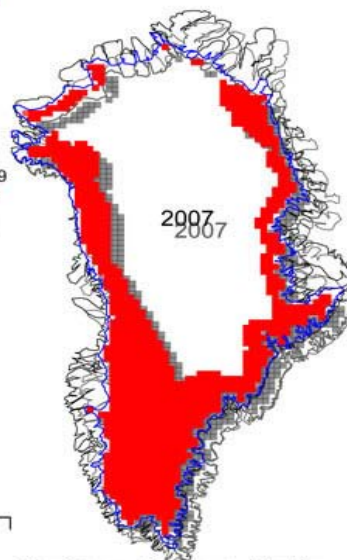
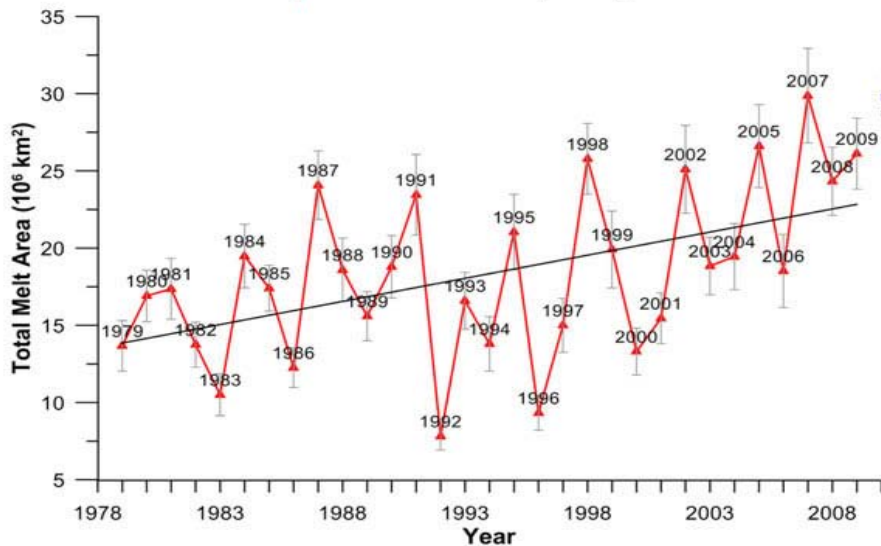
Warming is unequivocal – IPCC AR4

Annual Mean Surface Temperature Linear Trend



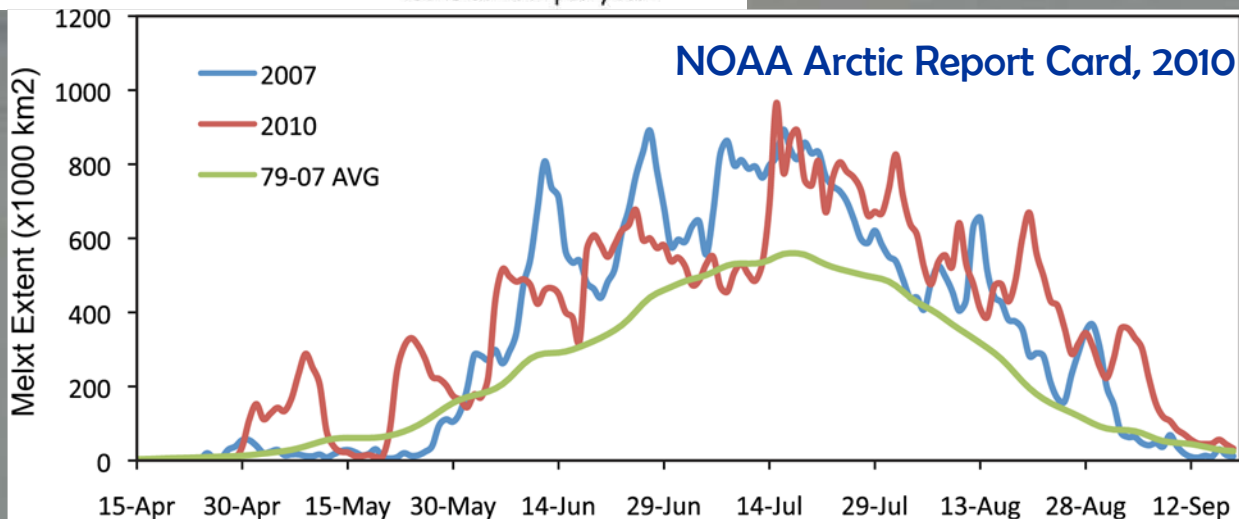
Greenland Total Melt Area: 1979-2009

Total Greenland ice sheet melt area increased 65% since 1979 over the 30 year record; on average 2%/year.

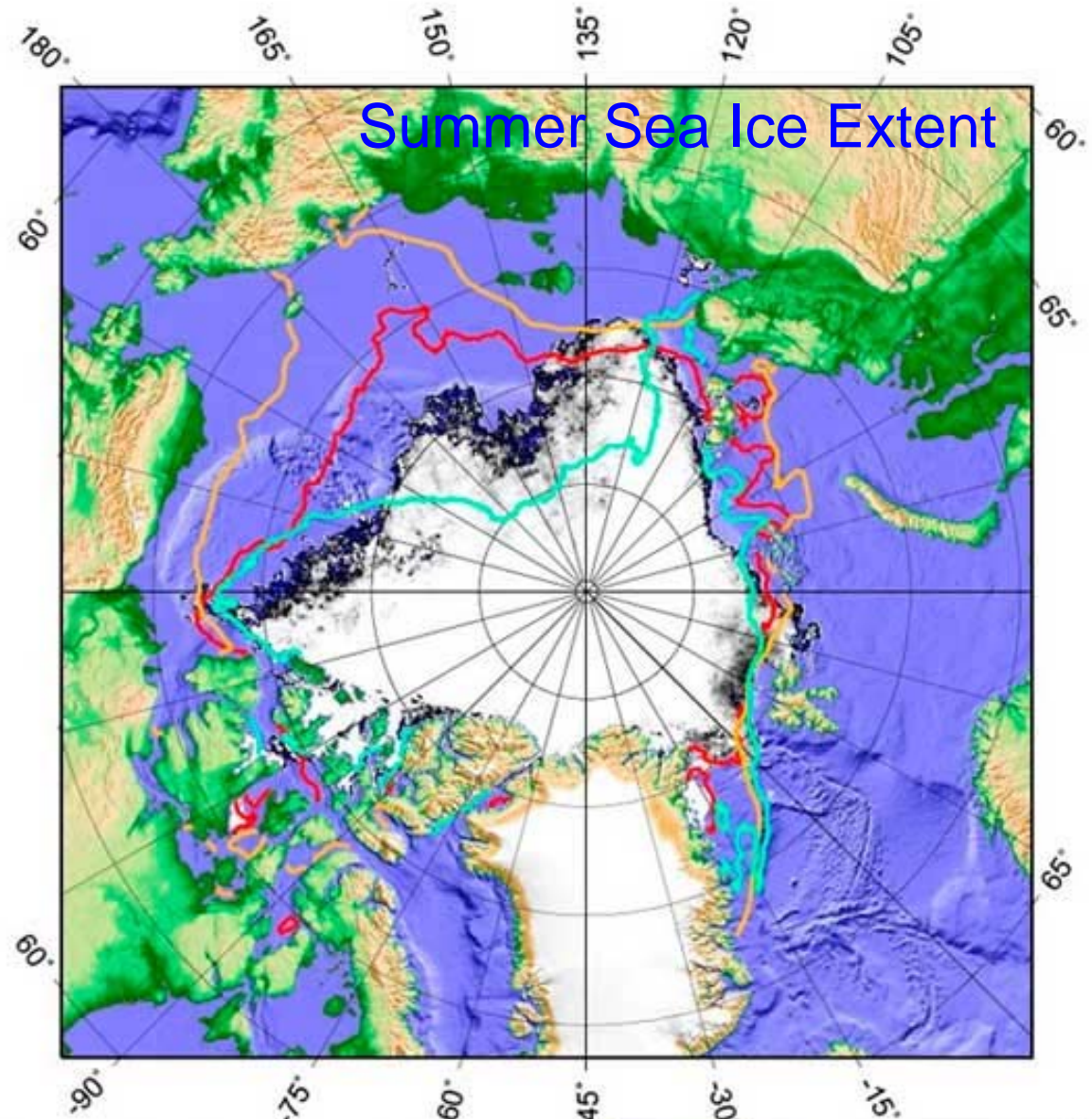


The increasing trend in the total area of melting bare ice is at 13% per year

<http://www.sciencepoles.org>

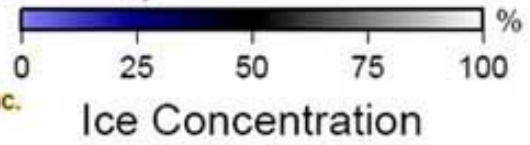


Summer Sea Ice Extent

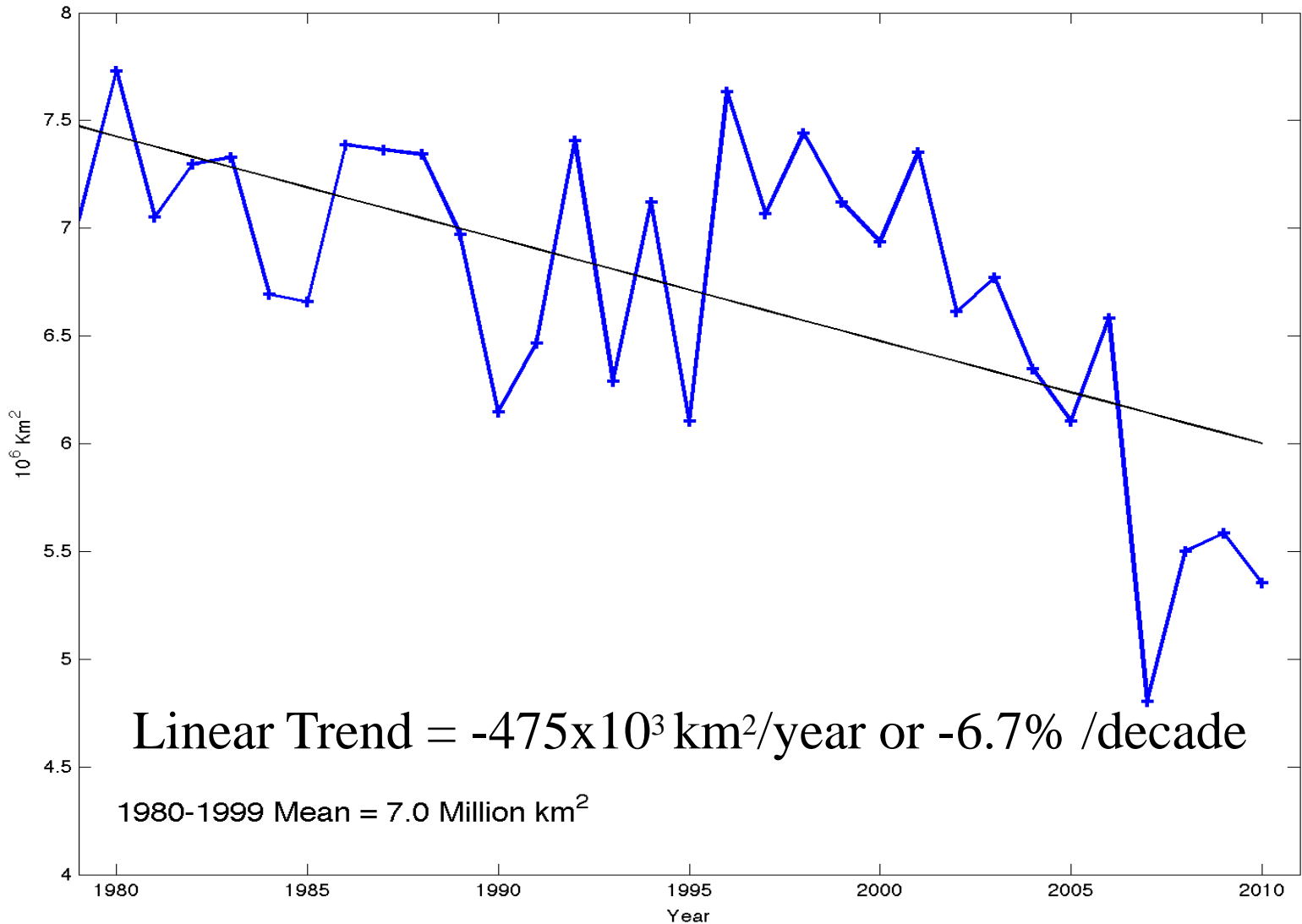


AMSR-E ASI 2008-09-18

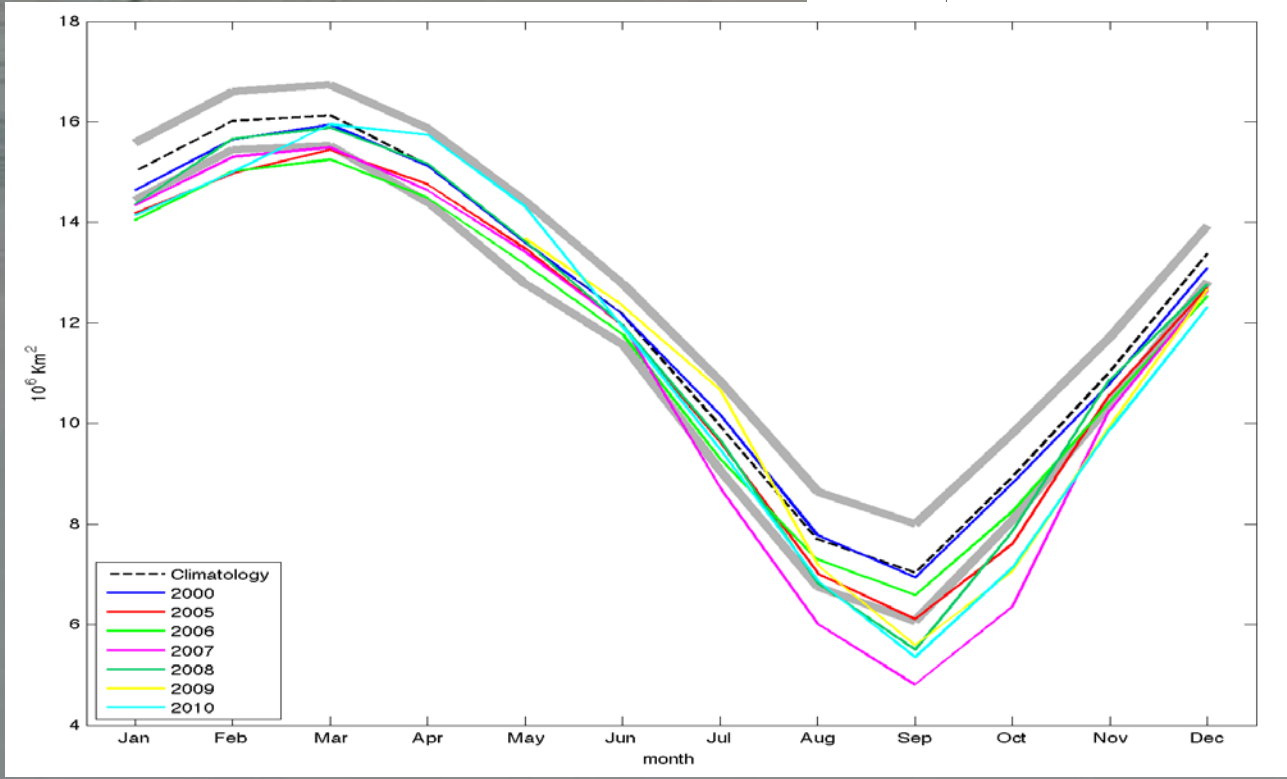
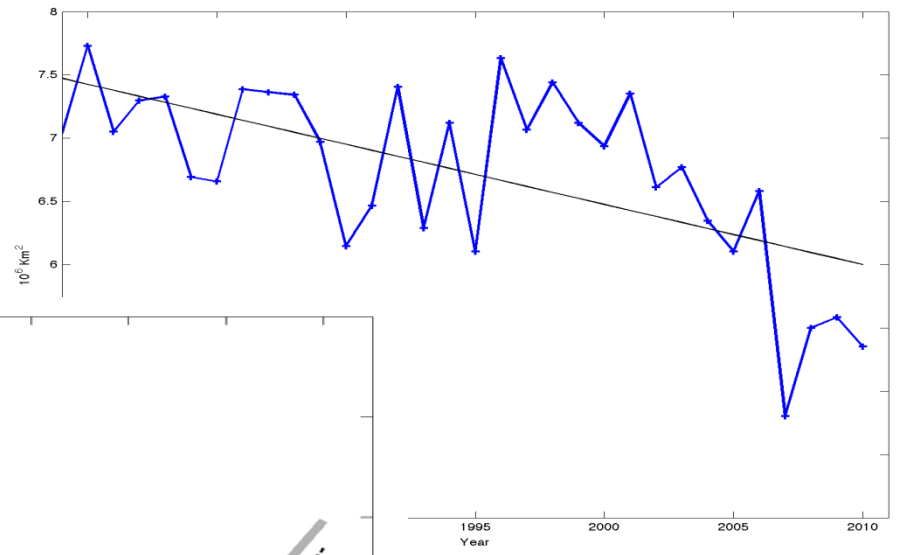
- orange: Sep 1979-1983 SMMR Bootstrap 50% ice conc.
- red: Sep 2002-2006 AMSR-E ASI 50% ice conc.
- green: Sep 2007 AMSR-E ASI 50% ice conc.



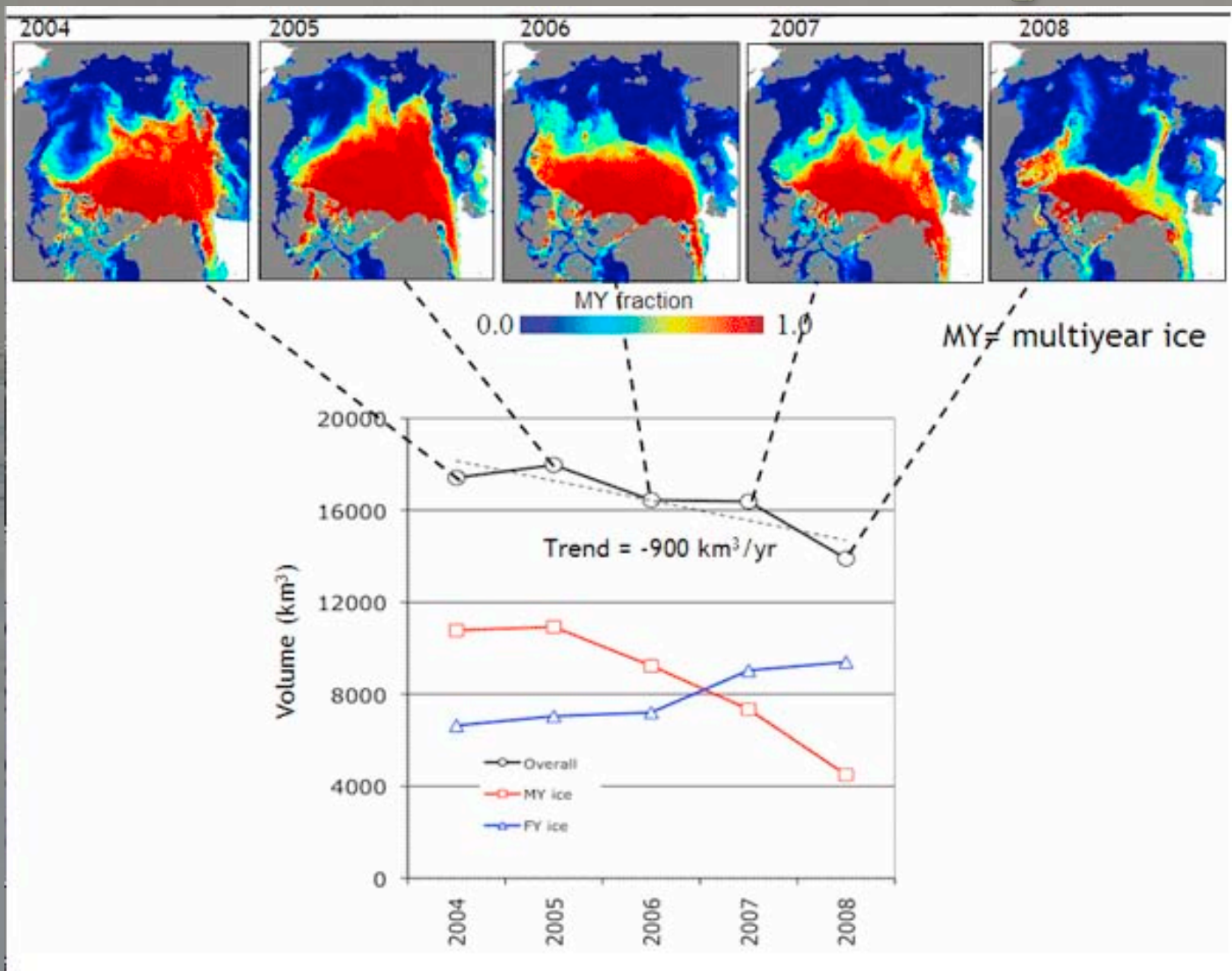
Arctic September Sea Ice Extent

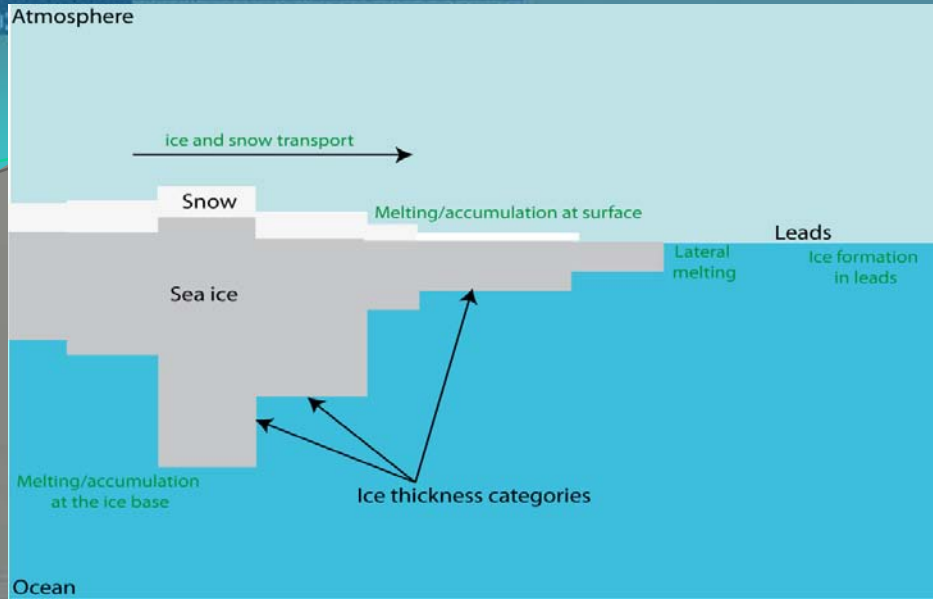


NH Sea Ice Extent



The Evolution of Multiyear Ice





Evolution of Sea Ice Thickness

http://stratus.astr.ucl.ac.be/textbook/chapter3_node12.xml

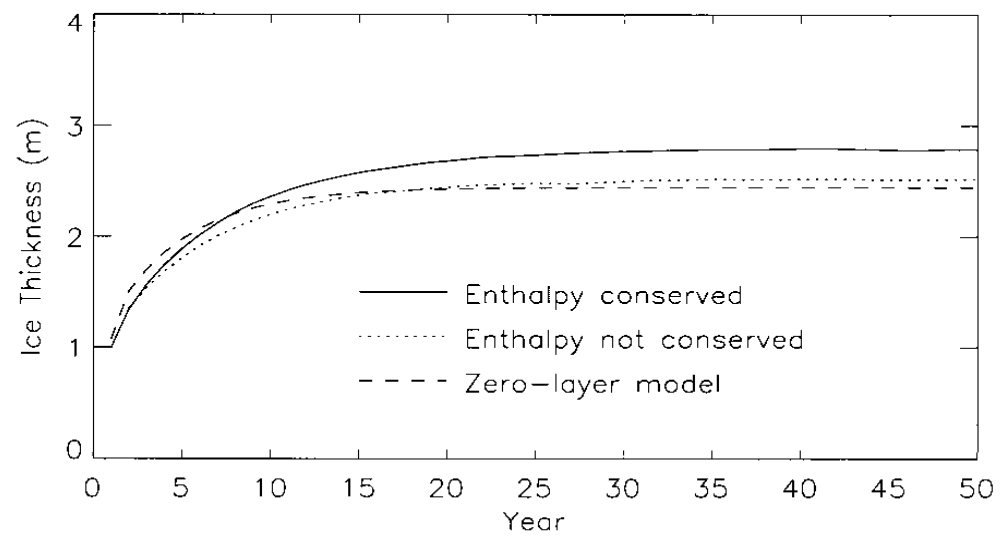


FIG. 5. Evolution of the simulated annual mean ice thickness.

Zhang and Rothrock, 2001

OUTLINE

- The observed changes in the Arctic
- Feedback mechanisms
- Model culling and projections of future Arctic Sea ice condition

THE FEEDBACKS

2 *a* : the partial reversion of the effects of a process to its source or to a preceding stage

b : the transmission of evaluative or corrective information about an action, event, or process to the original or controlling source; *also* : the information so transmitted

Merriam-Webster Dictionary

A system exhibiting **positive feedback**, in response to perturbation, acts to increase the magnitude of the perturbation.

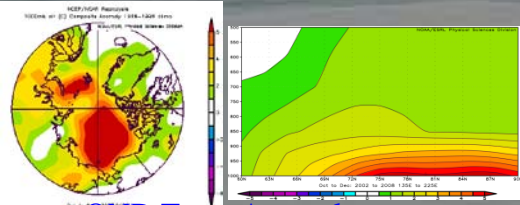
http://en.wikipedia.org/wiki/Positive_feedback

Arctic Climate System Feedbacks

Teleconnection and circulation pattern change

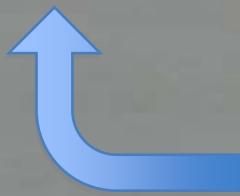


Arctic Atmosphere Warming

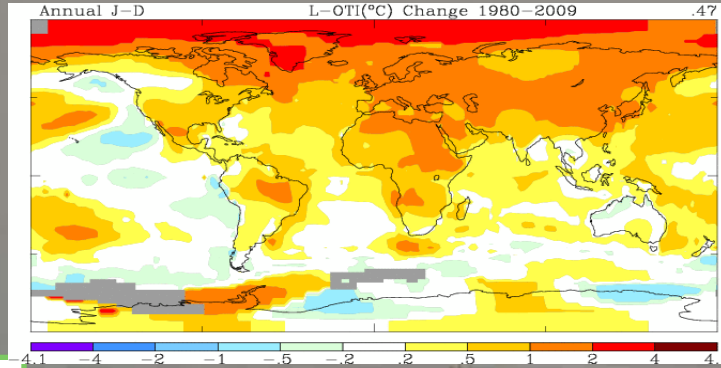


OND Temp Anomaly

Heat releases to atmosphere in the fall.



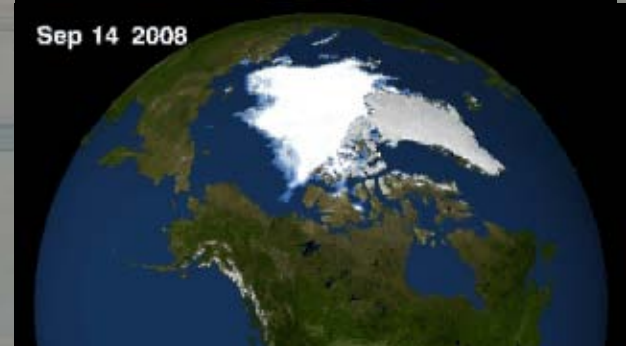
Global Warming



Arctic amplification



Reduction of Arctic Sea Ice

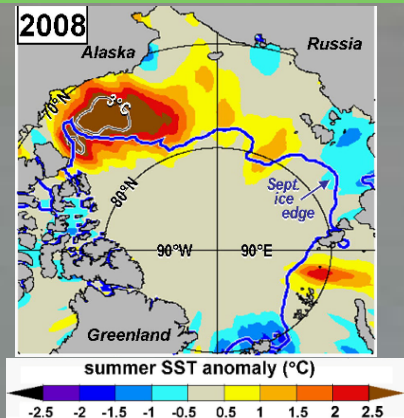


Sept Sea Ice Extent

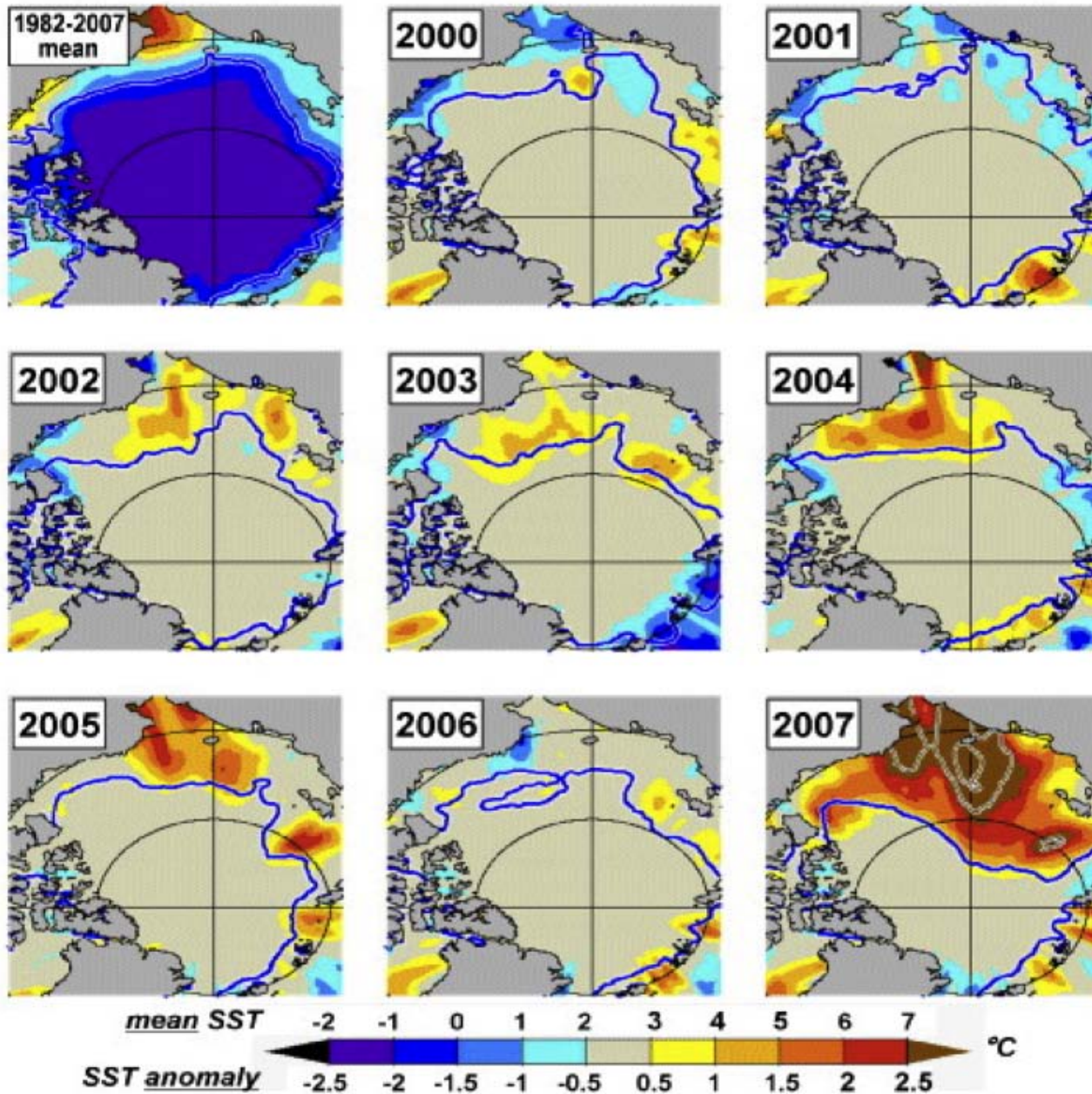
Surface albedo decrease => less sunlight being reflected from surface



Ocean Absorbs More Heat

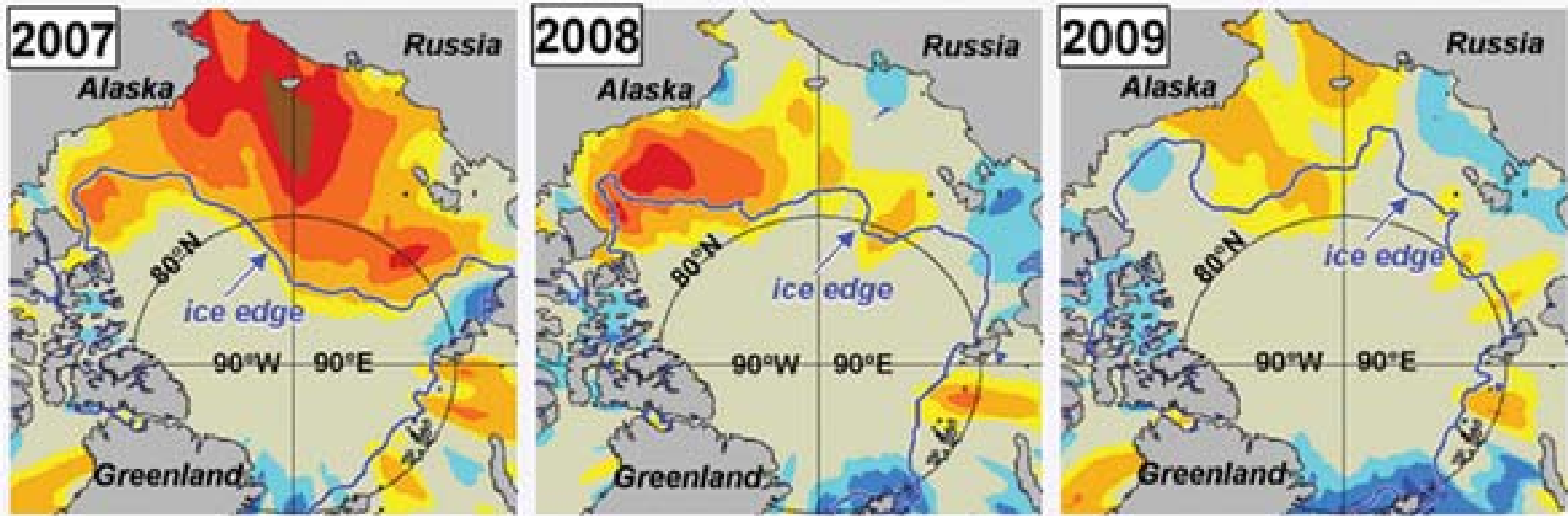


JAS SST Anomaly



JAS Sea Surface Temp Anomaly

Relative to 1982-2006 mean



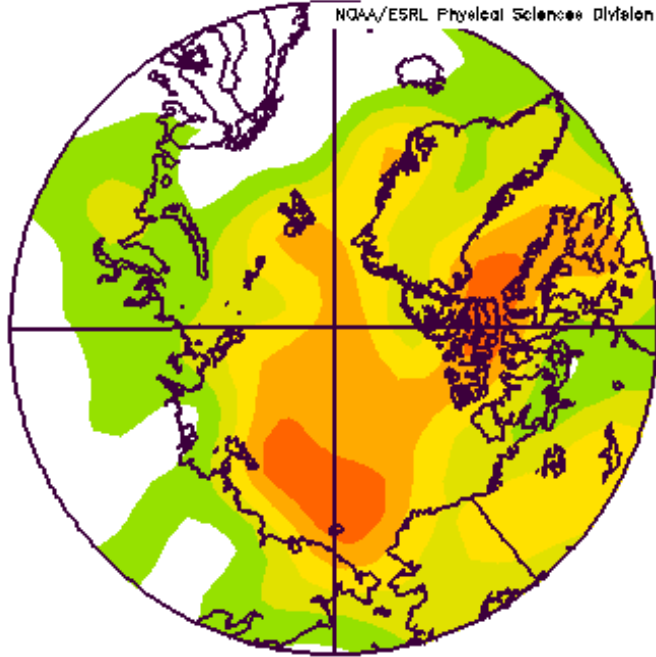
summer SST anomaly (°C)



Recent (2002-2008) Central Arctic Fall Temperature Anomalies Greater Than +5°C

NCEP/NCAR Reanalysis
1000mb air (C) Composite Anomaly 1968-1998 climo

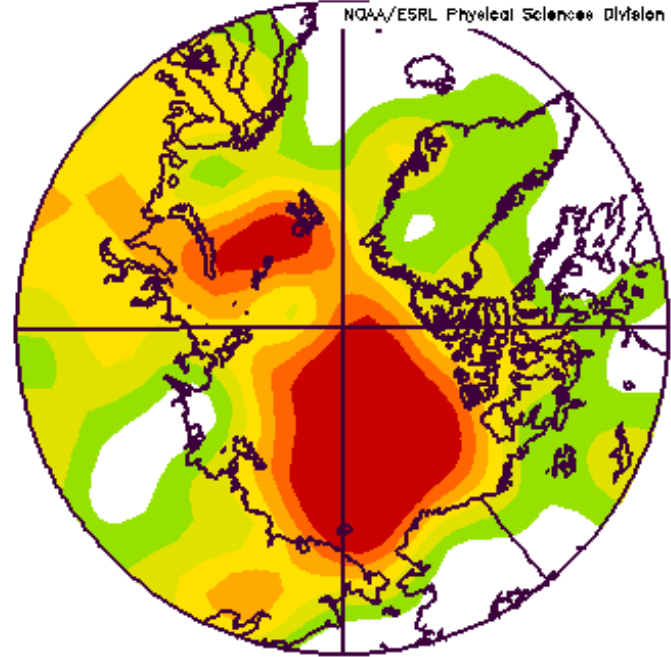
NOAA/ESRL Physical Sciences Division



Oct to Dec: 2002 to 2005

NCEP/NCAR Reanalysis
1000mb air (C) Composite Anomaly 1968-1998 climo

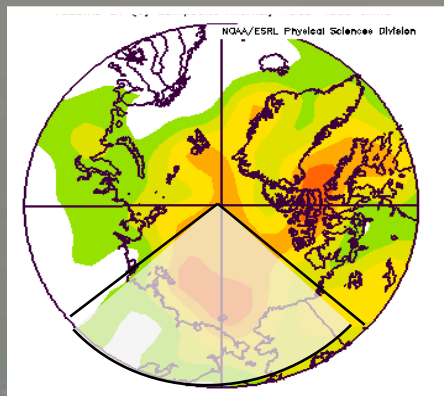
NOAA/ESRL Physical Sciences Division



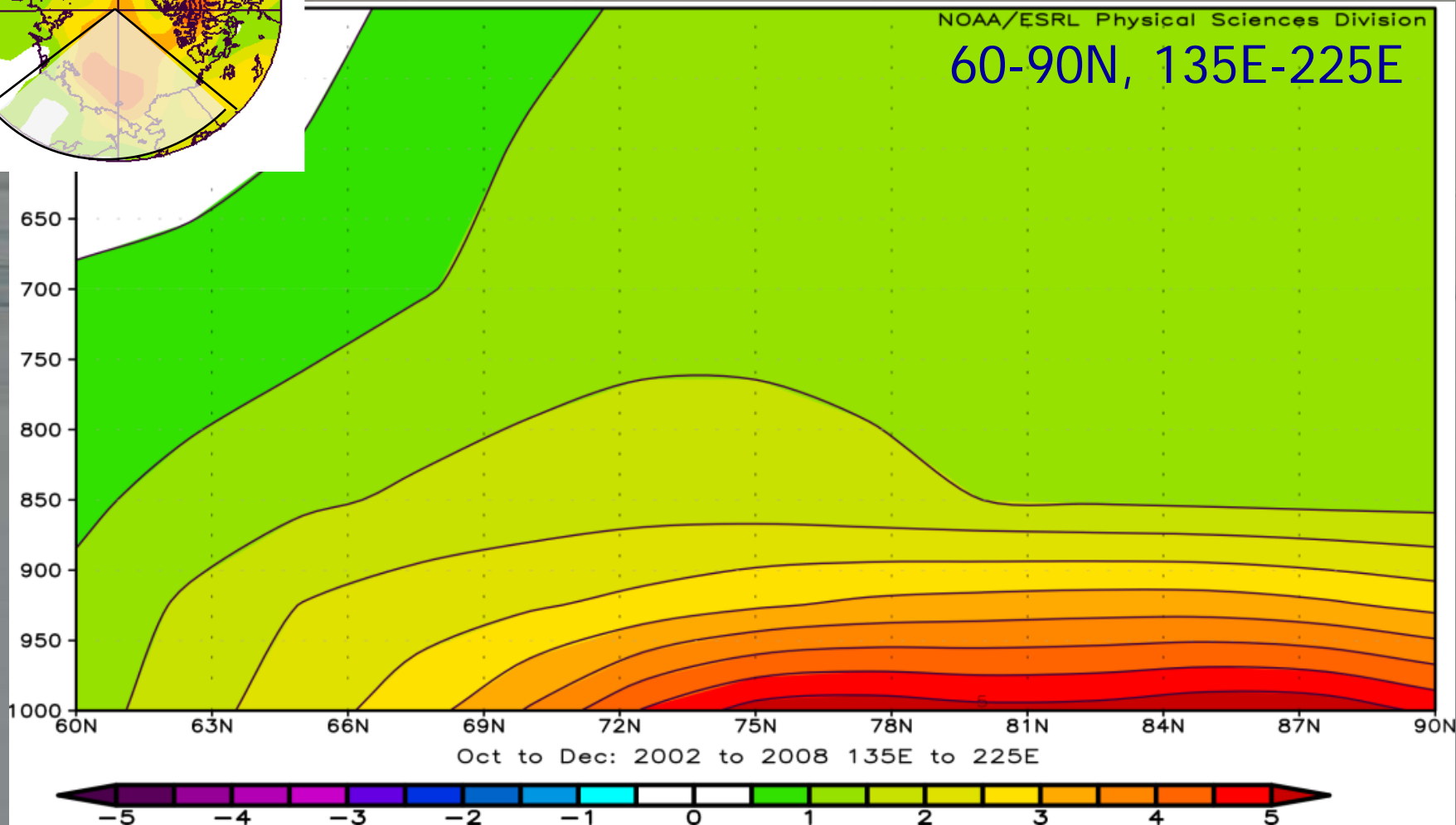
Oct to Dec: 2007,2008

Vertical Cross Section of Temp Anomaly

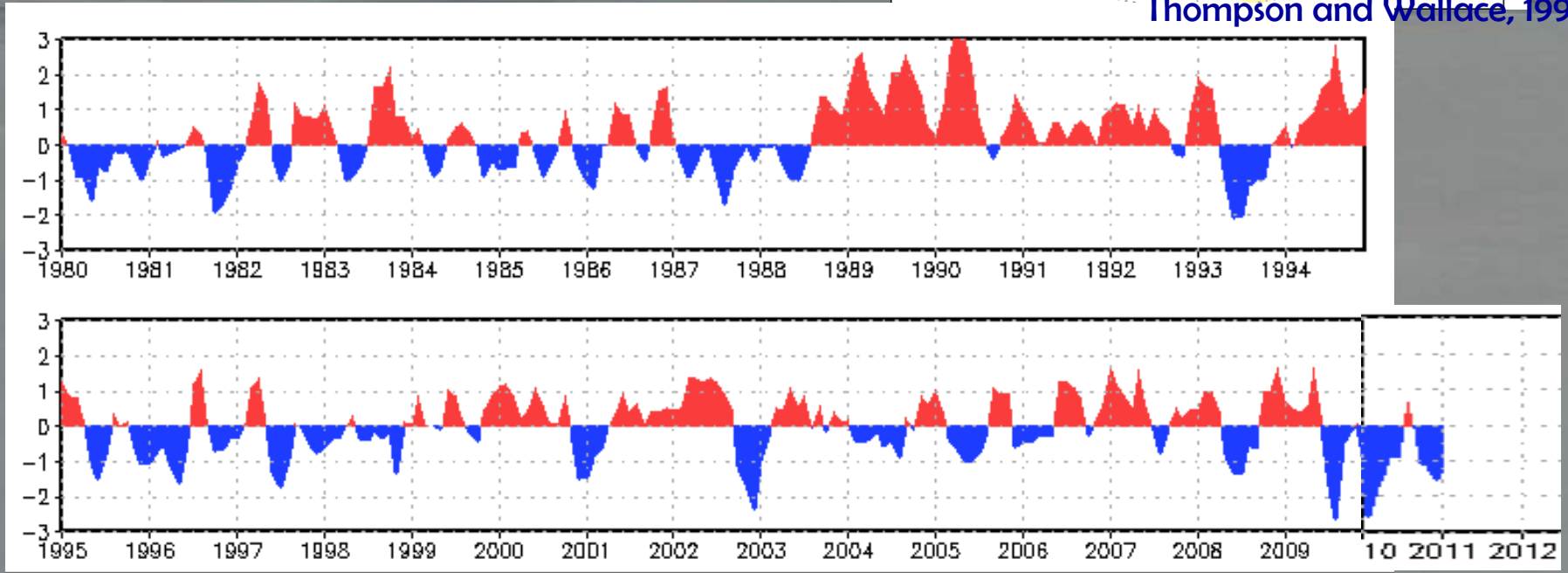
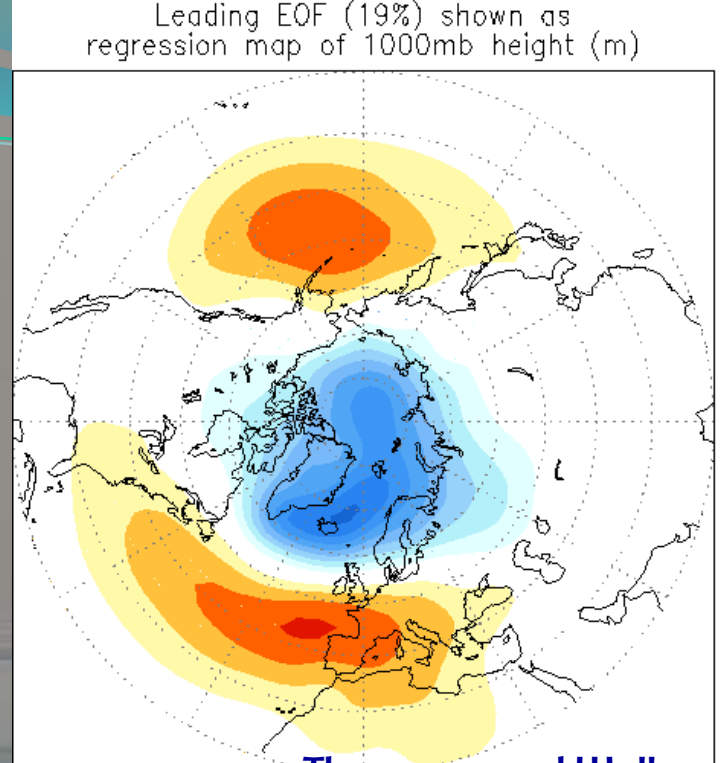
Oct. – Dec. 2002-2008



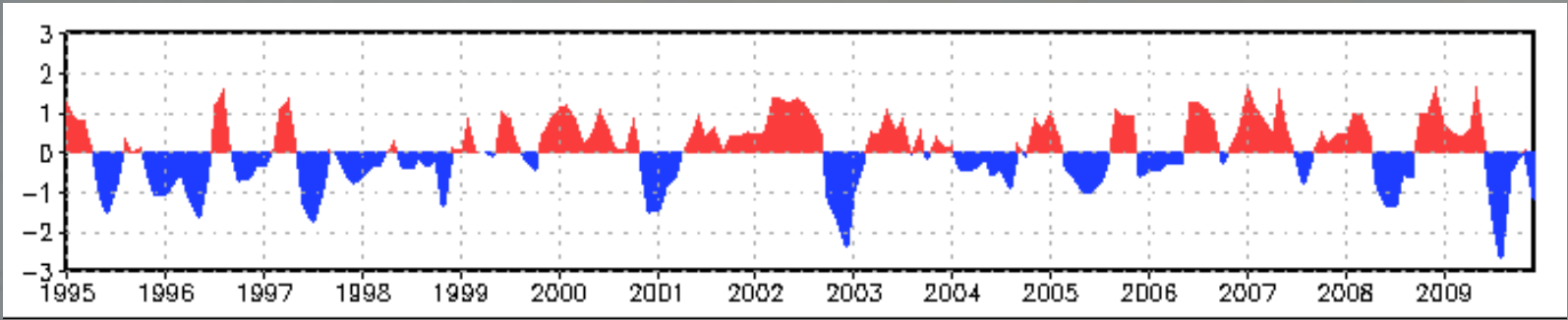
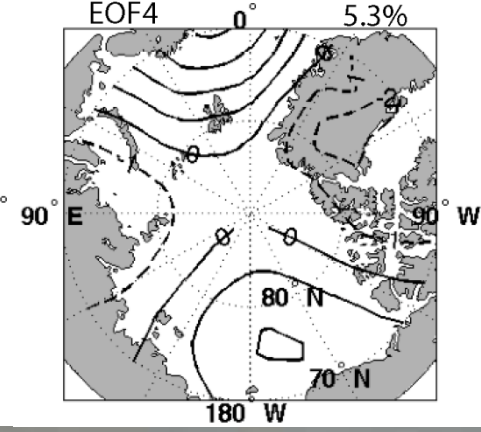
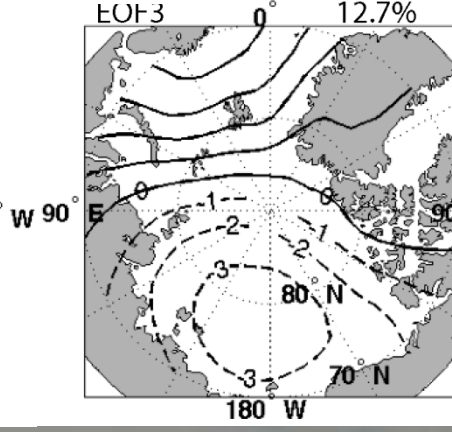
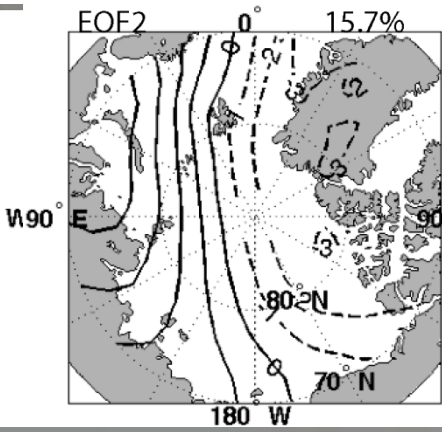
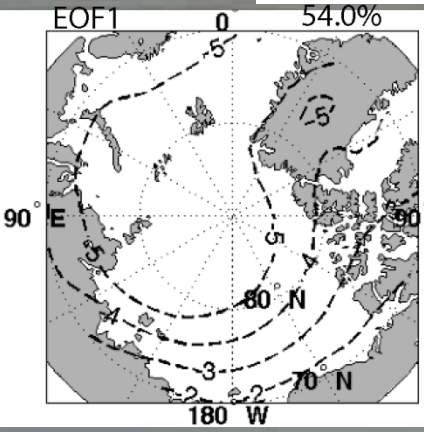
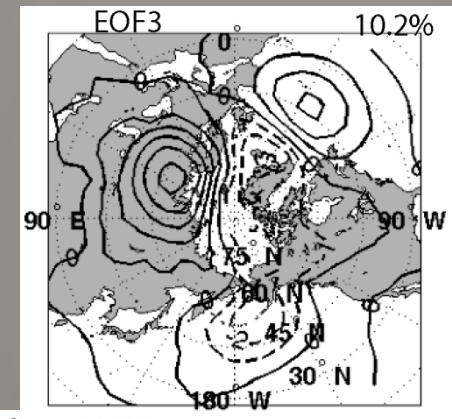
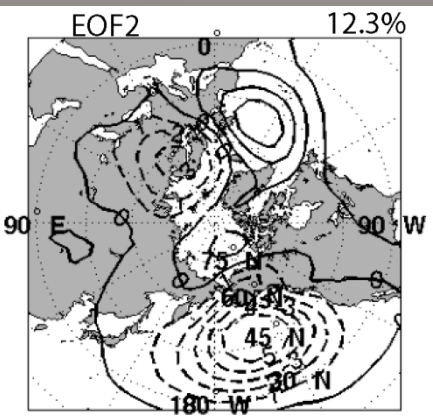
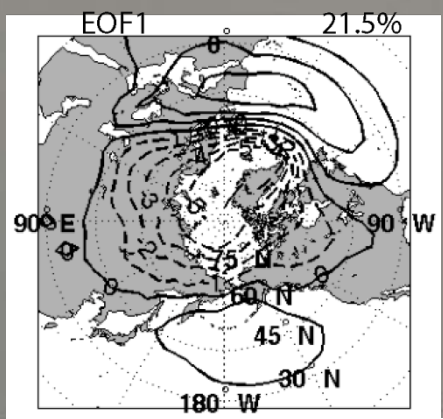
NOAA/ESRL Physical Sciences Division
60-90N, 135E-225E



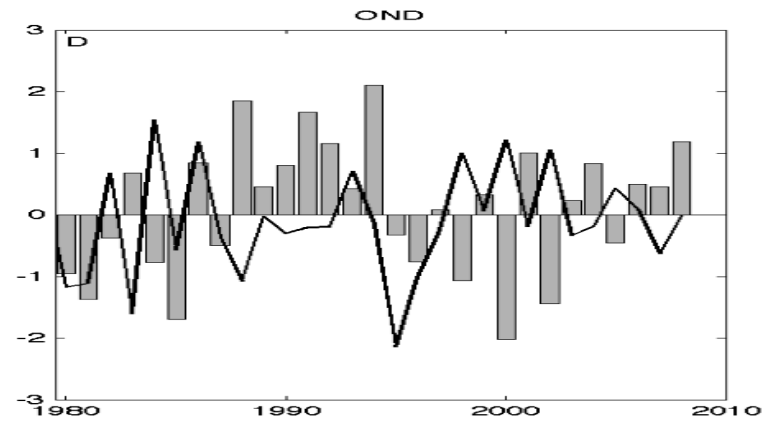
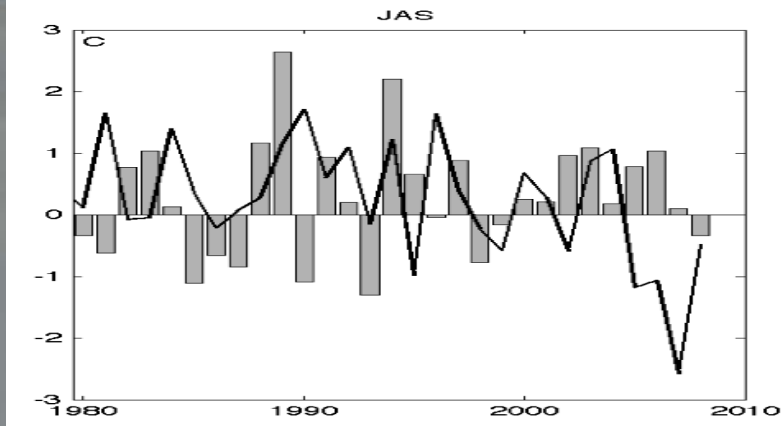
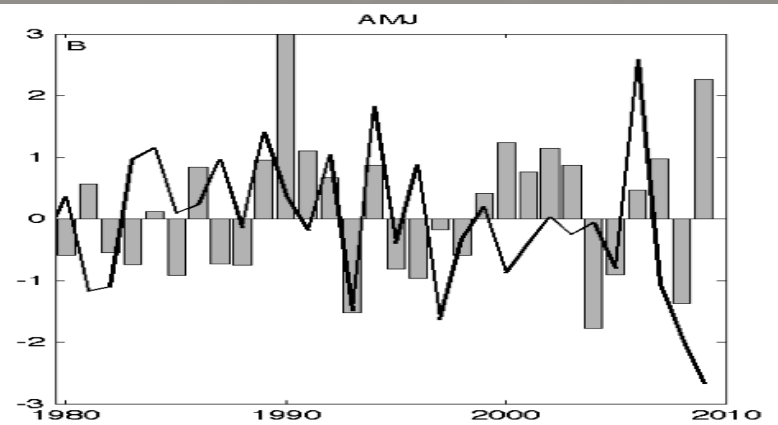
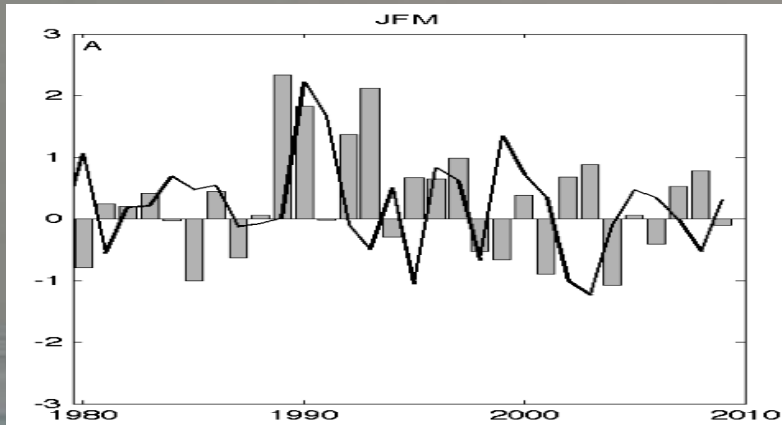
The Arctic Oscillation (AO) Northern Annular Mode (NAM)



Arctic Oscillation & Dipole Anomaly



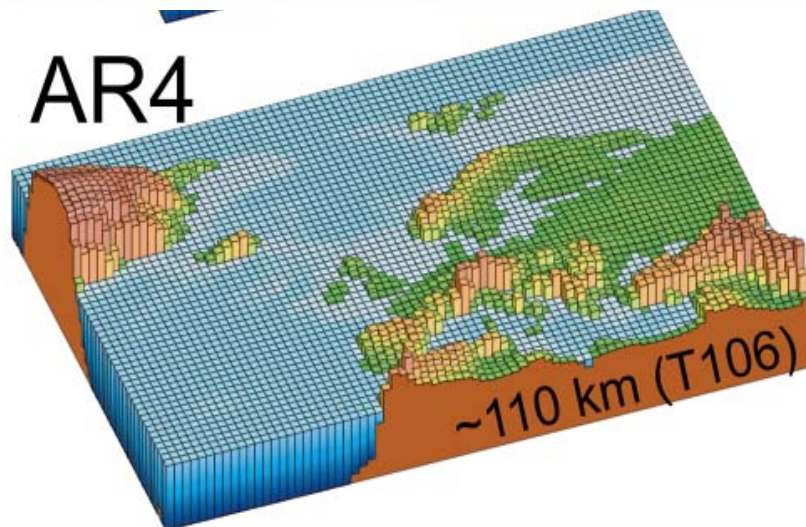
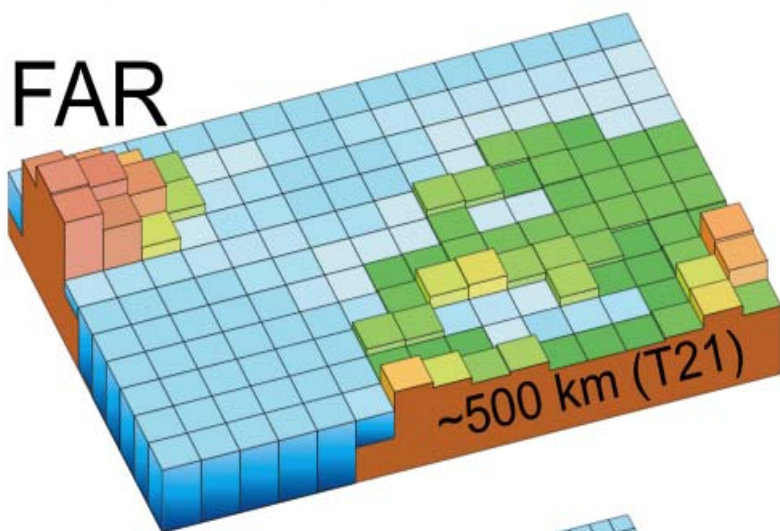
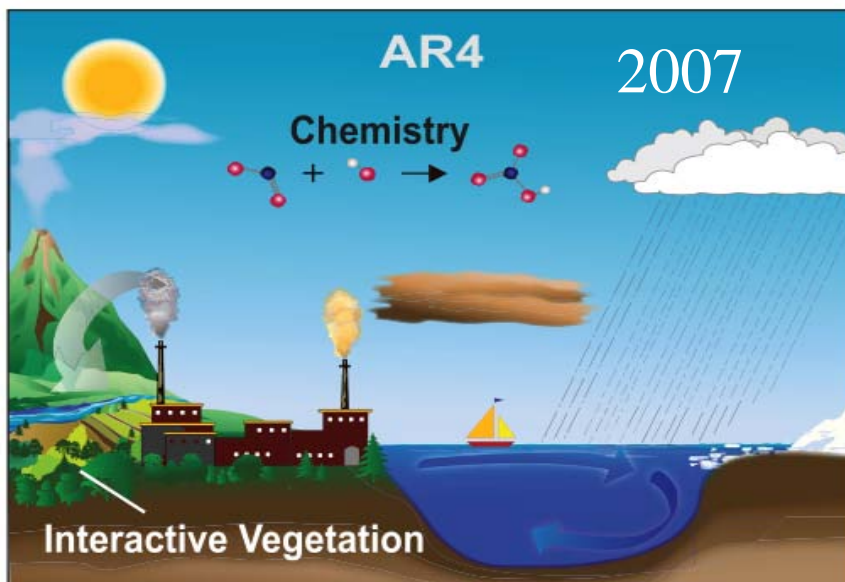
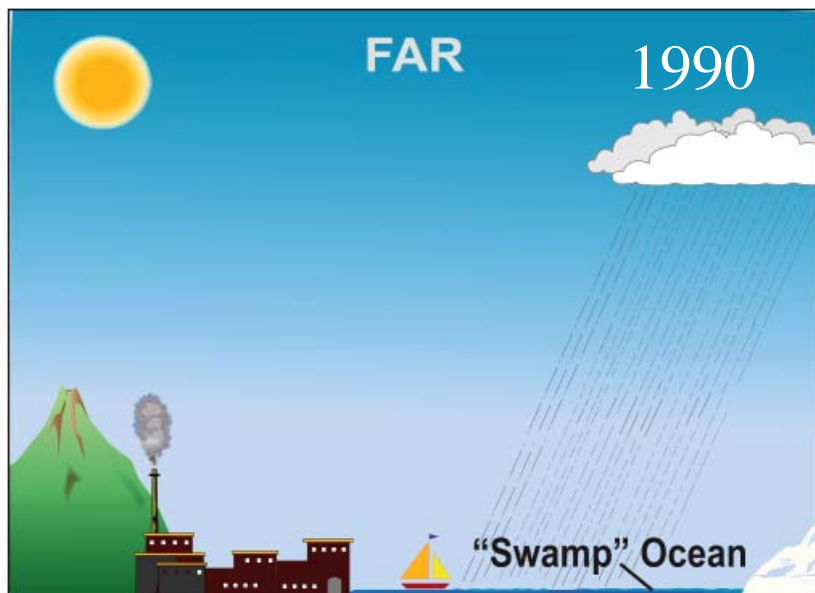
Seasonal Mean PC Series



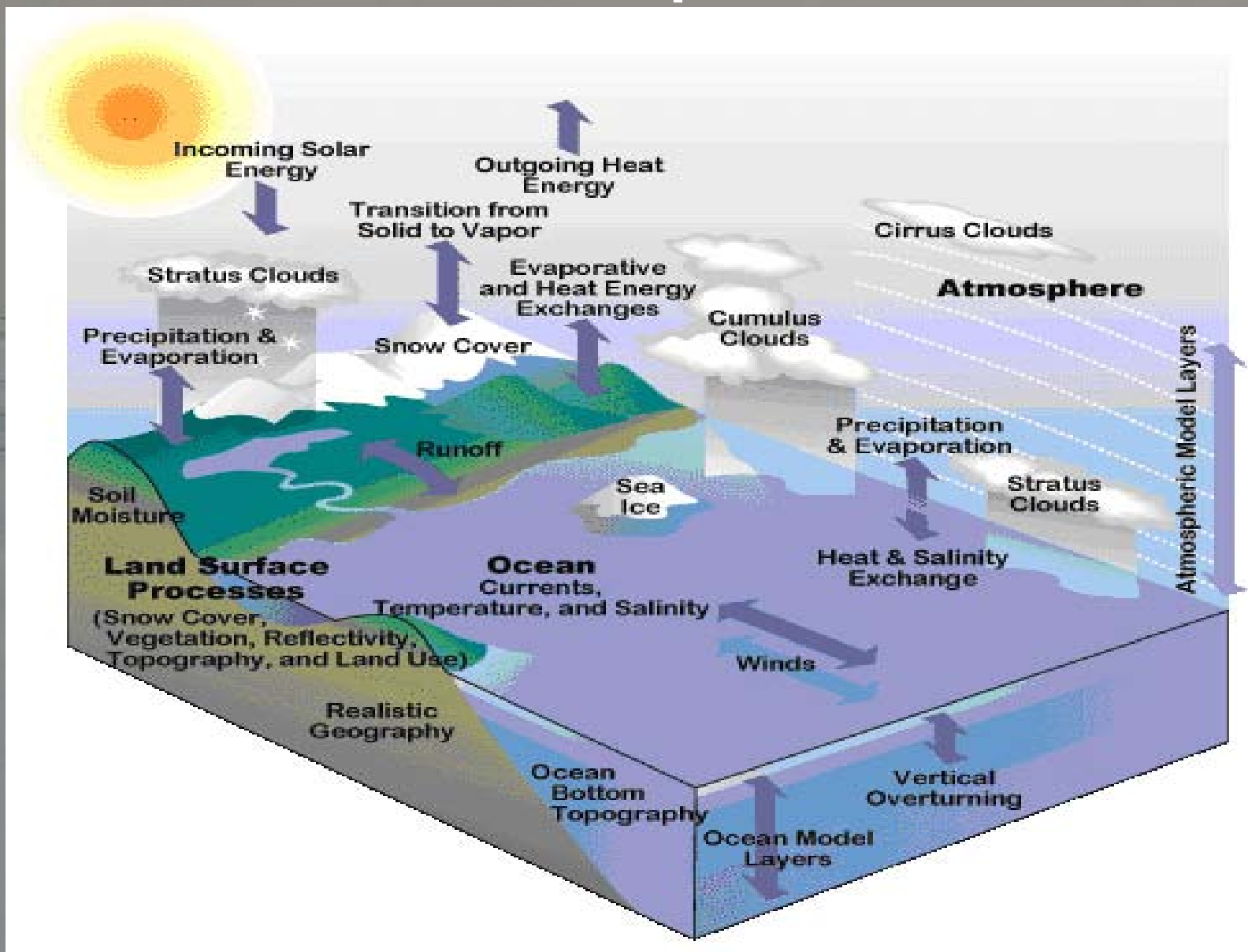
OUTLINE

- The observed changes in the Arctic
- Feedback mechanisms
- Model culling and projections of future Arctic Sea ice condition

The Progress of Climate Models



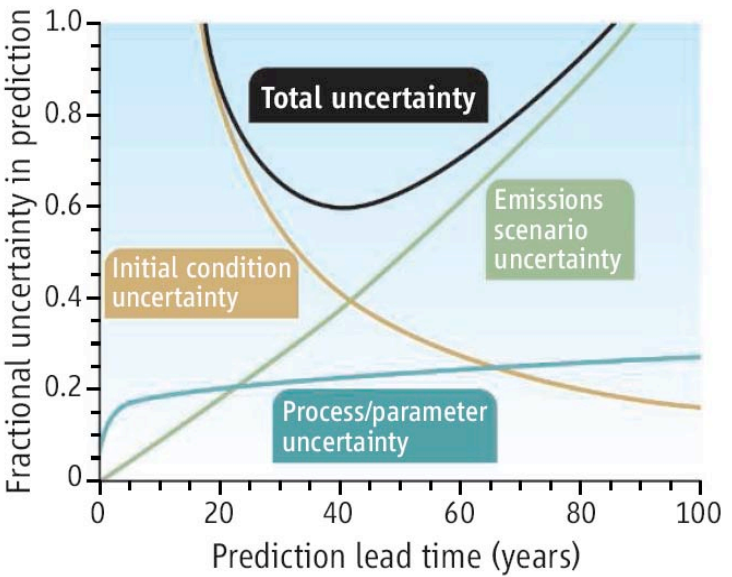
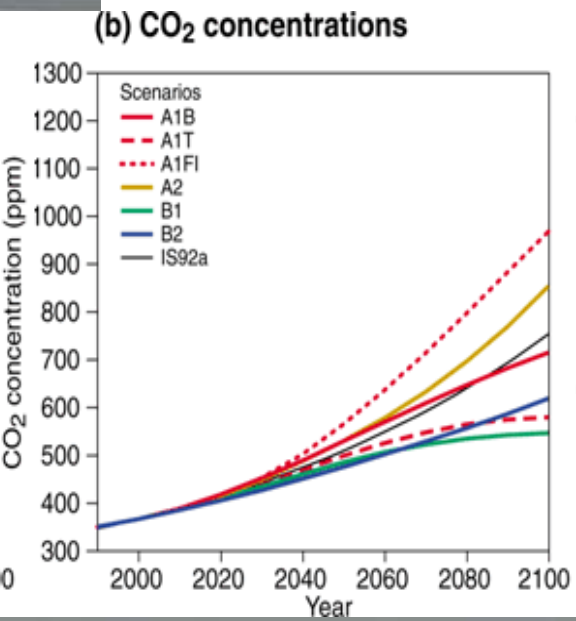
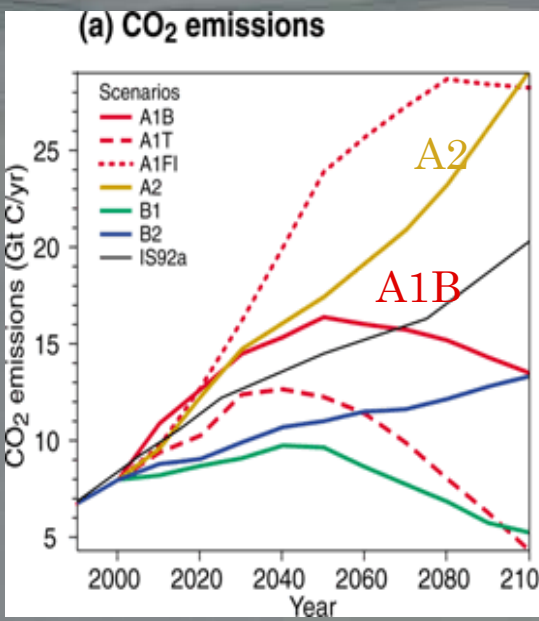
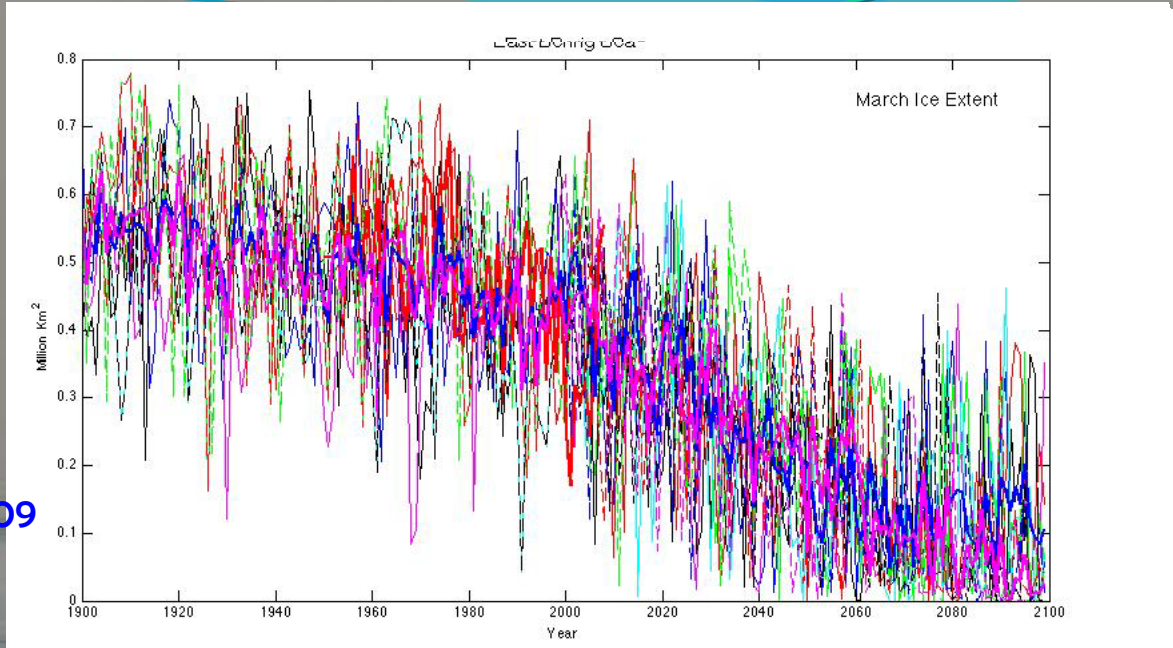
Model Components



Source of Projection Uncertainties

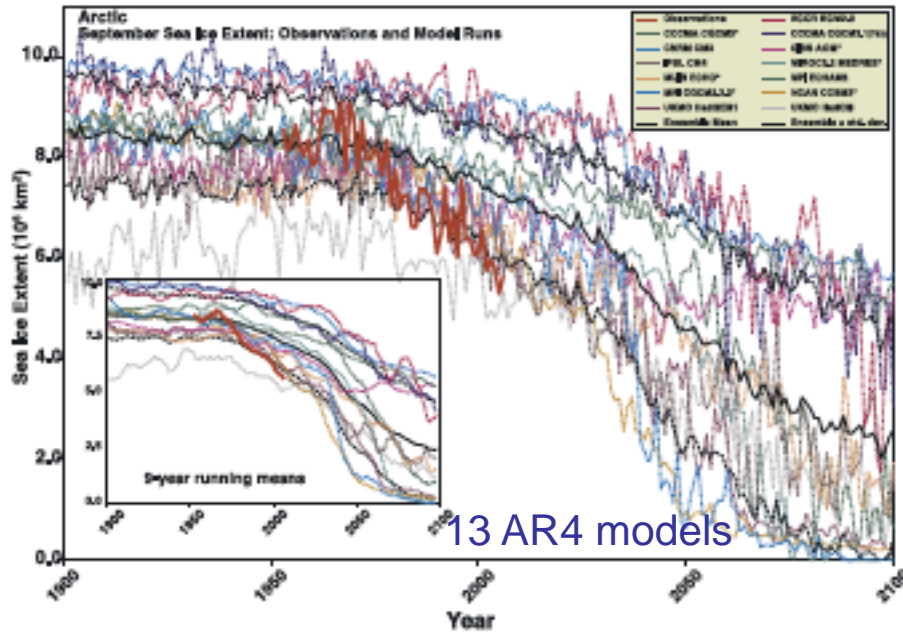
- Internal Variability of climate system
- Model Uncertainty
- Emission Scenarios

Hawkins and Sutton, 2009

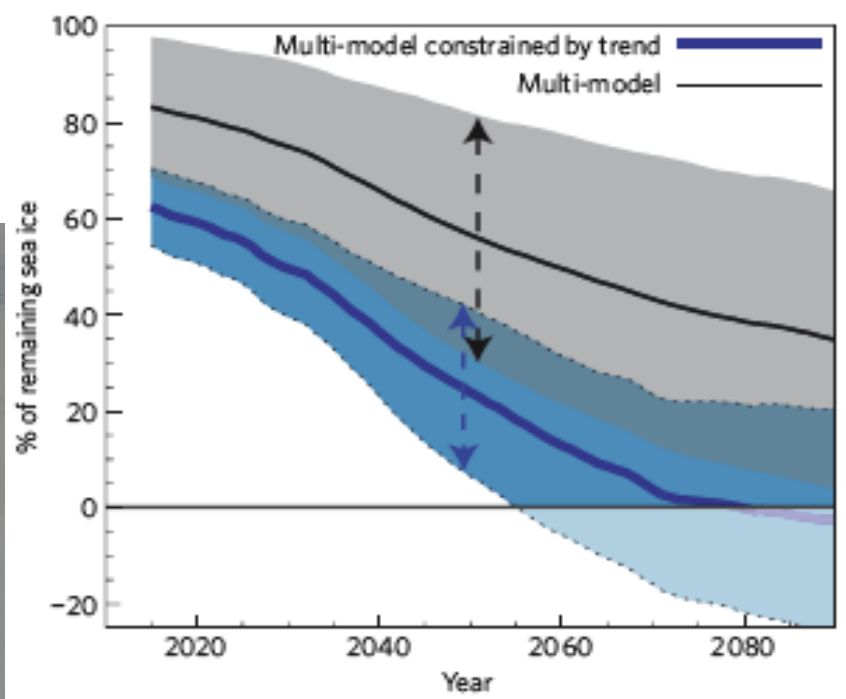


Cox and Stephenson, 2007

Model Culling

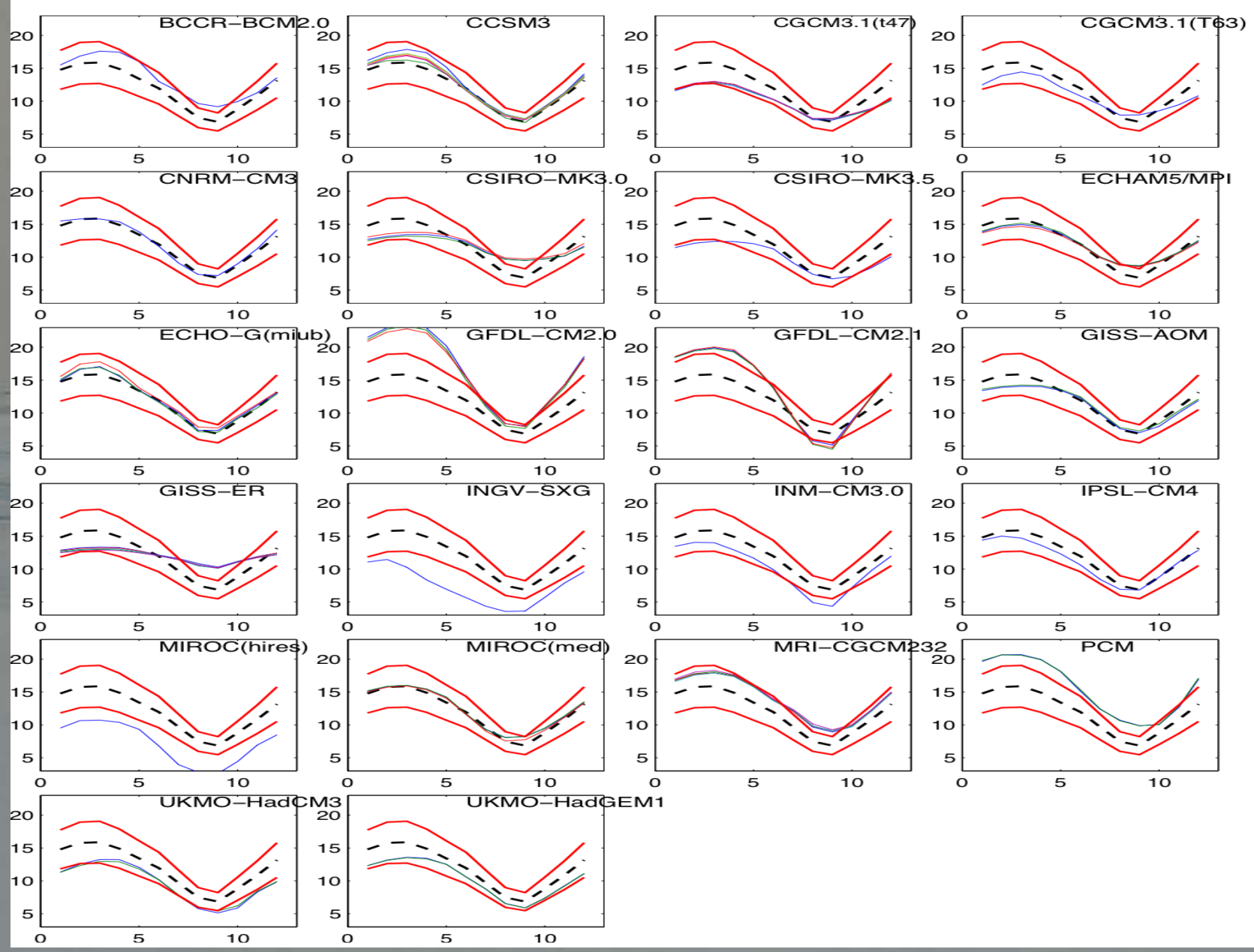


Stroeve et al. 2007

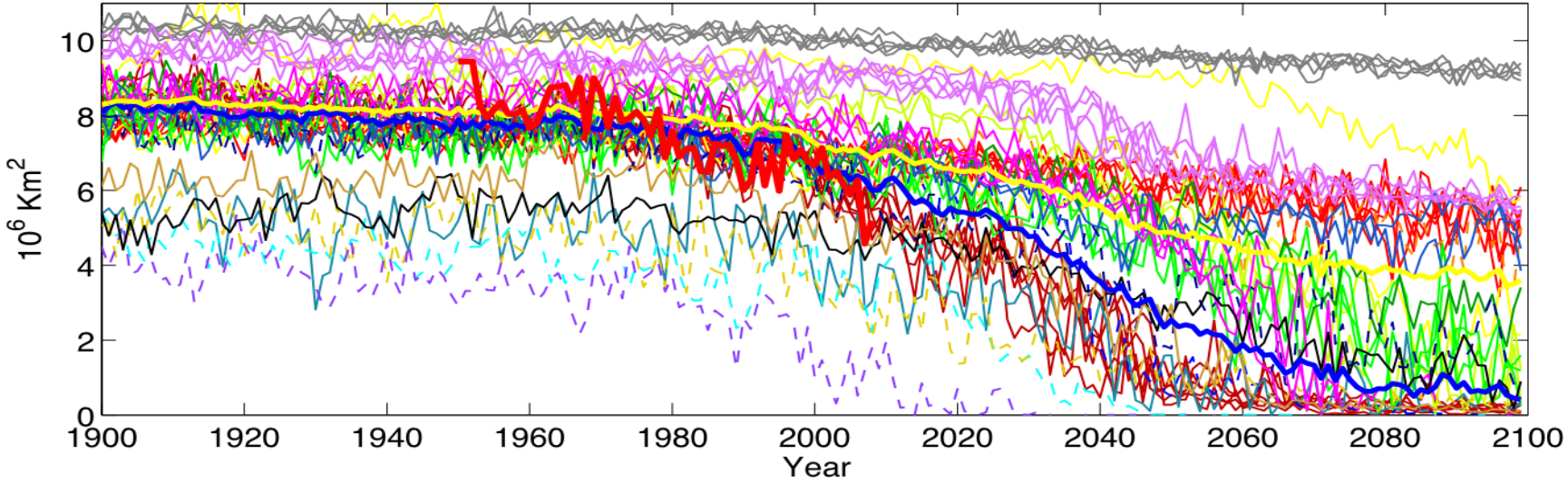


Boe et al. 2009

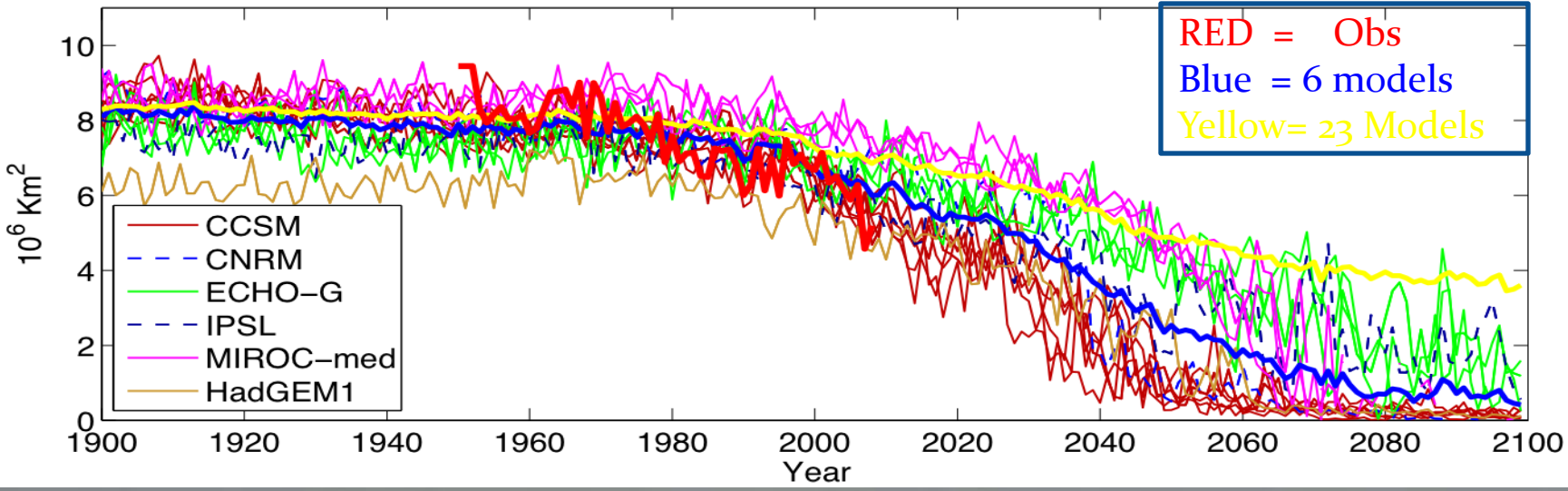
Climatology/Seasonal Cycle



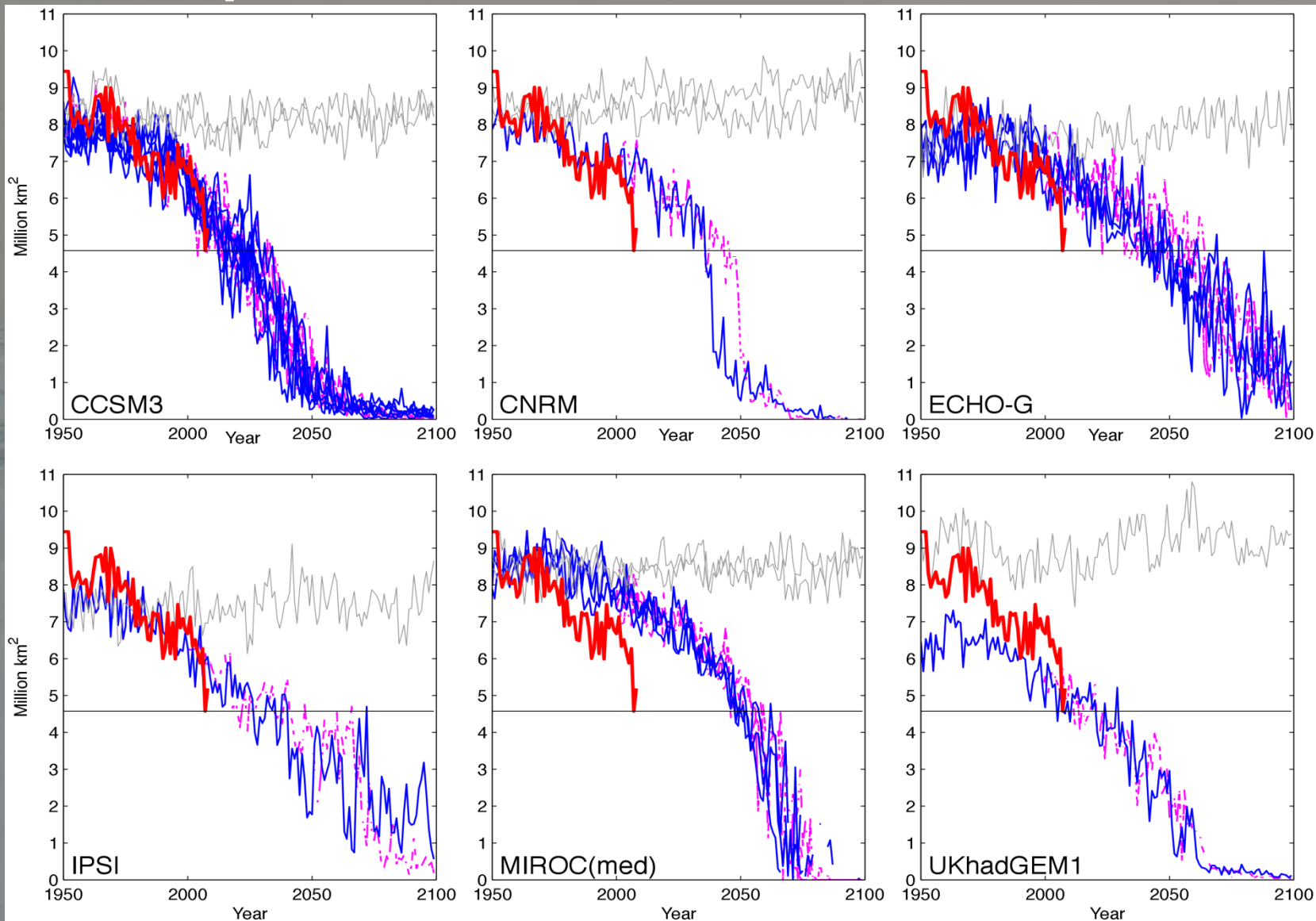
NH Sept Sea Ice Extent

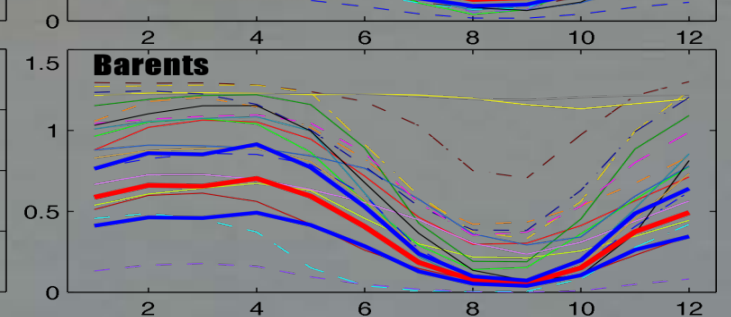
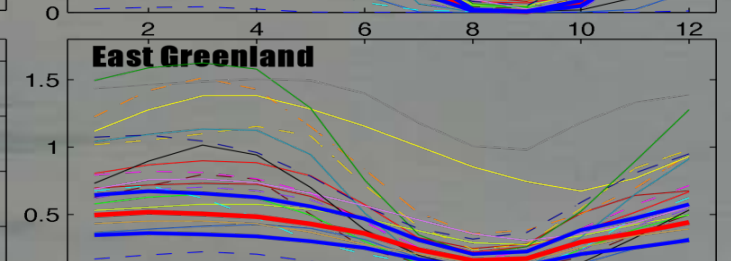
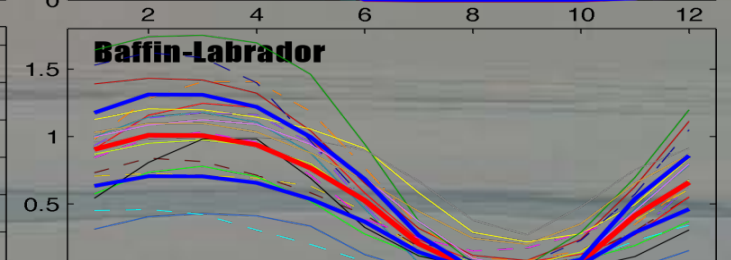
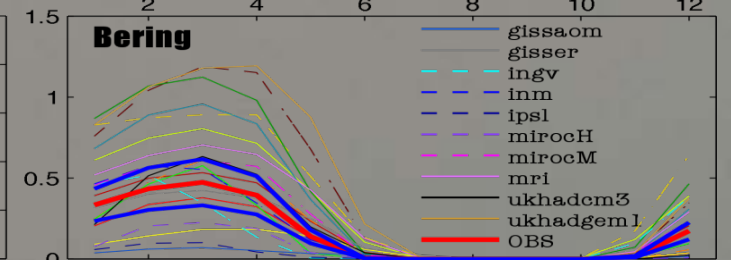
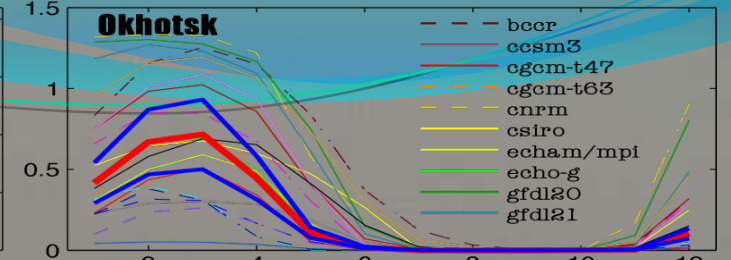
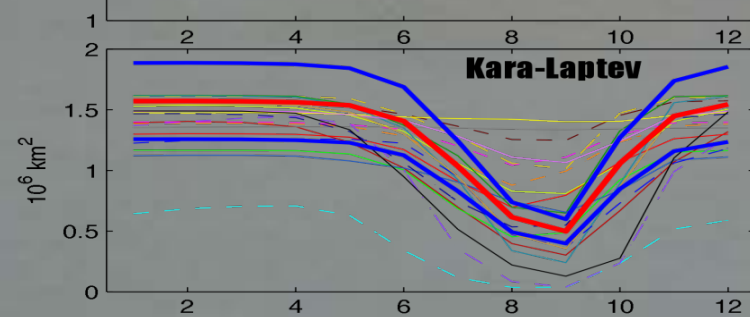
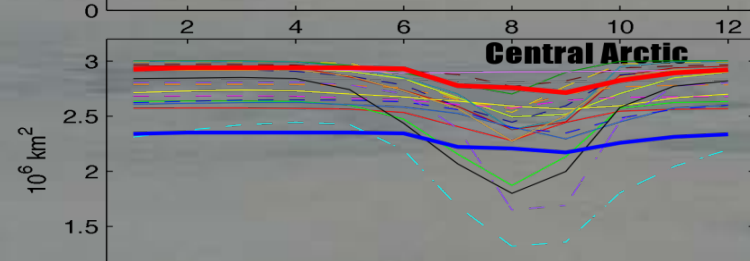
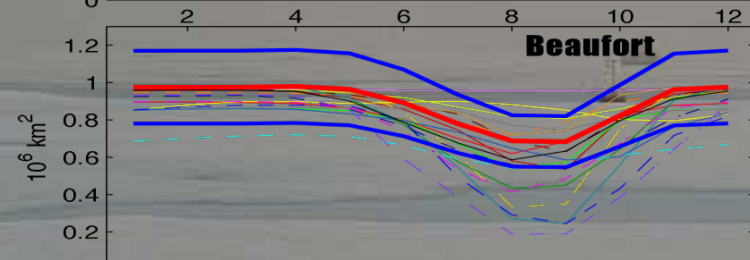
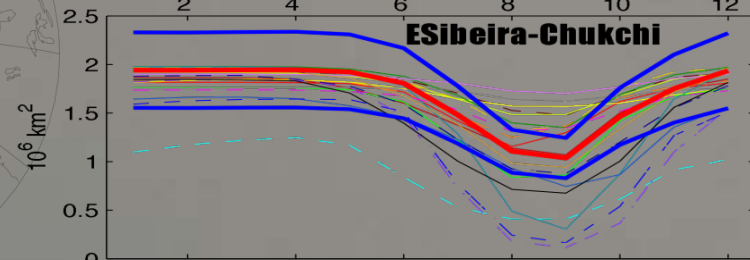
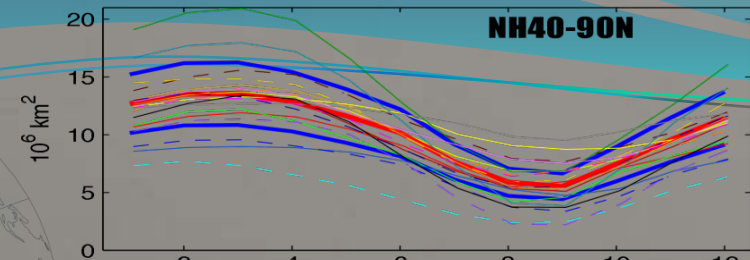
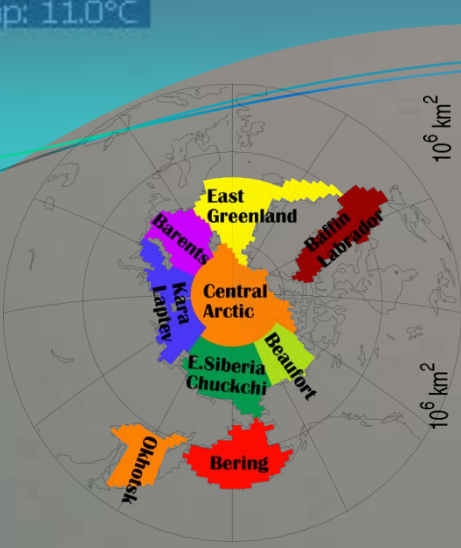


A1B Sep NH Ice Extent

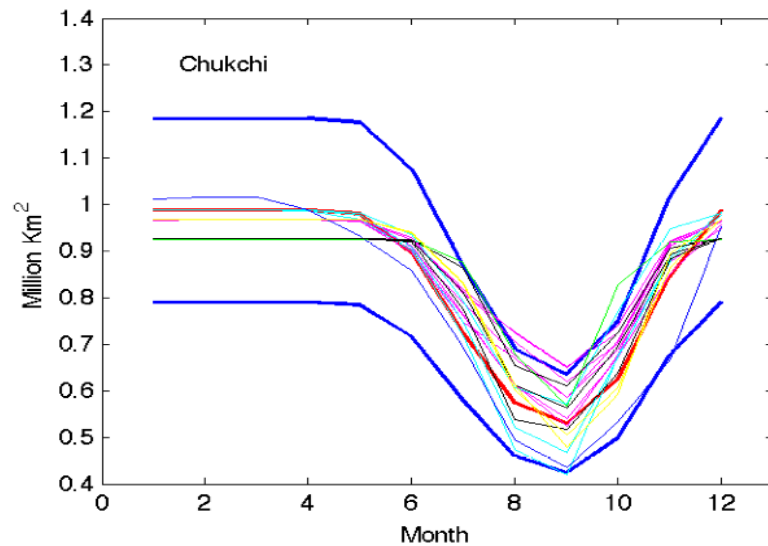
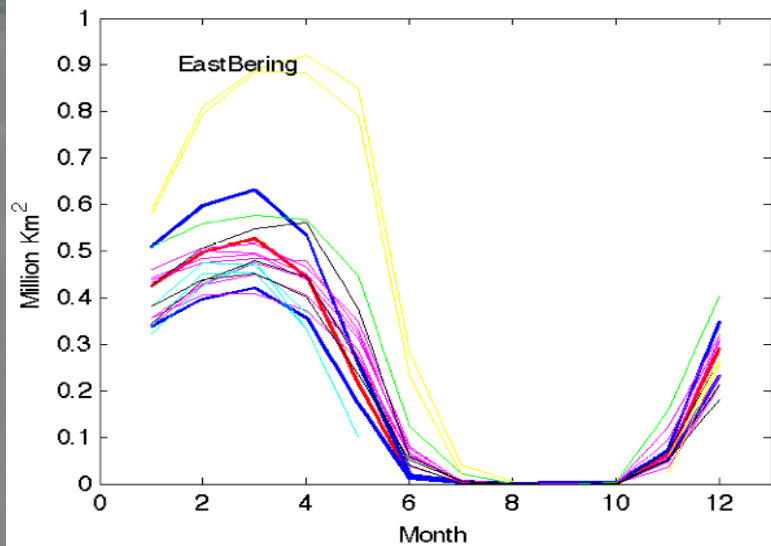
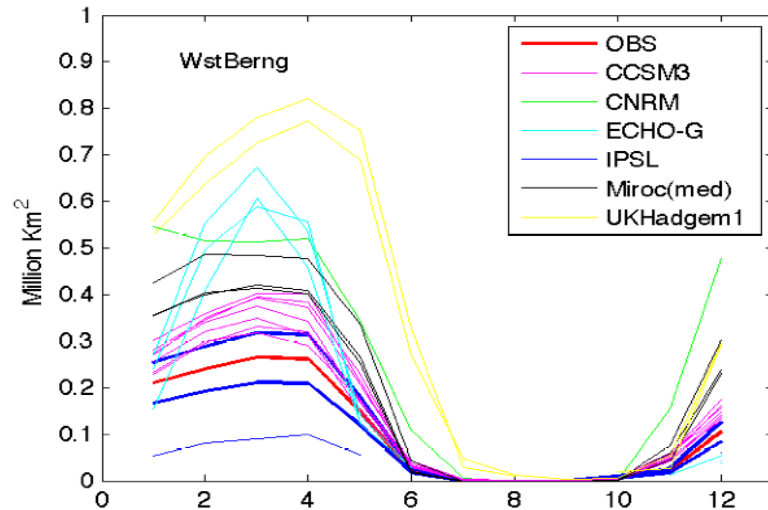
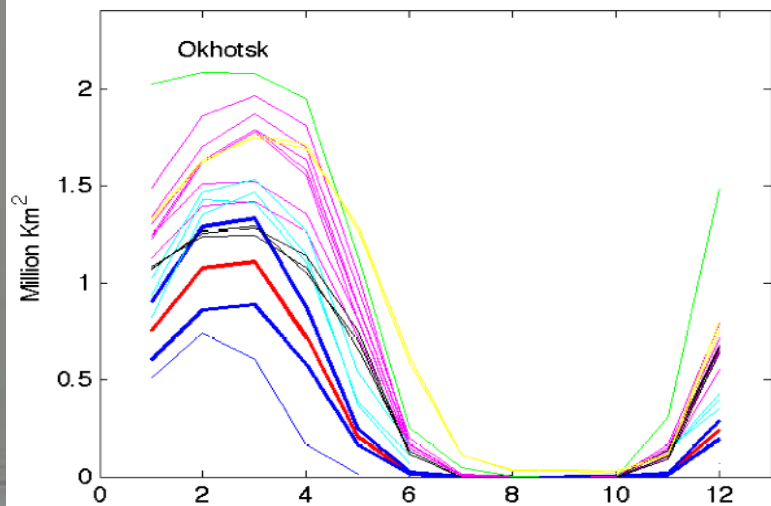


September Sea Ice Extent

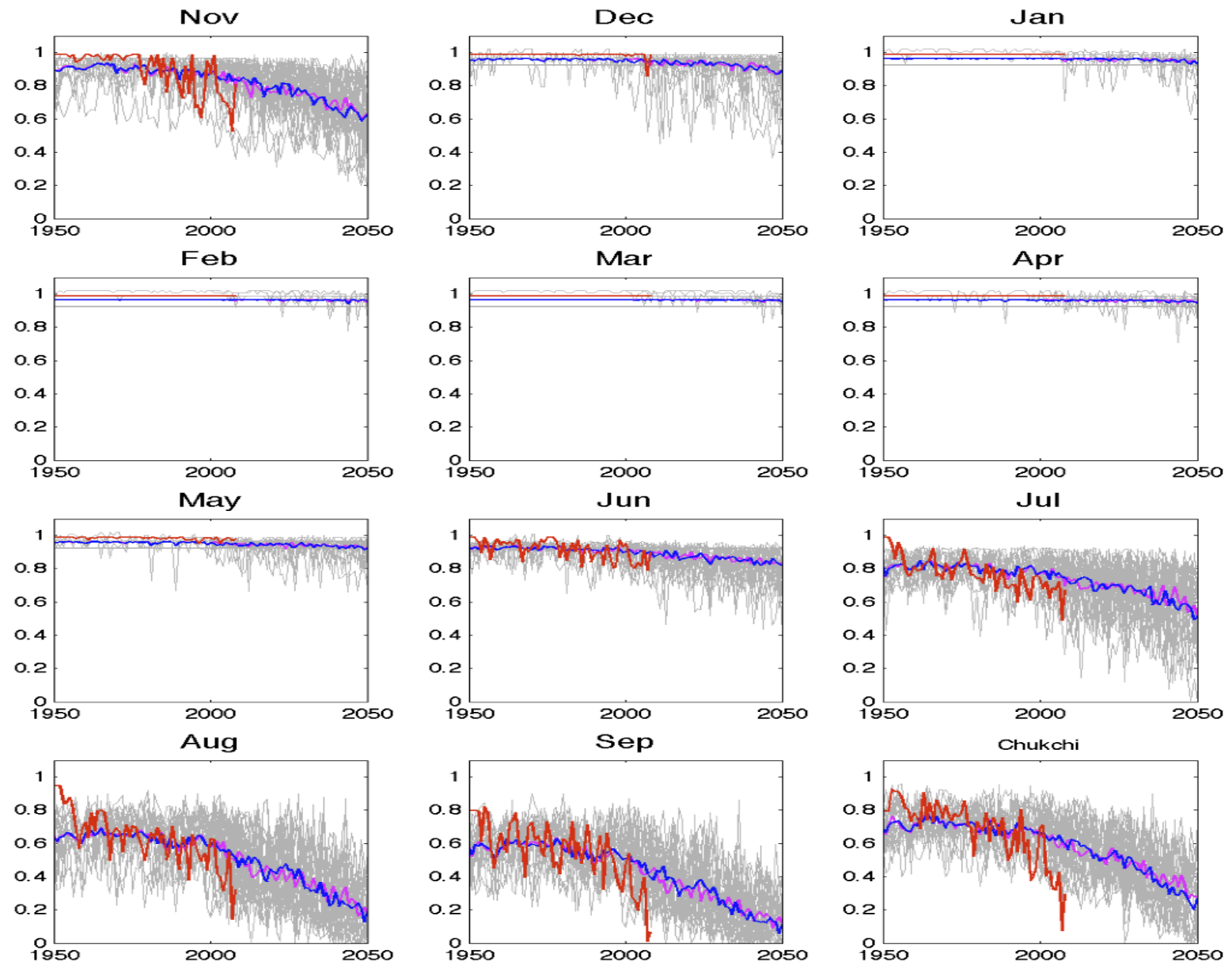




Seasonal Cycle of Sea Ice Extents

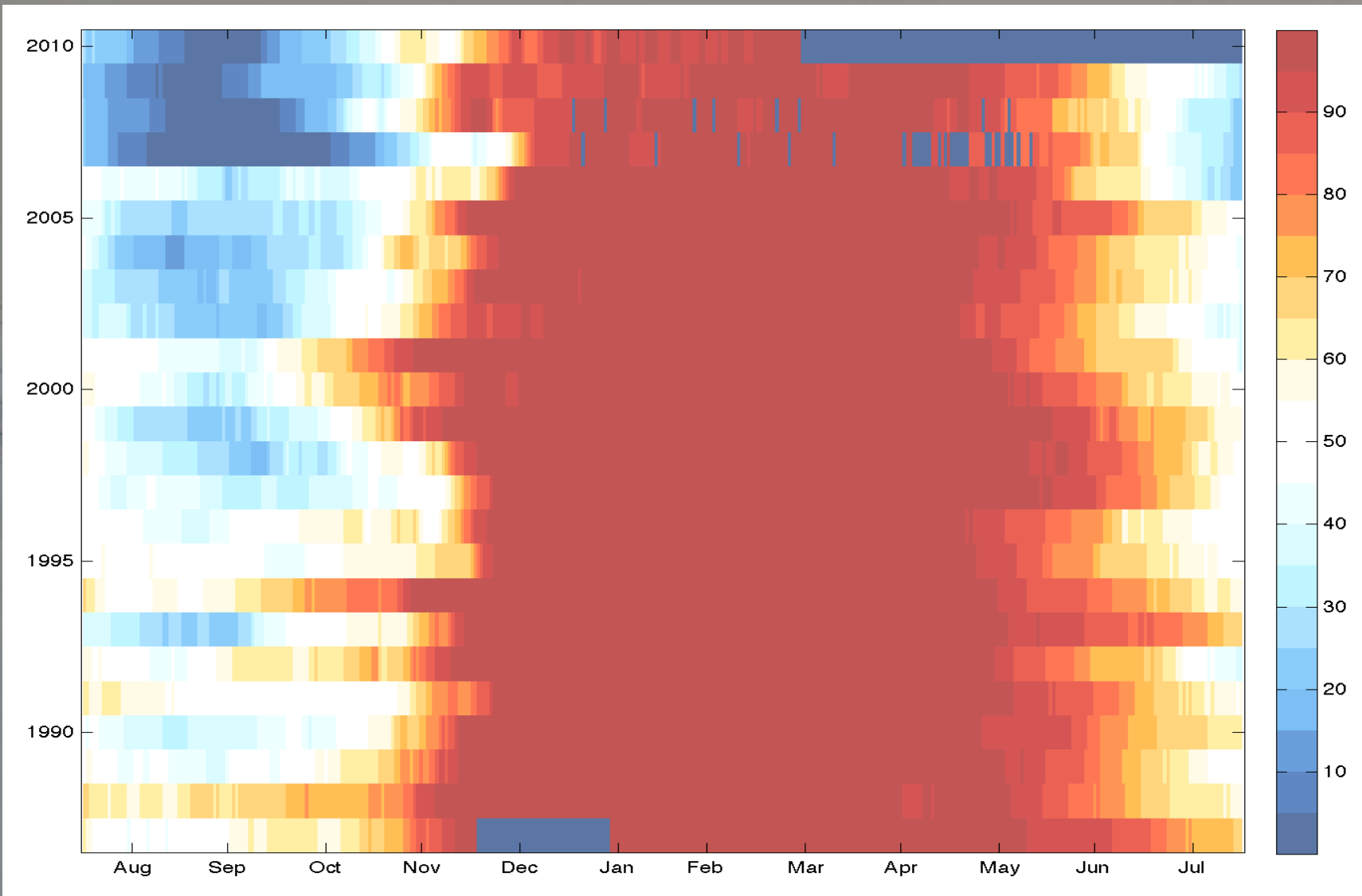


Monthly Sea Ice Extent over Chukchi Sea

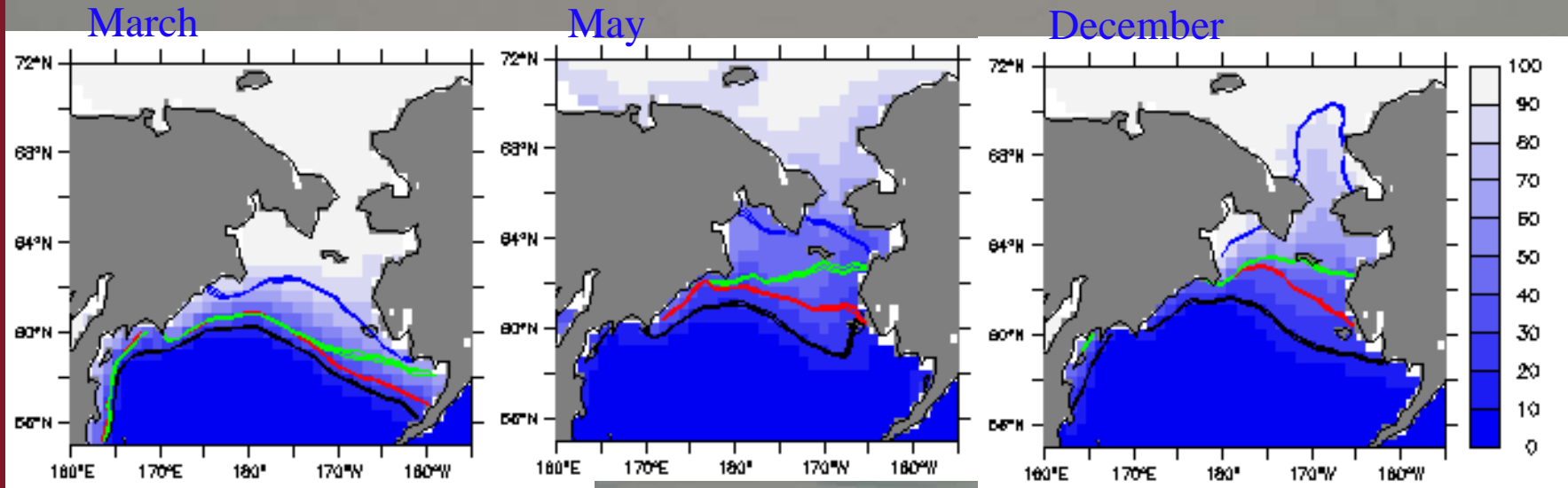


Mean Sea Ice Concentration over the Chukchi Sea

65-80N, 180-157W



DECADAL MEAN ICE EDGE under A1B scenario

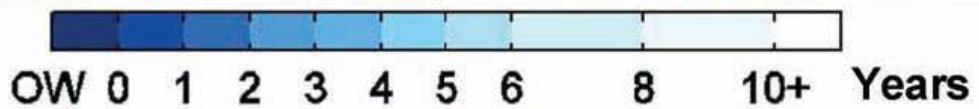
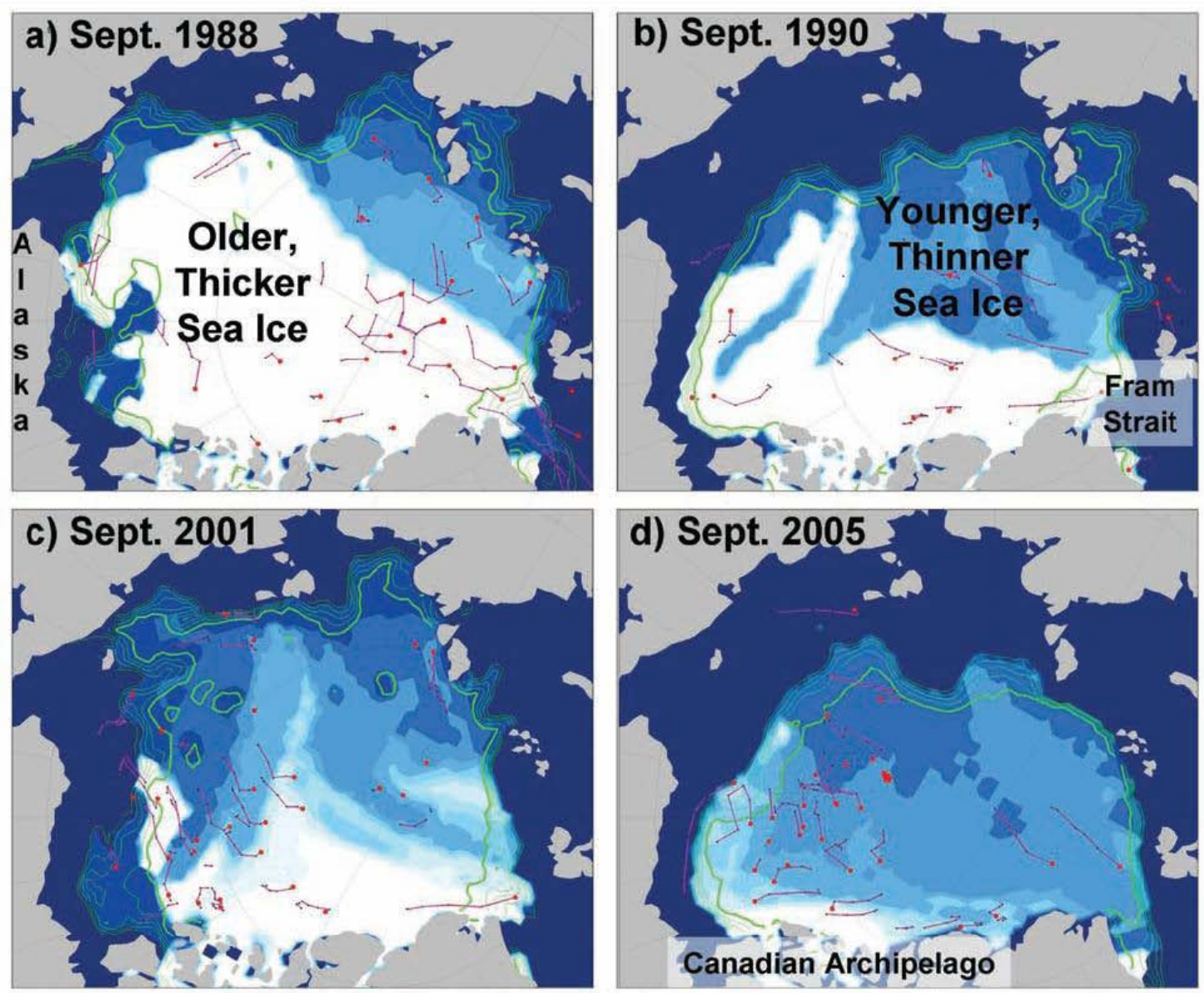


— Current — 2011-19 — 2031-39 — 2051-59

CCSM3



Buoy drift and Ice Concentration



Ice Export Faster ...

STEADY AS SHE FLOWS

In 1896, Fram became the first vessel to have ridden the Transpolar Drift Stream - one of the Arctic's ice currents. This year, Tara was the second, making the journey in less than half the time. The Transpolar Drift Stream is pushed along by westerly winds, while the other major ice current in the Arctic is the clockwise-circulating Beaufort Gyre, generated by the rotating winds created by a high-pressure atmospheric system





Arctic Report Card: Update for 2010

Tracking recent environmental changes

Home Atmosphere Sea Ice Ocean Land Greenland Biology

Return to previous Arctic conditions is unlikely

Record temperatures across Canadian Arctic and Greenland, a reduced summer sea ice cover, record snow cover decreases and links to some Northern Hemisphere weather support this conclusion

Arctic Report Card 2010



■ Atmosphere
 ■ Biology
 ■ Greenland
■ Sea Ice
 ■ Ocean
 ■ Land

Red boxes: Consistent evidence of warming.
Yellow boxes: Many indications of warming.

Atmosphere

Arctic climate is impacting mid-latitude weather, as seen in Winter 2009-2010

Sea Ice

Summer sea ice conditions for previous four years well below 1980s and 1990s

Ocean

Upper ocean showing year-to-year variability without significant trends

Land

Low winter snow accumulation, warm spring temperatures lead to record low snow cover duration

Greenland

Record setting high temperatures, ice melt, and glacier area loss

Biology

Rapid environmental change threatens to disrupt current natural cycles

About the Report Card

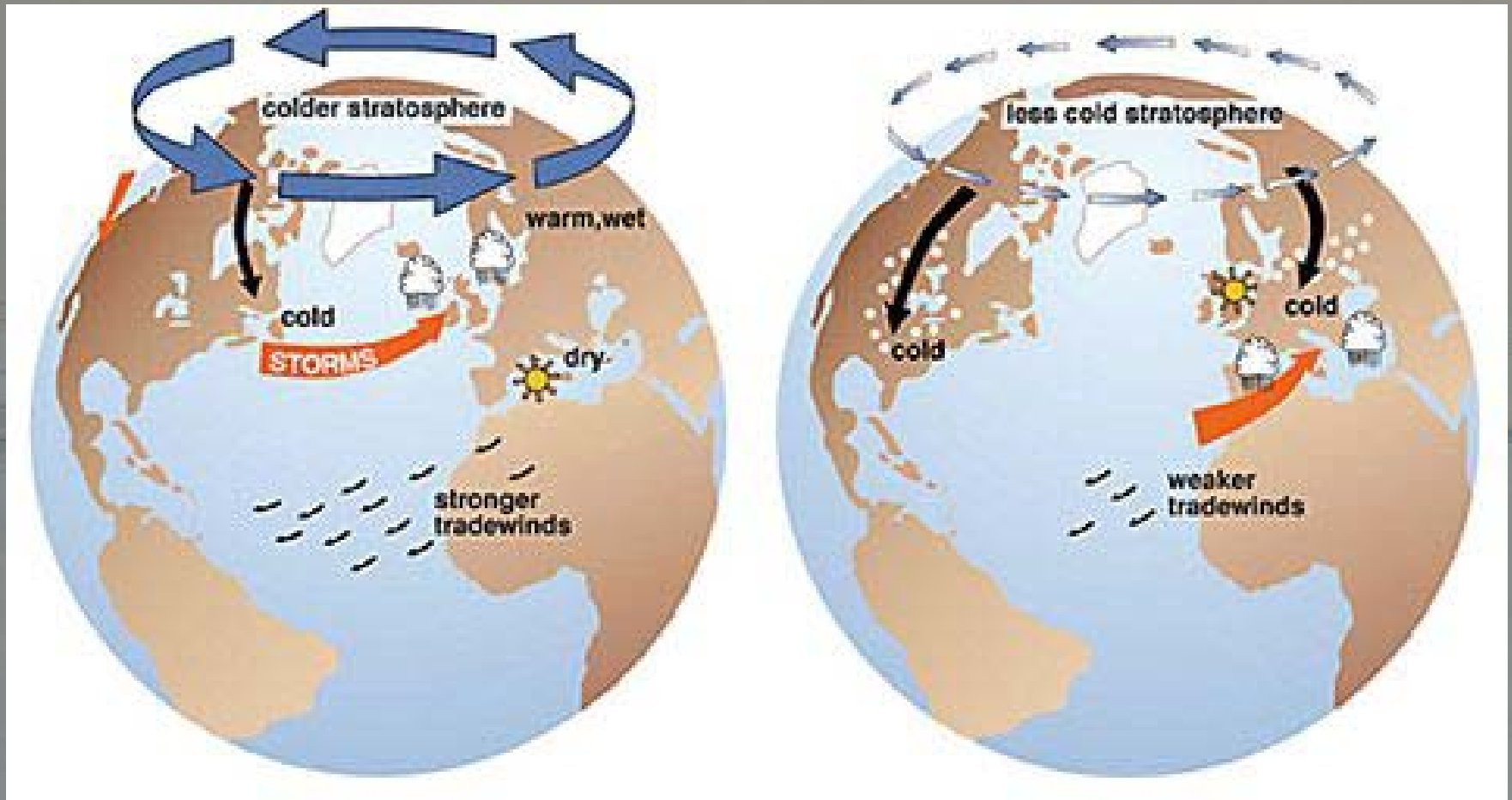
Printable Handout :: Executive Summary :: Full Arctic Report Card (PDF)
NOAA Arctic Theme Page



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Disclaimer | Privacy Policy | Webmaster

<http://www.arctic.noaa.gov/reportcard/>

The AO and Weather Pattern



Courtesy of J. Wallace

Toward producing reanalysis wind field over the Chukchi/Beaufort Seas via data assimilation and analysis nudging

Jing Zhang¹, Jeremy Krieger², Fuhong Liu¹,
Martha Shulski³, Xiangdong Zhang⁴

¹NOAA ISET Center, NC A&T State University,

²ARSC, UAF,

³High Plains Regional Climate Center, UNL

⁴IARC, UAF

Introduction

- When performing long-term regional simulations, techniques must be used to constrain the model and maintain numerical stability.
- Data assimilation and analysis nudging are two powerful methods for improving model performance.
- Data assimilation introduces observed, high resolution information into the model simulation, increasing accuracy.
- Analysis nudging constrains the model solution and prevents errors from growing too large.

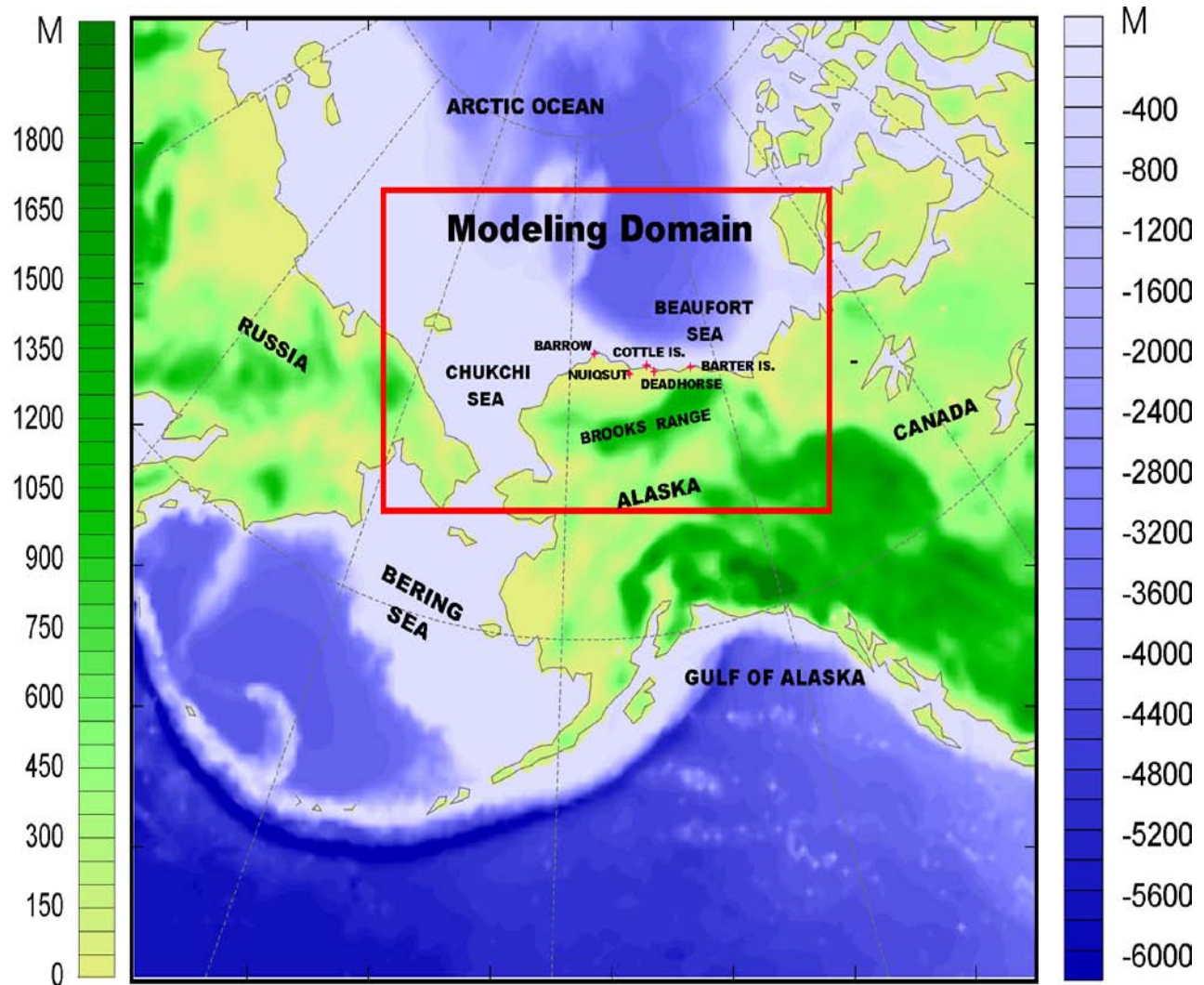
Model Configuration

Horizontal Grid
Spacing: **10 km**

49 vertical levels

test periods:

Jun 30 – Aug 30



A little about data assimilation

- Assume:

$$T_m = 18^\circ \text{C} \text{ (model temperature)}$$

$$T_o = 21^\circ \text{C} \text{ (observed temperature)}$$

T_t true temperature, but we don't know...

- Assume:

$$\sigma_m = 2^\circ \text{C} \text{ (model error)}$$

$$\sigma_o = 1^\circ \text{C} \text{ (observational error)}$$

σ_m & σ_o are uncorrelated

- T will be corrected as:

$$T = a T_m + b T_o$$

- The most straightforward way to decide a & b is to minimize the mean square error of T:

$$E[(T - T_t)^2] = E[(a(T_m - T_t) + b(T_o - T_t))^2] = a^2 \sigma_m^2 + b^2 \sigma_o^2$$

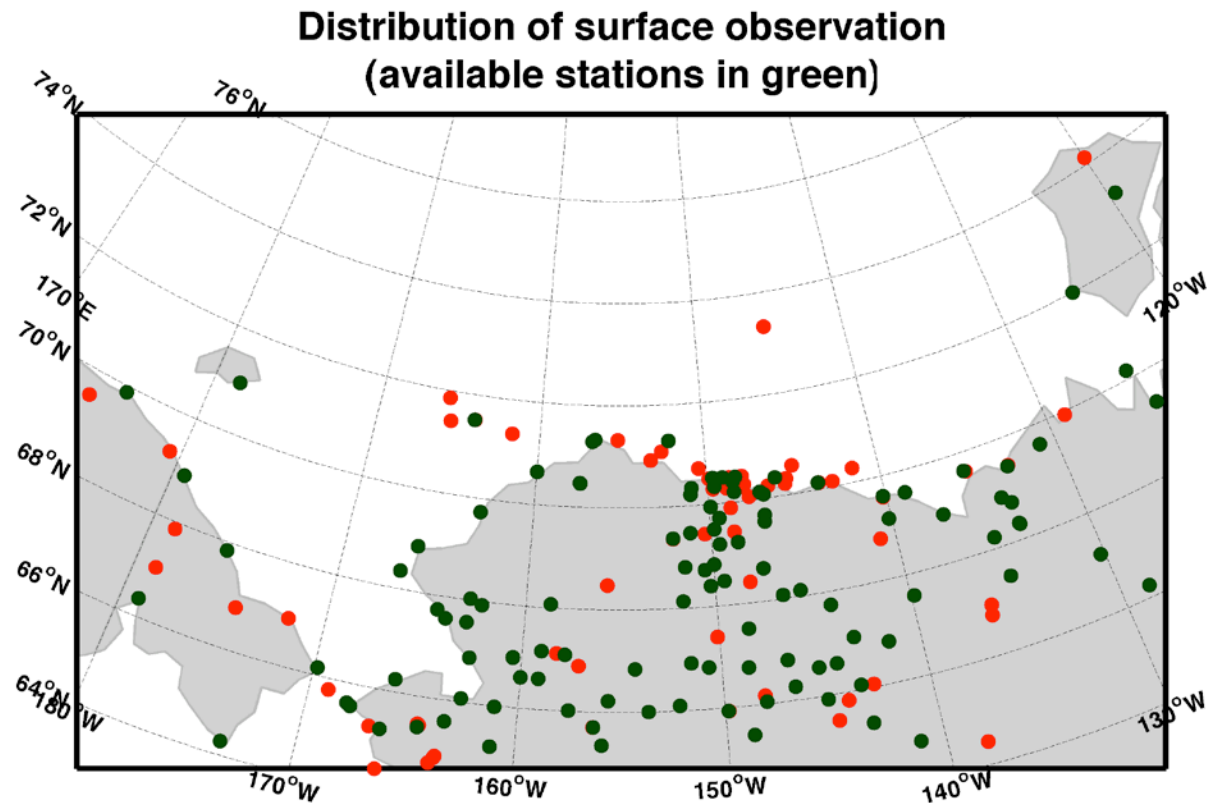
$$a = \frac{\sigma_o^2}{\sigma_o^2 + \sigma_m^2} \quad b = \frac{\sigma_m^2}{\sigma_o^2 + \sigma_m^2}$$

Model Errors

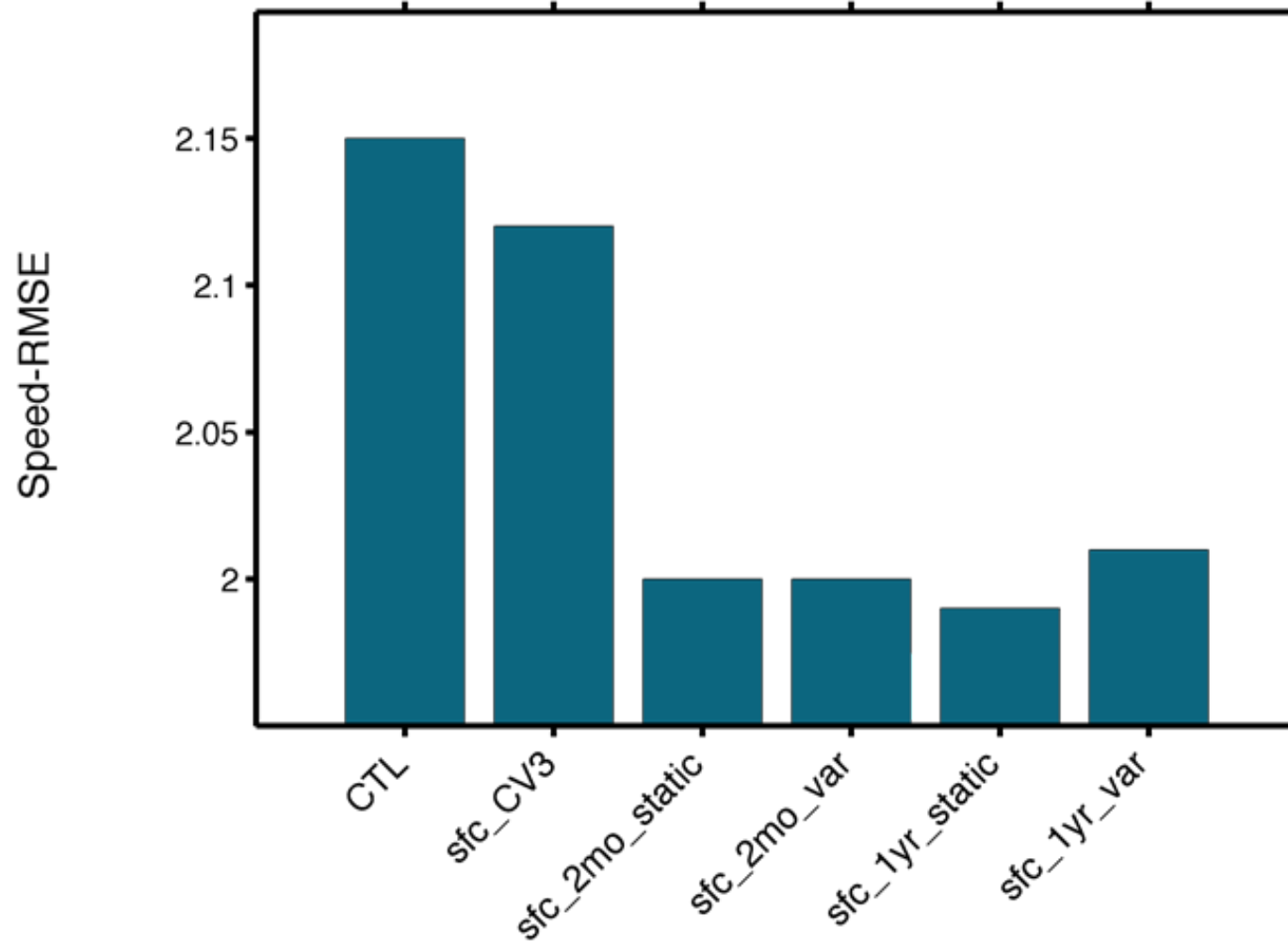
- Several model background errors (BEs) are tested, including
 - CV3 (default)
 - CV5 (customize)
 - 2-month simulation, no diurnal cycle ([sfc_2mo-static](#))
 - 2-month simulation, with diurnal cycle ([sfc_2mo-var](#))
 - 1-year simulation, no diurnal cycle ([sfc_1yr-static](#))
 - 1-year simulation, with diurnal cycle ([sfc_1yr-var](#))

Surface Observation

- Available/Total: 119/195 stations
- Once hourly (most of them)



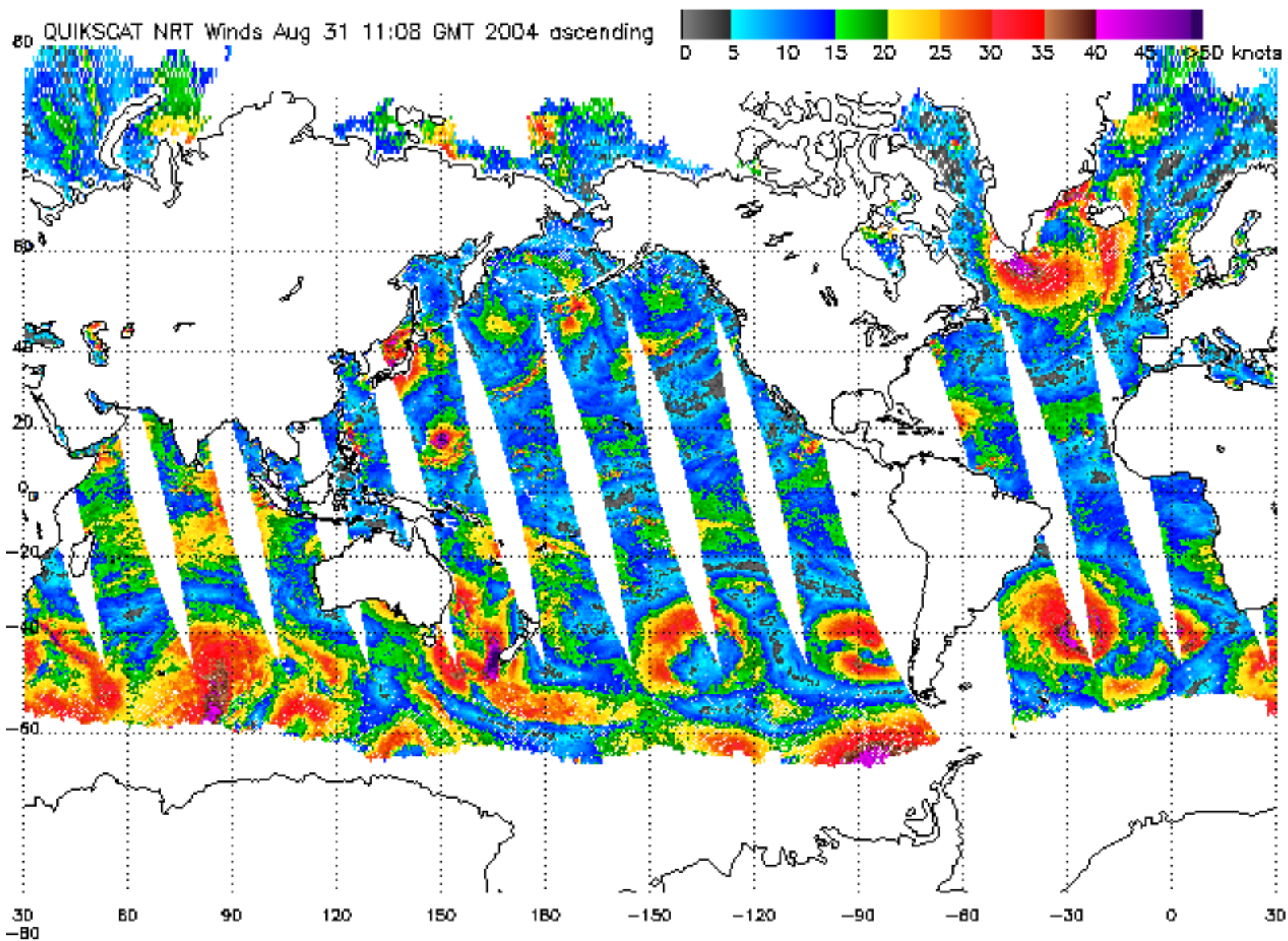
The effects of BE



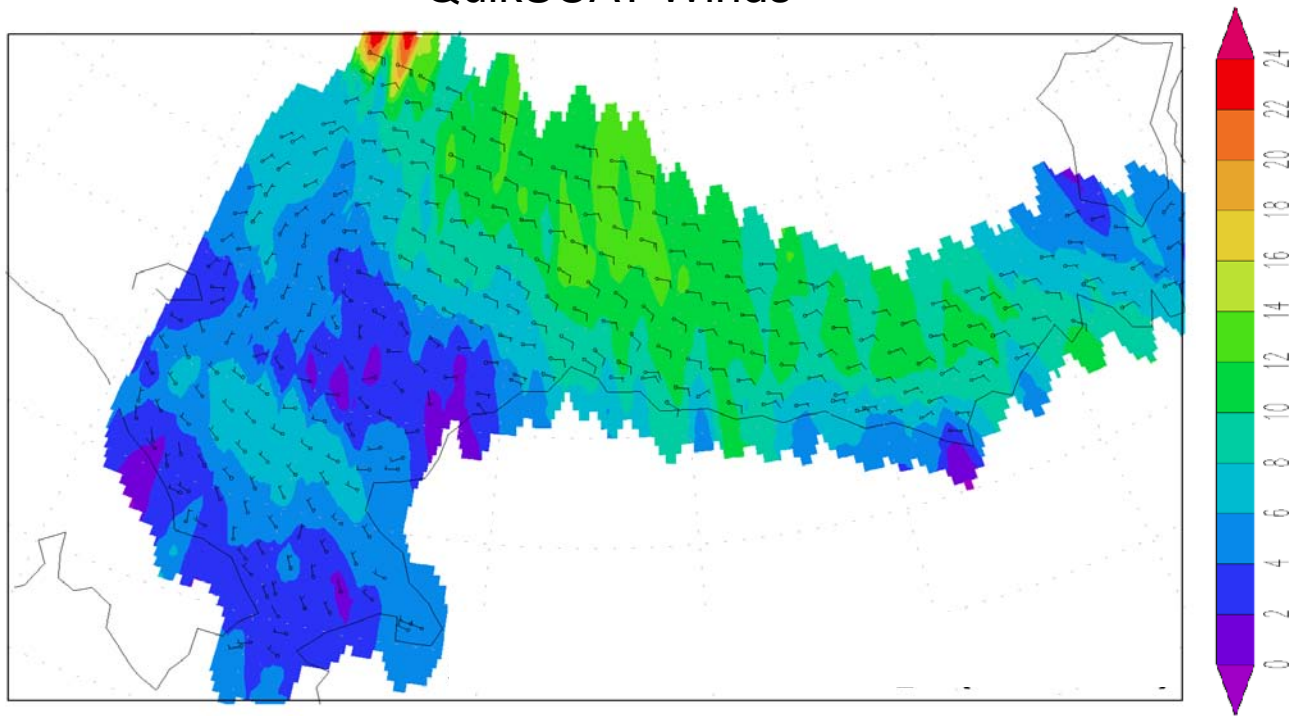
QuikSCAT Winds & Errors

- Ocean surface winds at 10 m height retrieved using observation data from NASA/JPL's SeaWinds scatterometer
- Data is available from **19 July 1999** through **21 November 2009**
- Temporal resolution: multiple times daily
- Spatial resolution: **12.5 km**
- Different Wind speed errors are tested
 - 1 m/s (**qscat_err1**)
 - 4 m/s (**qscat_err4**)

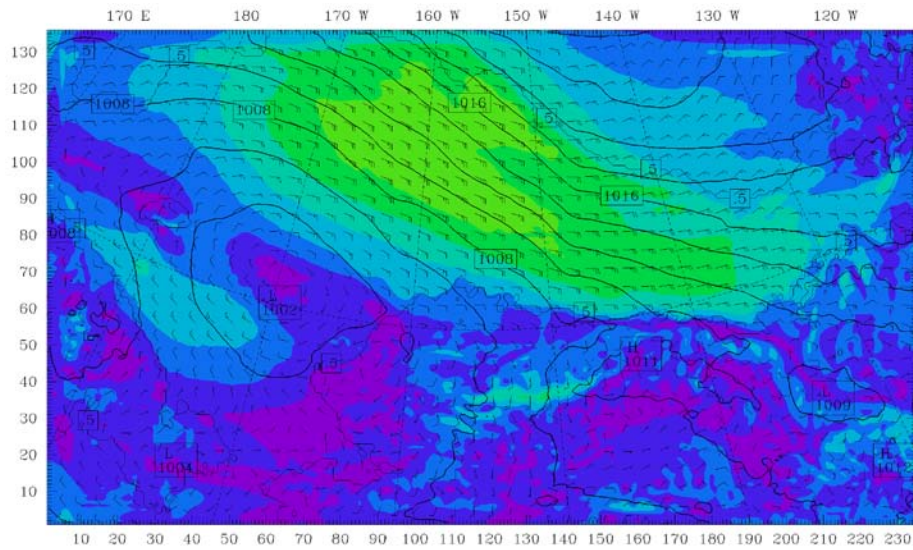
QuikSCAT Winds



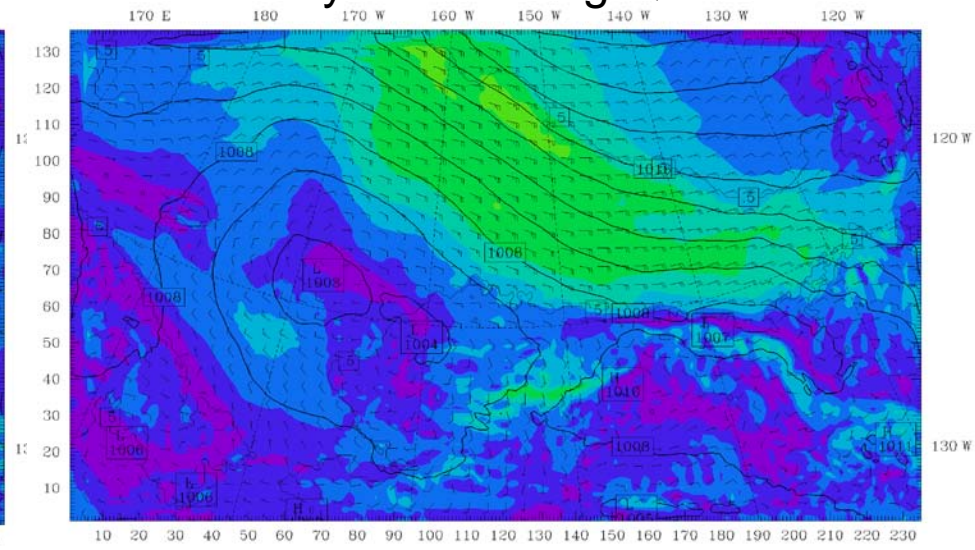
QuikSCAT Winds



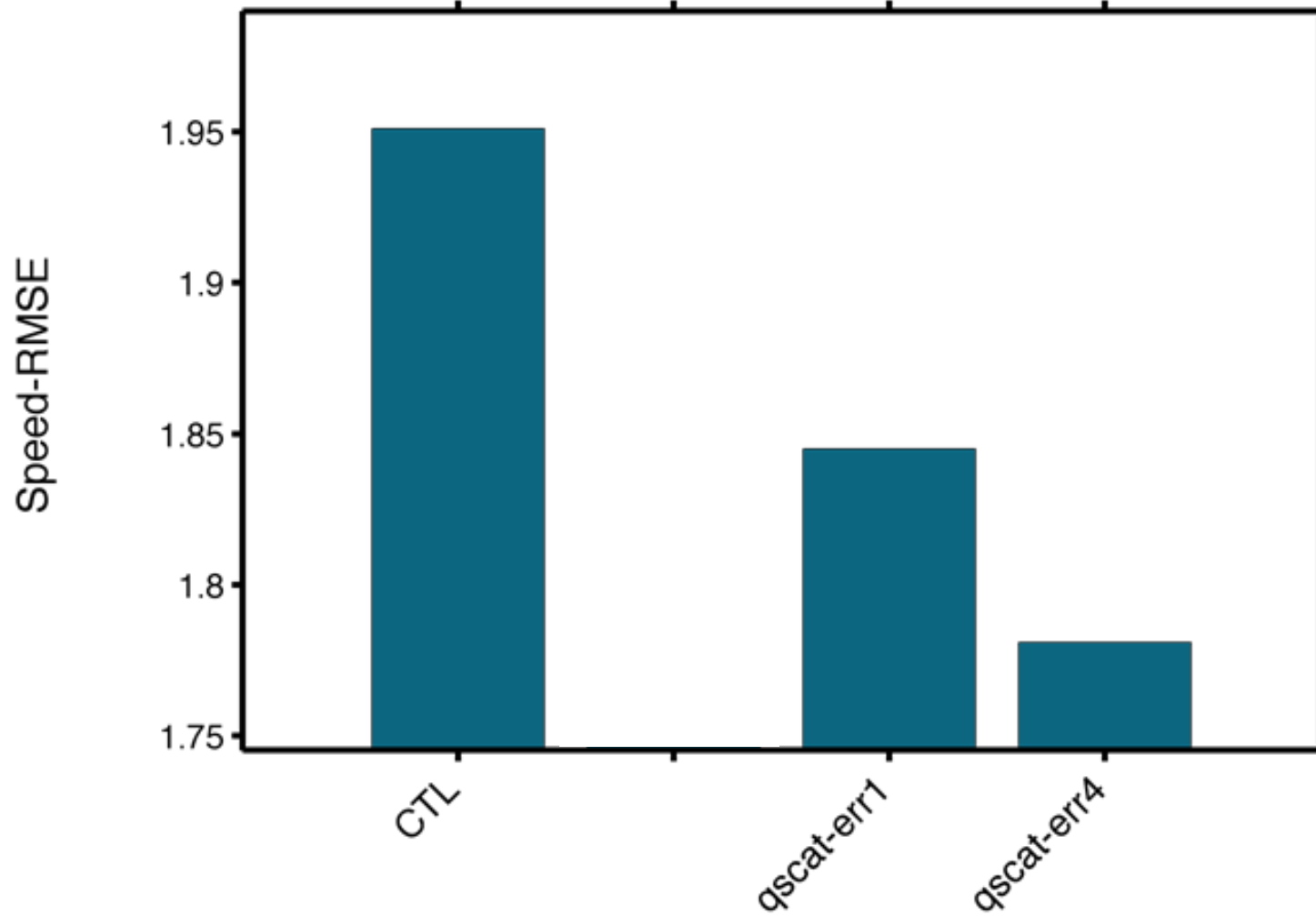
WRF initial winds



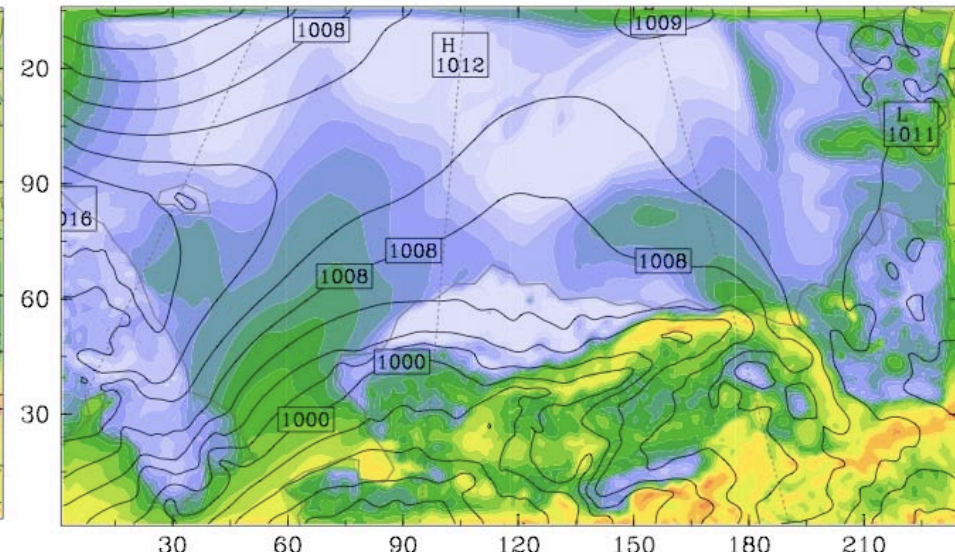
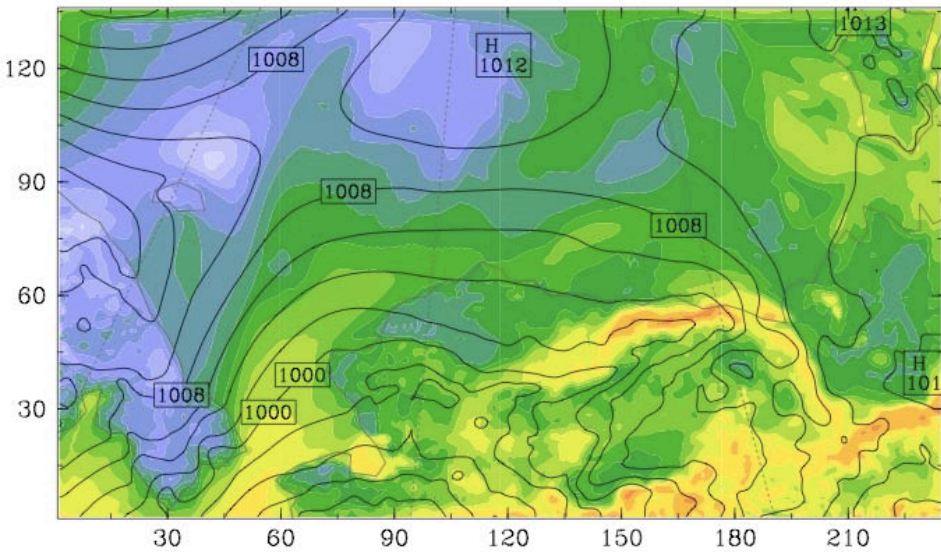
WRF re-initialization by assimilating QuikSCAT winds



The effects of Data Errors



Comparisons between Reanalysis (left) and WRF 1-yr Simulation without Assimilation and Nudging (right)

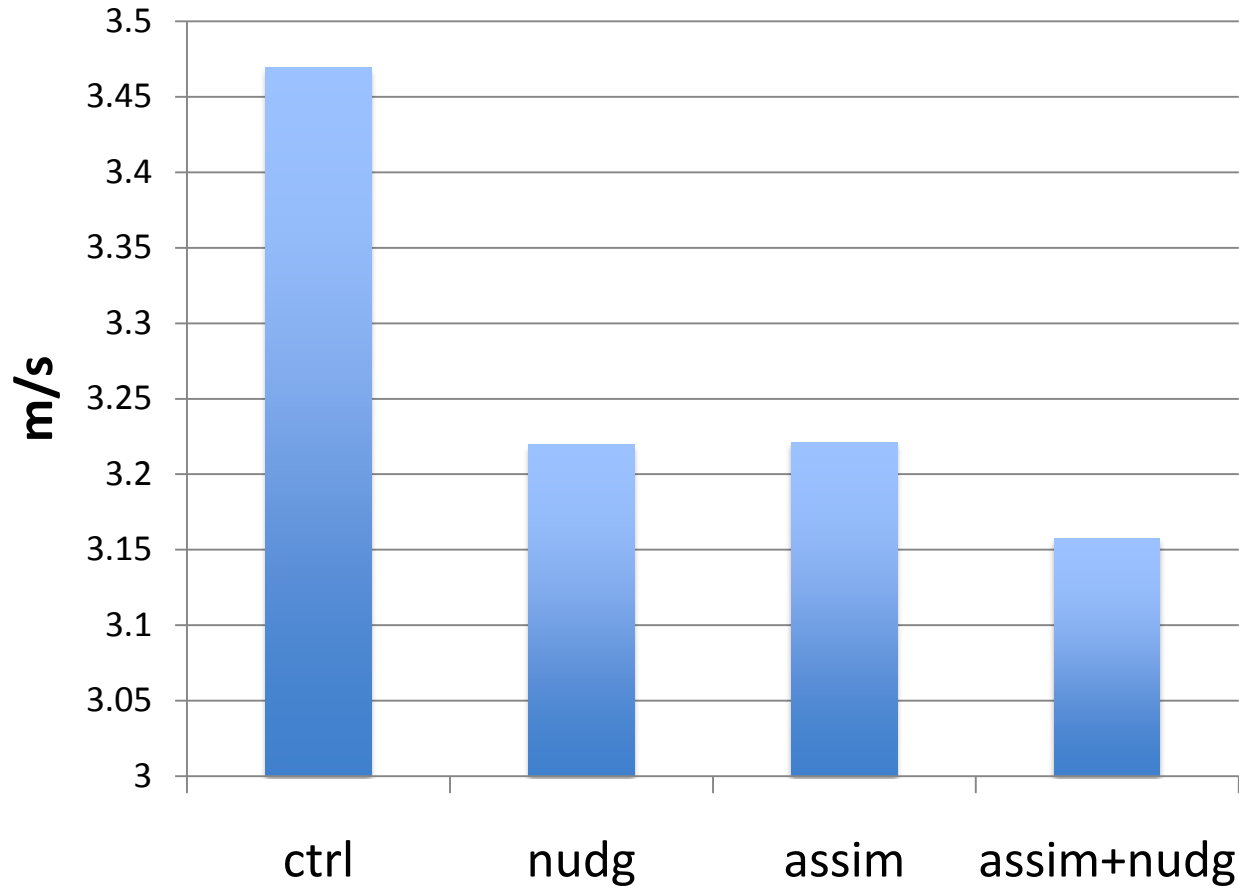


Analysis Nudging

- Analysis nudging continually guides model solution towards existing reanalysis to prevent errors from growing too large, which is very necessary for a regional climate modeling.
- WRF contains two 3-D analysis nudging options:
 - Gridpoint nudging – each gridpoint is nudged toward input dataset
 - ✓ – Spectral nudging – only nudges coarser scales; better allows smaller-scale model information to be retained

Assimilation + Spectral Nudging

Aug09 case W RMSVE



Final Configuration

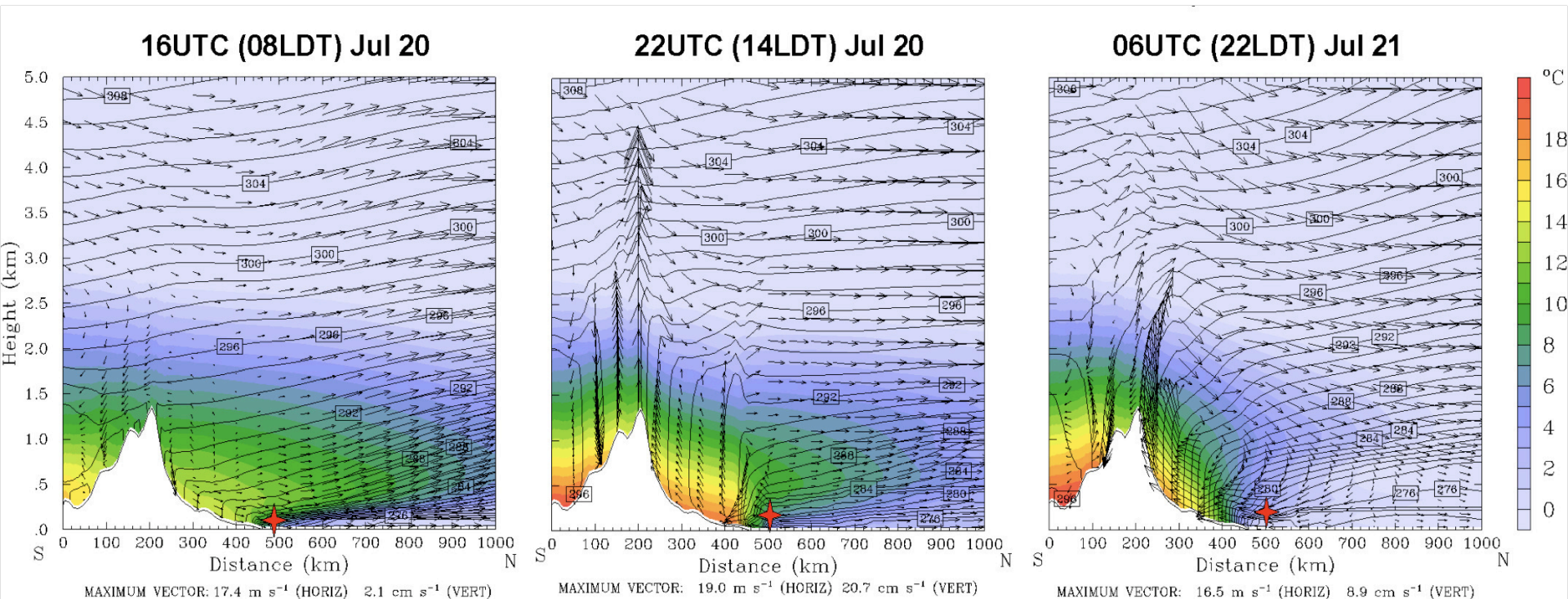
- Data assimilation of:
 - Surface stations
 - Radiosondes
 - QuikSCAT
 - MODIS profiles
 - COSMIC profiles
- Spectral nudging of all vars / levels
- Updated surface condition:
 - CMC snow depth analysis
 - AMSR-E sea ice analysis

Mesoscale Modeling – sea breeze

Unique environmental conditions:

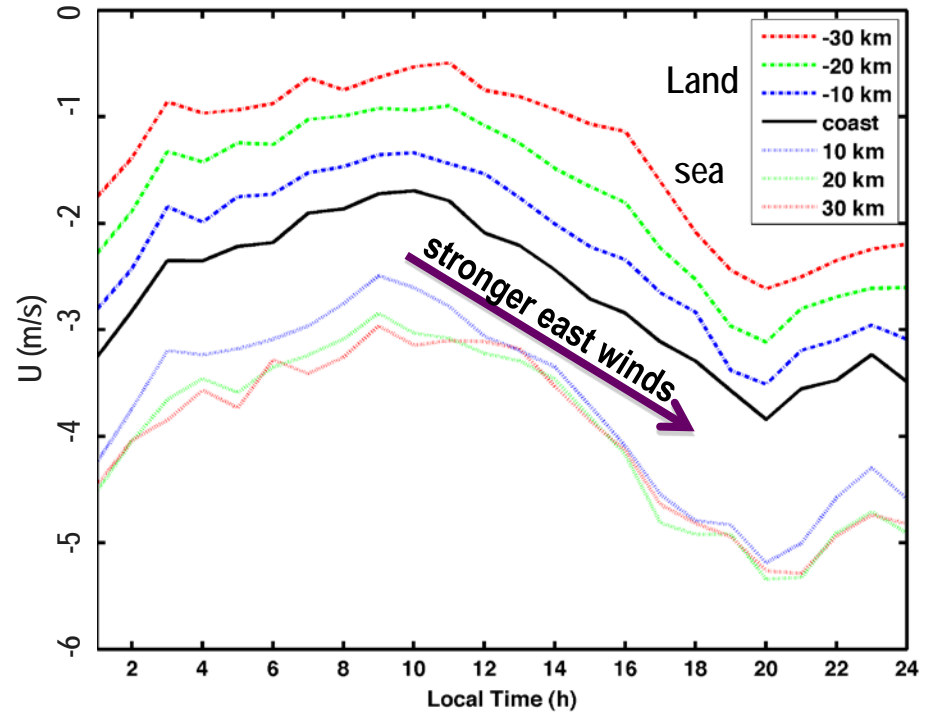
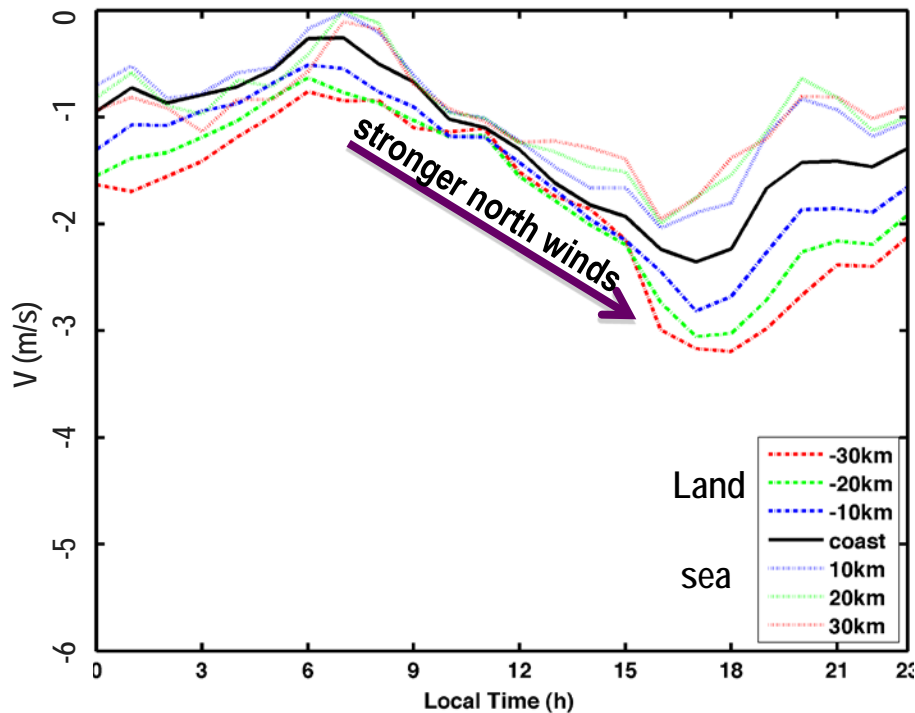
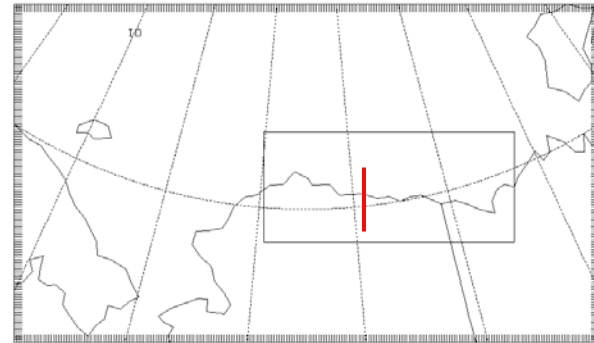
- Continuous solar radiance
- Greater Coriolis effect

Strongest sea breeze (+mountain breeze) occurs late evening due to continuous solar radiance



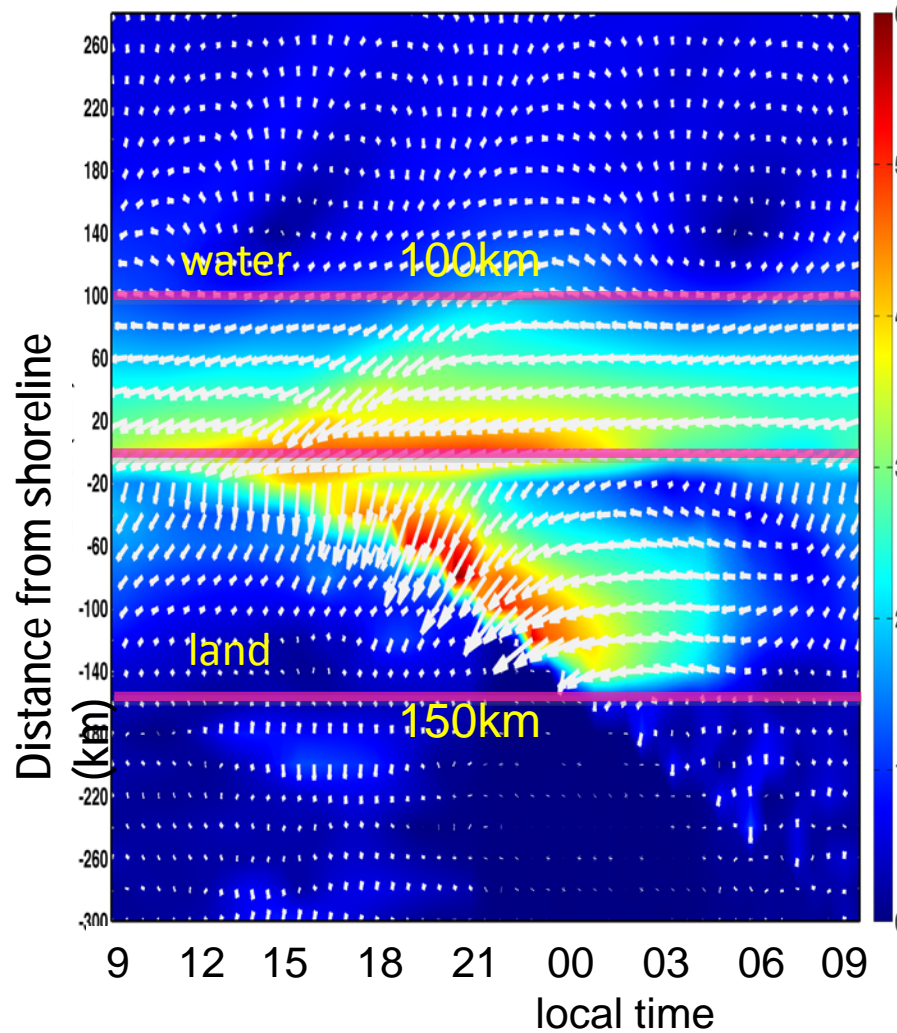
Mean diurnal variation of V and U

Wind fields along a north-south cross section were analyzed to show the averaged diurnal variation of U and V at different distance from the shoreline, negative values signify north and east winds



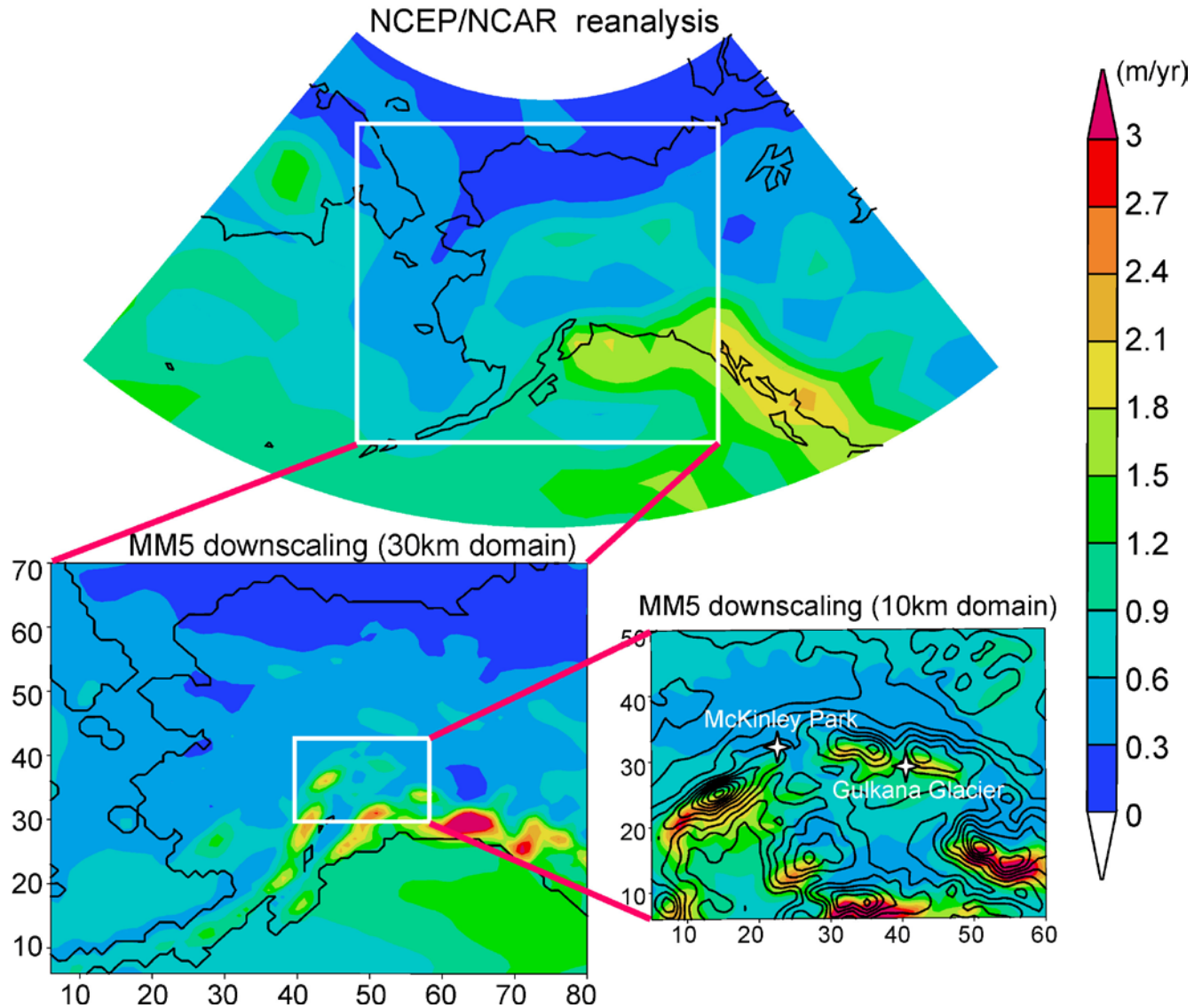
Sea Breeze Diurnal Variation in Idealized Simulation

latitude = 70°N



Beyond Wind Field Reanalysis

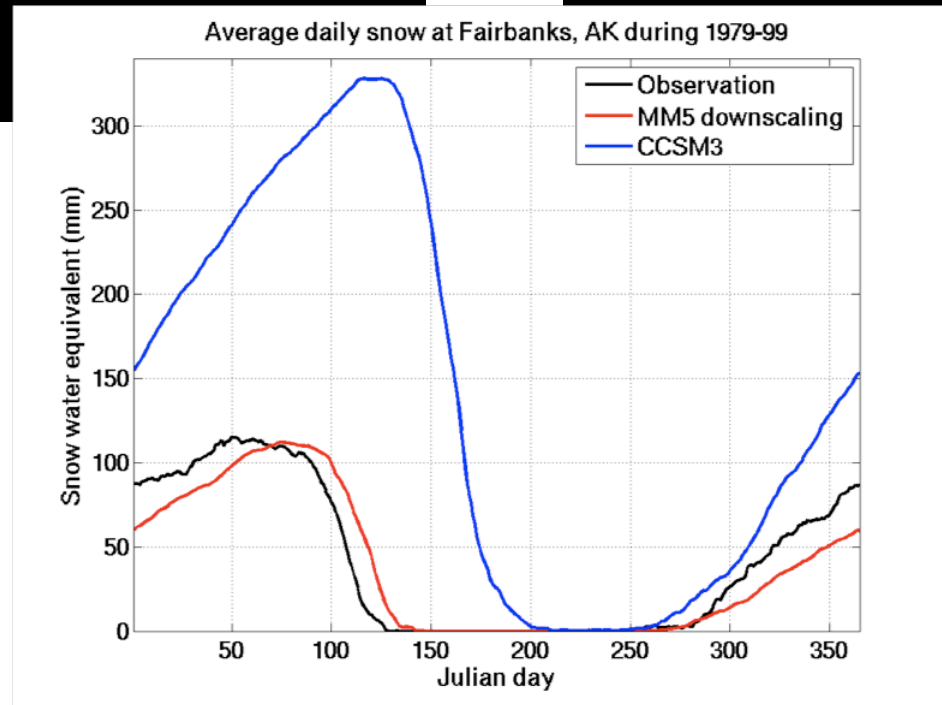
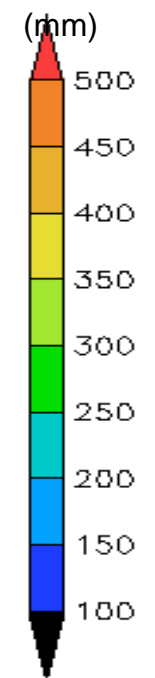
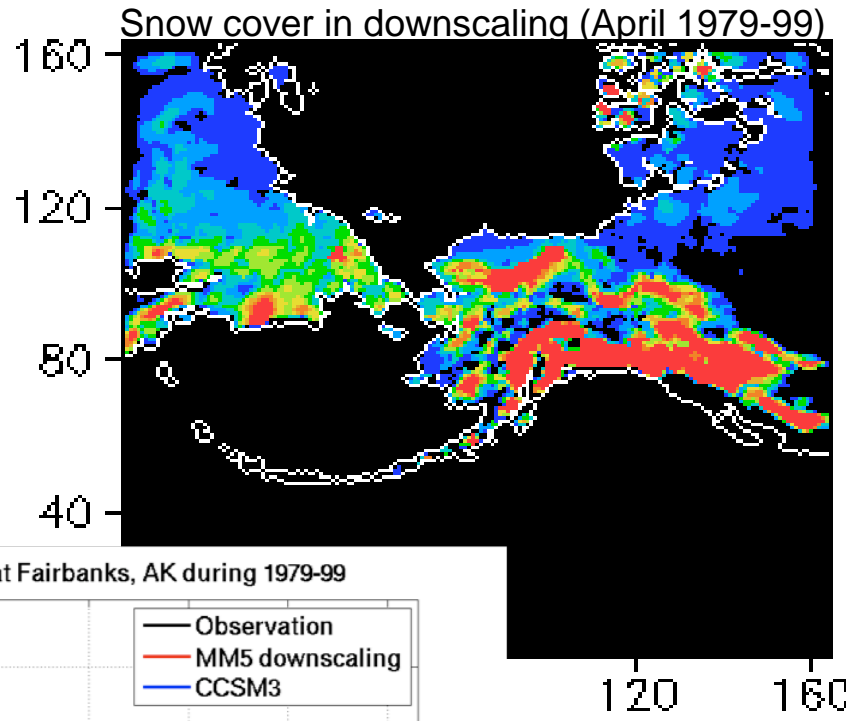
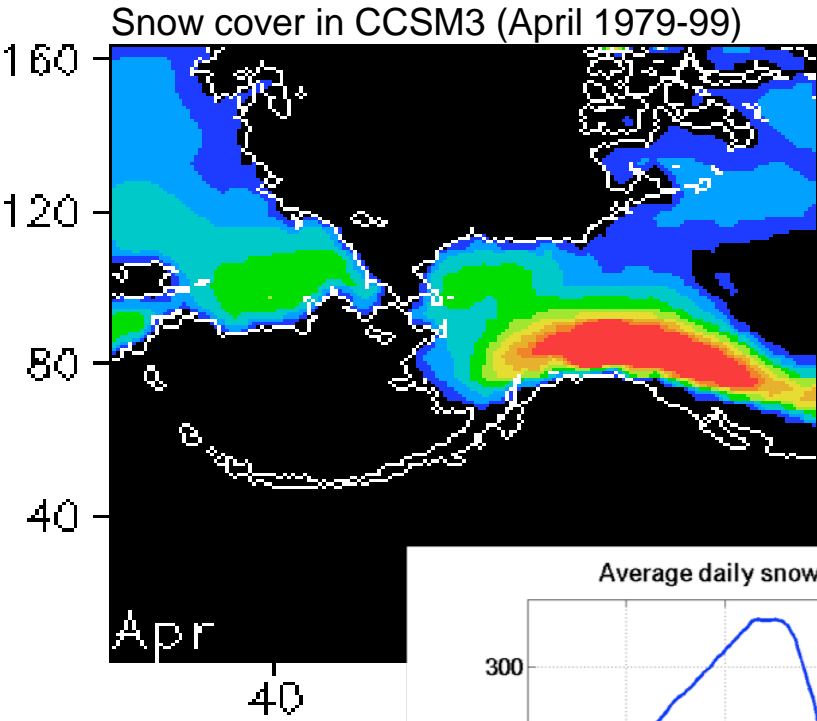
Why downscaling, not just GCM?



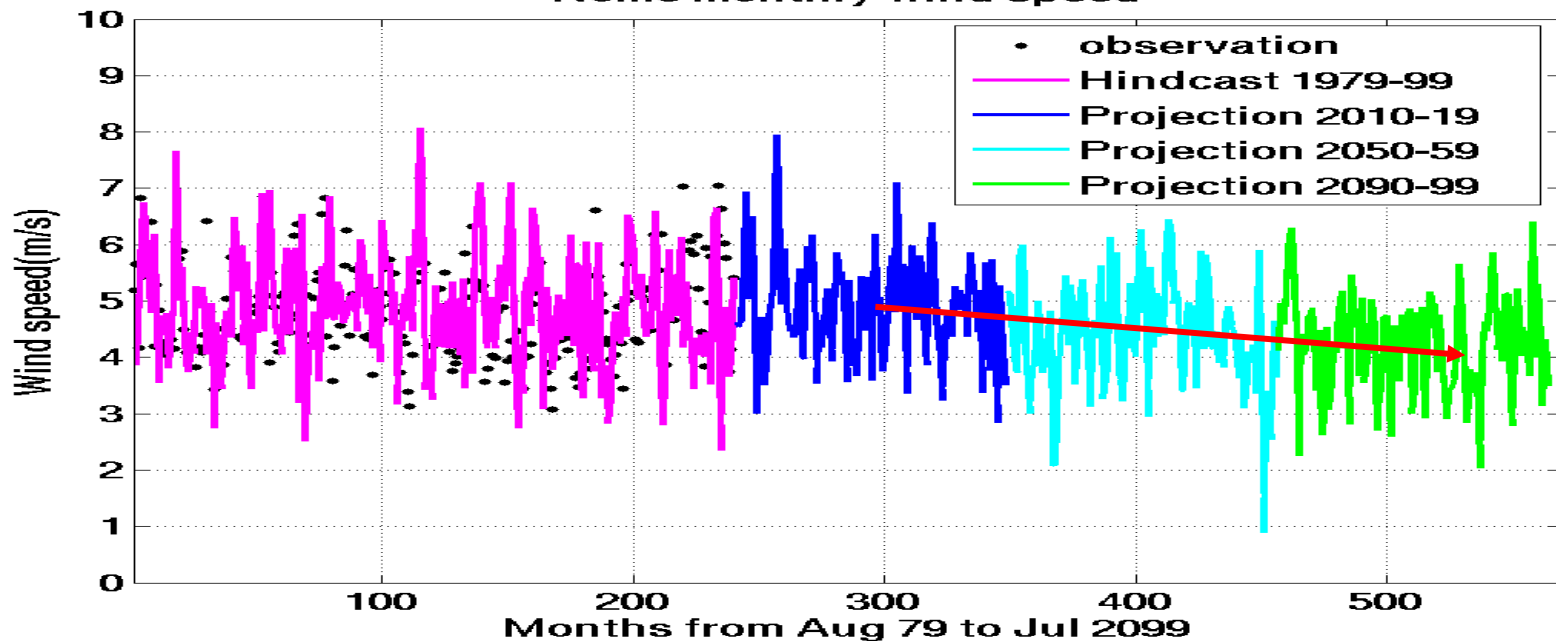
10-yr (Oct.94-Sep.04) annual mean precip.(m/yr)

(Zhang et al 2007, GRL)

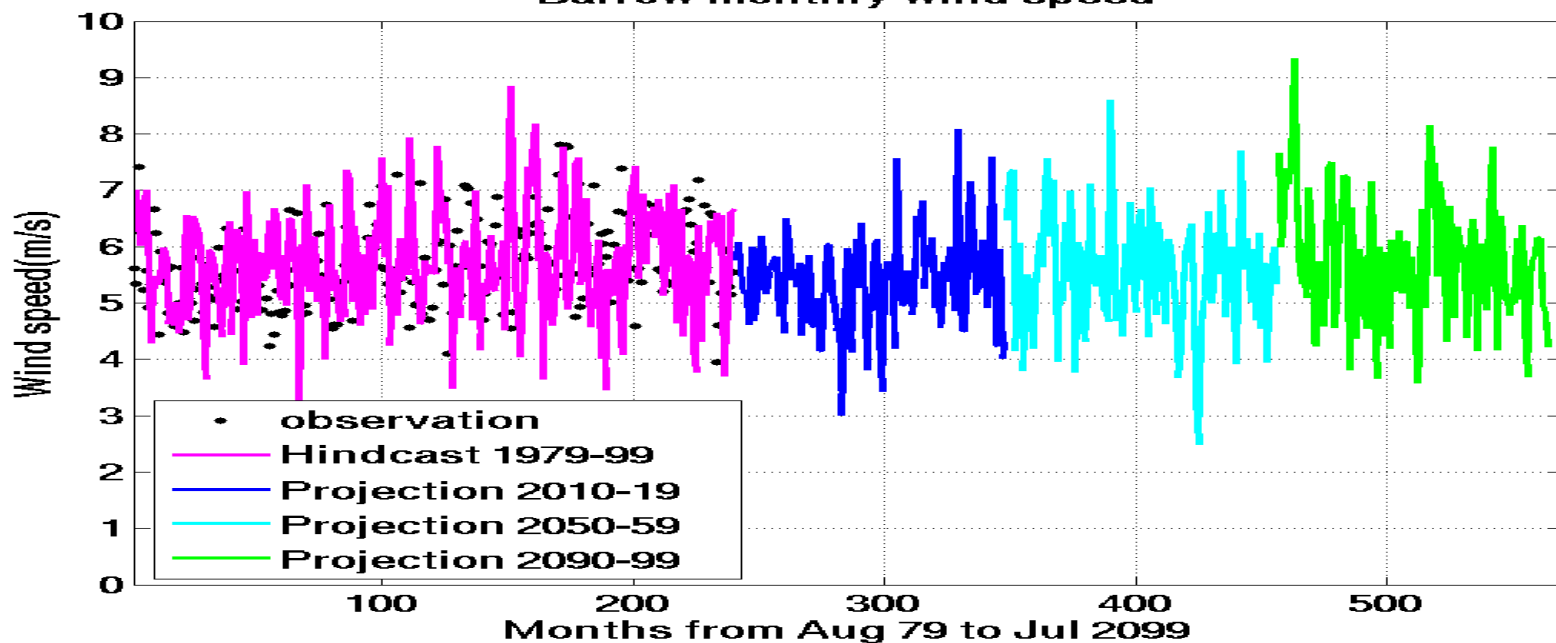
Snow Cover in CCSM3 (left) and Downscaling (right)



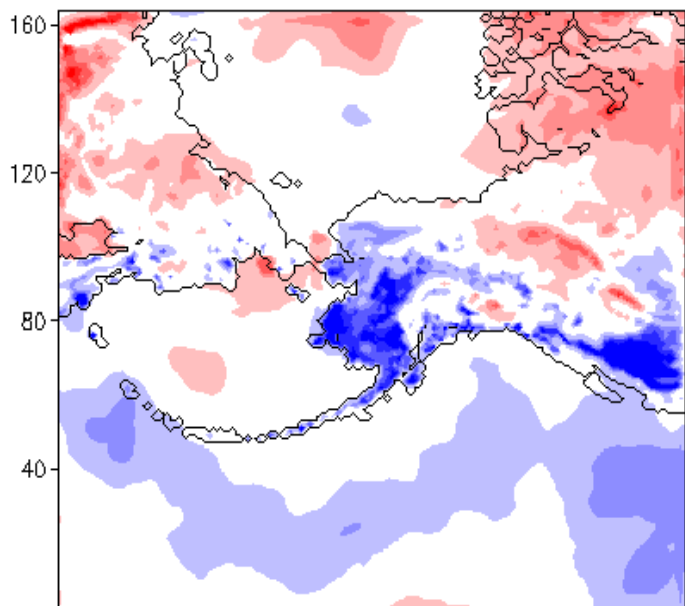
Nome monthly wind speed



Barrow monthly wind speed

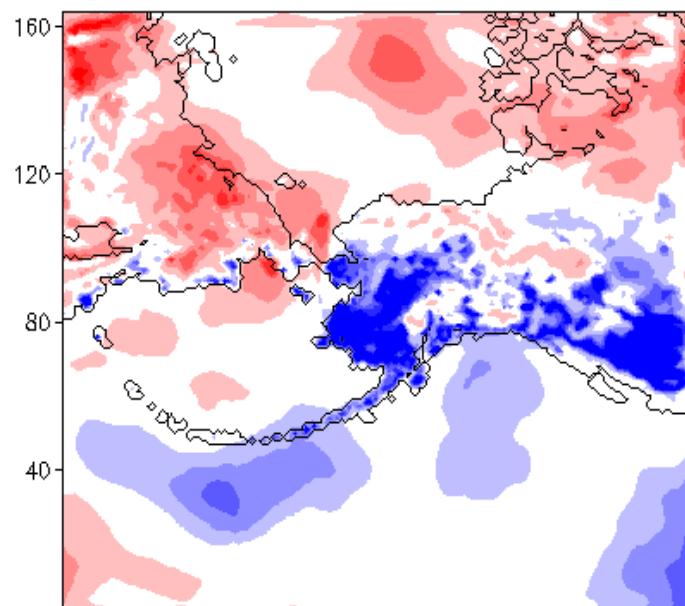


DJF wind changes in 2050s

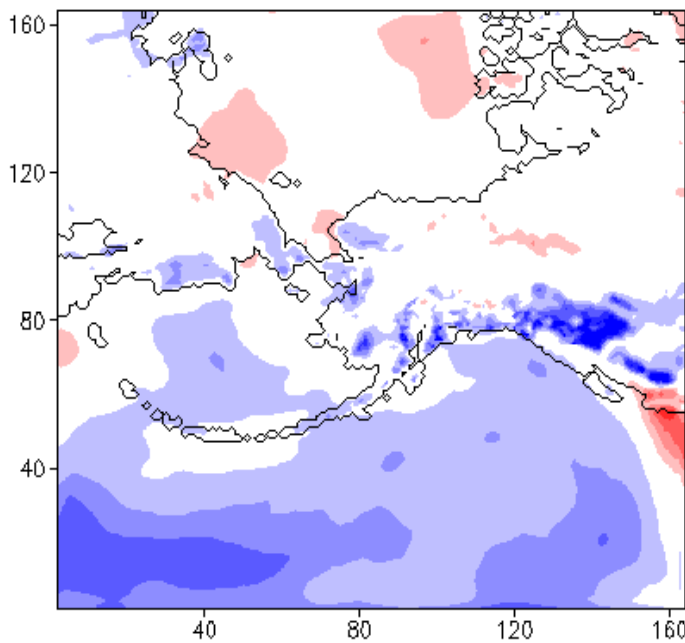


DJF

DJF wind changes in 2090s

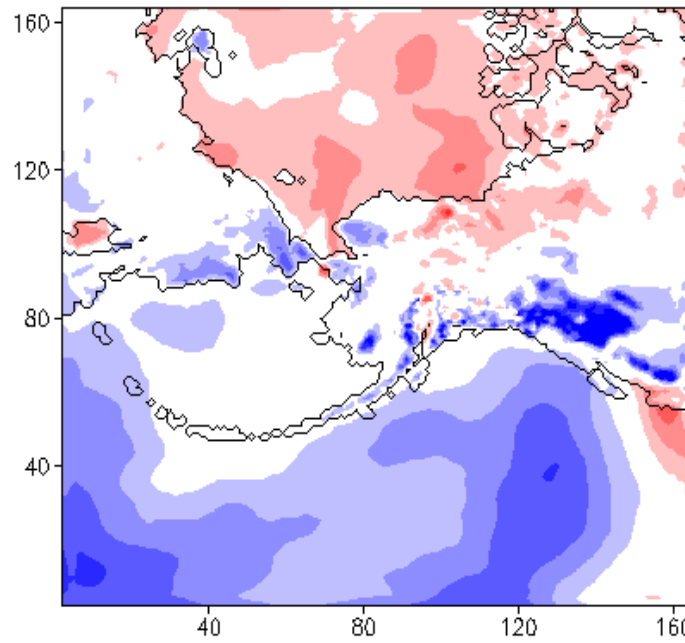


JJA wind changes in 2050s



JJA

JJA wind changes in 2090s



(m/s)

1

0.8

0.6

0.4

0.2

-0.2

-0.4

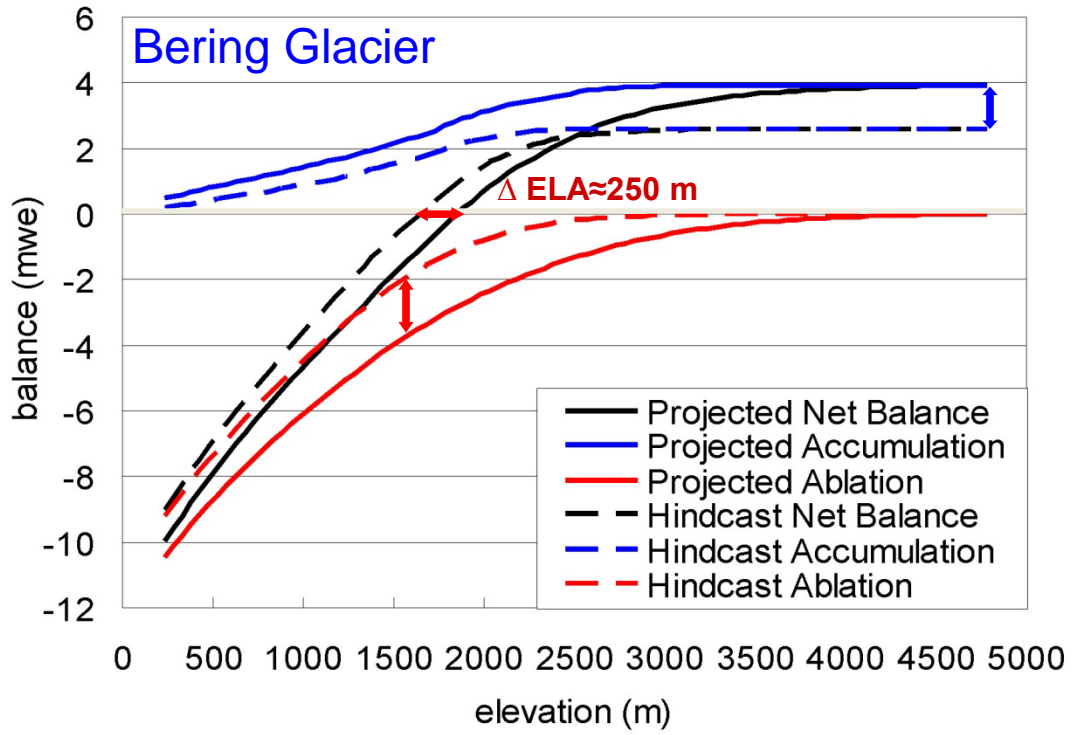
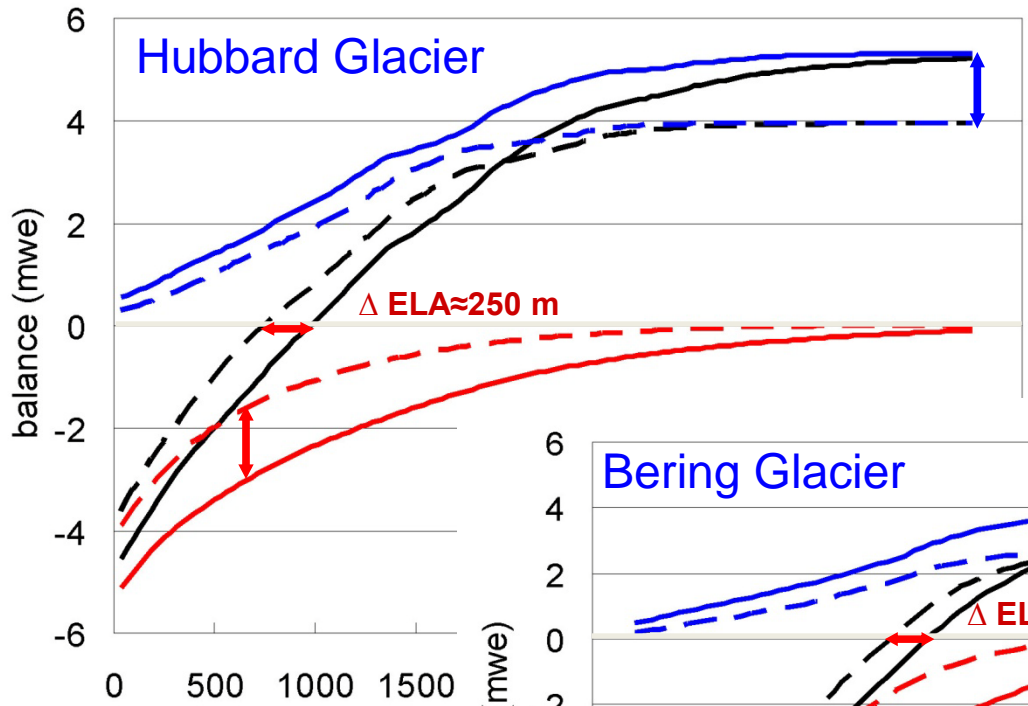
-0.6

-0.8

-1

Projected (2010-19) and Hindcast (1994-04) Glacier Mass Balance

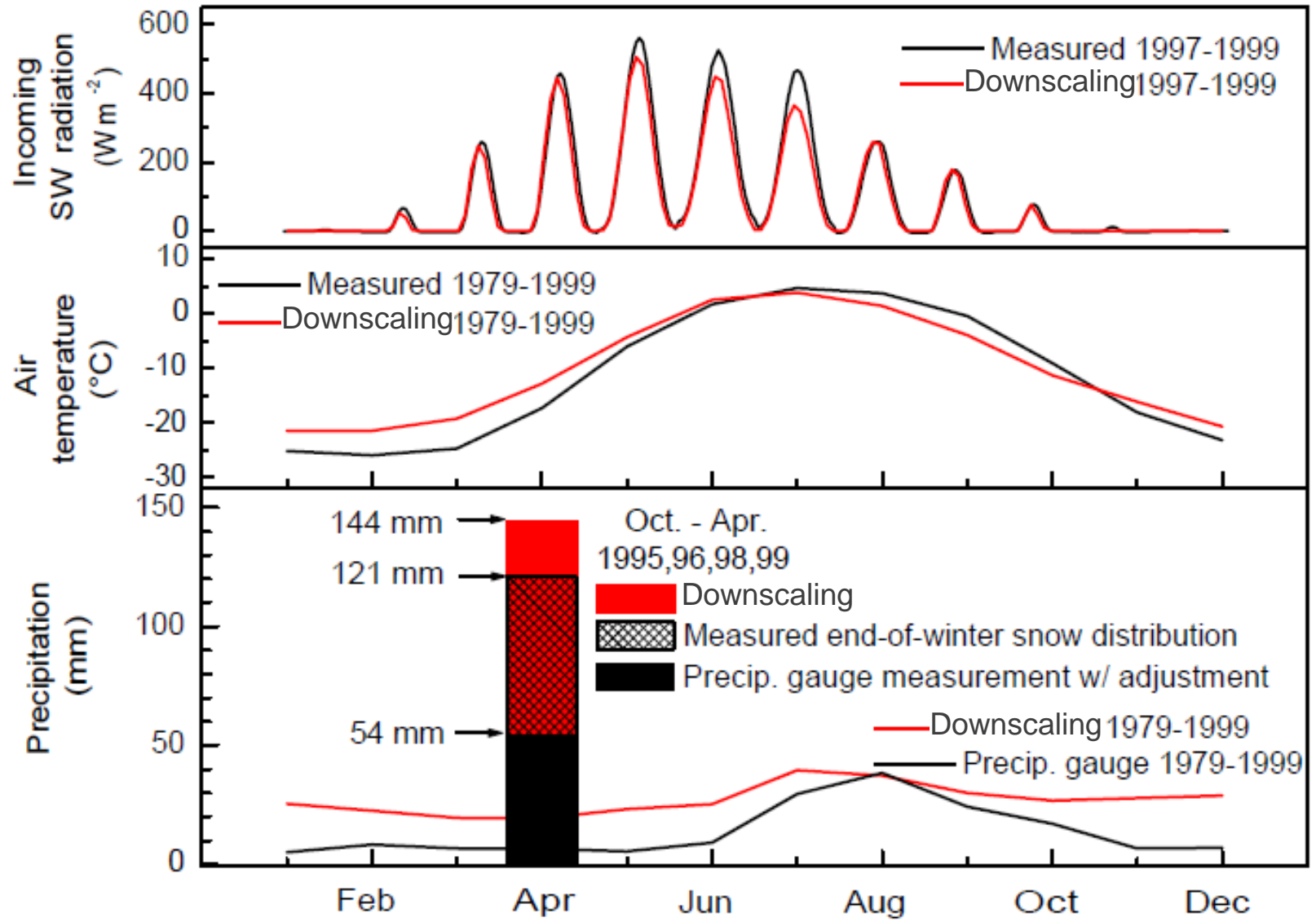
(Zhang et al. 2006)



- Projected Net Balance
- Projected Accumulation
- Projected Ablation
- - Hindcast Net Balance
- - Hindcast Accumulation
- - Hindcast Ablation

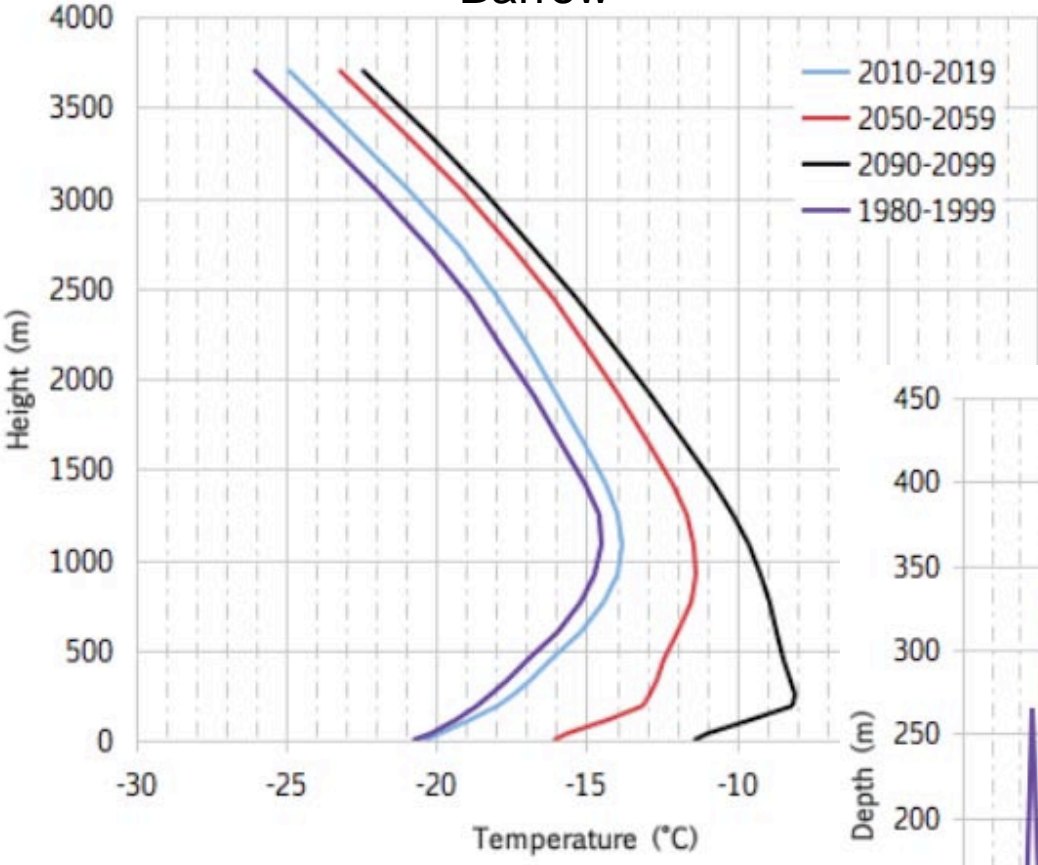
WATER BALANCE OF AN ARCTIC COASTAL WETLAND, BARROW, ALASKA: END-OF 21ST CENTURY PROJECTIONS

Anna K. Liljedahl et al 2011
(in preparation)

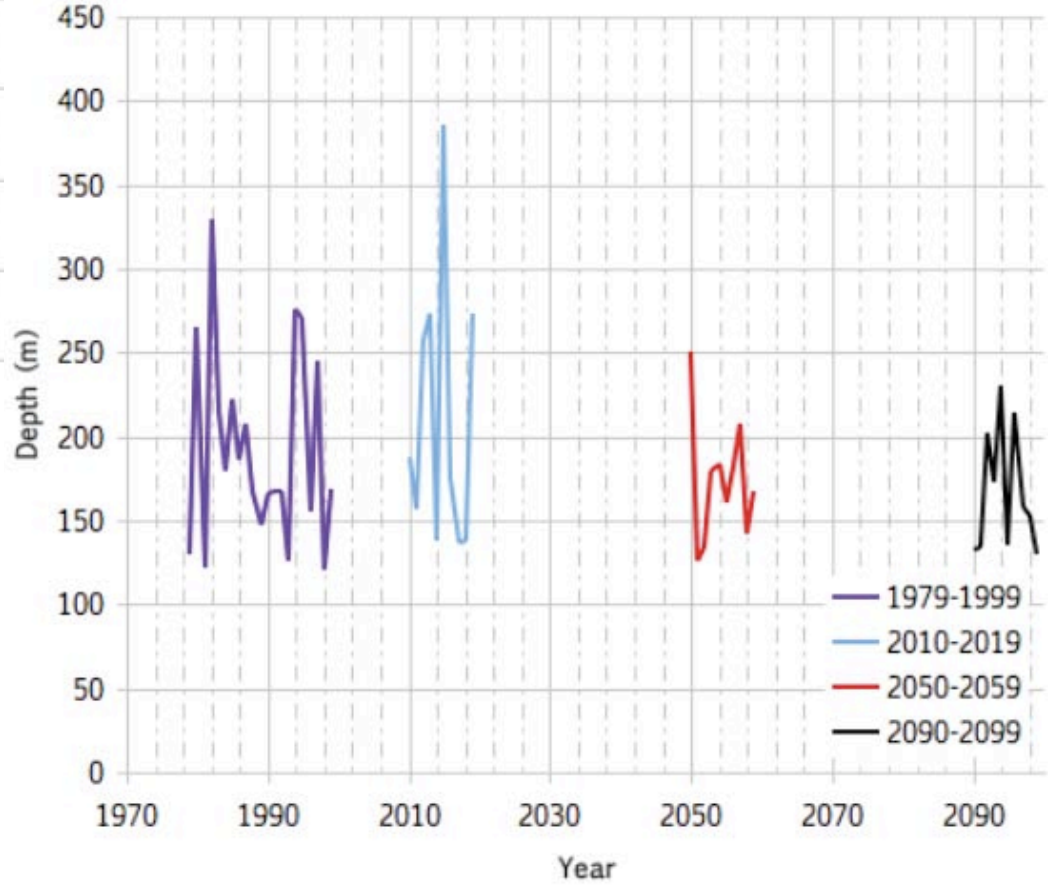


FUTURE INVERSION CLIMATE IN ALASKA DURING 21 CENTURY

Barrow



Fairbanks

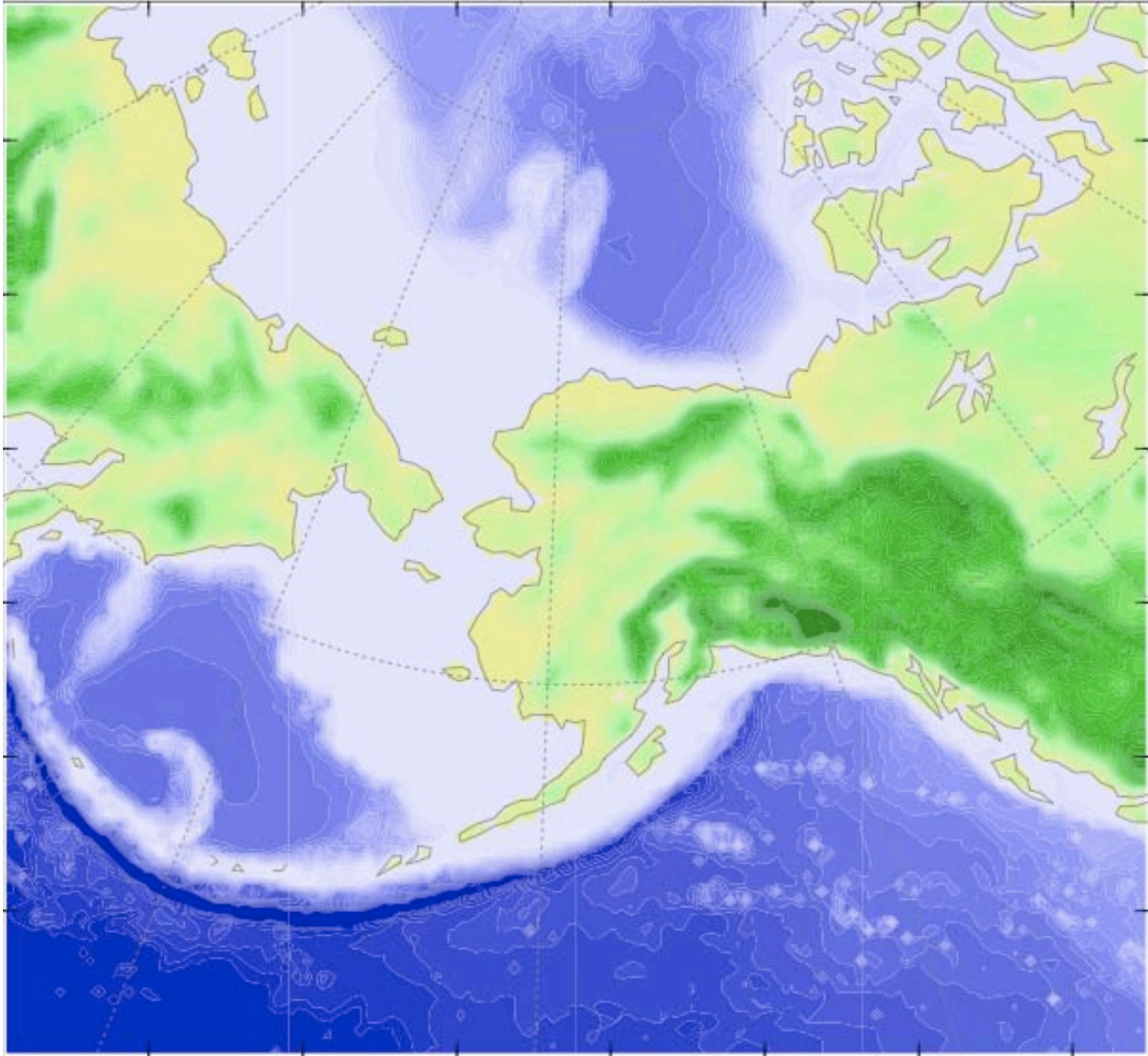


S. Bourne et al. 2009

FUTURE CONTRIBUTION OF GLACIER RUNOFF TO FRESHWATER DISCHARGE INTO GOA

ongoing project...

Downscaling Domain



Questions?
email: jzhang1@ncat.edu